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WATERSHED MODELLING IN THE CHEMUNG
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USING THE US ARMY CORPS OF ENGINEERS
COMPUTER PACKAGE HEC-1

#### A Thesis

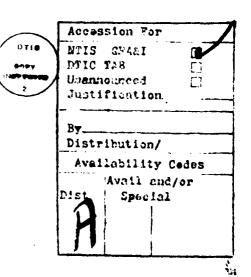
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in Partial Fulfillment of the Requirements for the Degree of

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#### ABSTRACT

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#### BIOGRAPHICAL SKETCH

The author, James C. Styron III, was born in Hobart, Oklahoma on September 29, 1951. He was graduated from the United States Military Academy at West Point, New York in June 1973 and was commissioned as an officer in the U.S.Army at that time. James has attended and graduated from several Army schools including a German course at the Defense Language Institute in Monterey, California, the field artillery officers advanced course at Fort Sill, Oklahoma, and Airborne and Ranger courses in Georgia.

James is an Army Captain who will be working with the U.S.Army Corps of Engineers upon leaving Cornell.

### DEDICATION

To my wife and partner

#### **ACKNOWLEDGEMENTS**

I would like to thank Professor Wilfried H. Brutsaert, my Chairman and thesis advisor. I thank Professor Thomas D. O'Rourke for serving on my Special Committee. A special thanks to Steve Sather, Cornell Computer Services, for his time and effort in making the program compatable with the Cornell IBM computer.

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# CHAPTER ONE INTRODUCTION

Floods are headline grabbers. Every year floods show us how destructively nature can unleash its power. They receive extensive news coverage which describes the resulting damage and loss of life. Floods disrupt lives and productive endeavors. A great deal of effort has gone into studies of past floods and estimates of future floods in order to develop plans for flood damage reduction and prevention. Richards (1955) wrote, "The flood problem is one which affects many branches of engineering other than those dealing primary with the control of water." He also wrote that the flood estimation problem is a peculiarly difficult and complicated one. Although the natural laws governing floods are both recognizable and generally appreciated, the difficulty lies in the application.

The flood factors which are normally considered the most important are as follows: rain intensity, rainfall distribution, rainfall duration, drainage area, basin shape, basin slope, land cover, and, finally, the initial state of wetness in the soil (Richards, 1955, p. 8).

Hoyt and Langbein (1955) recognized the need for a national policy addressing the flood problem. They hypothesized that some future Congress or Congresses would recog-

nize that "a sound approach to floods must be broader than flood protection alone,...". In their ten point proposal number three is quoted, "Flood-forecasting services should be recognized as an essential part of flood management and be appropriately supported." Indeed, around 1955, the U.S. Weather Bureau (now the National Weather Service, NOAA) received appropriations for its flood-forecasting service averaging about \$900,000/year. The savings accredited to this service were roughly estimated at about \$27,000,000, implying a cost-benefit ratio of 1 to 30.

Flood-forecasting is certainly not a modern notion. Diodorus Siculus visited Egypt during the 1st Century B.C. (Biswas, 1967, p.125). On the Nile River, the stage of the Nile were maintained by using nilometers (gaging stations). These structures were used to observe the river stage and compare the current height with that of previous high water marks. Little evidence of actual technical co-operation between the temples charged with running the nilometers exists but according to Diodurus, flood warnings were sounded to the population from the nilometer at Memphis (the seat of the government) in case of emergency. The ancients' warning system worked as follows; with the approach of the flood season, the stage of the Nile was carefully watched and compared with the markings of the previous year. Swift rowers were sent from the furtherest upstream gaging stations, one after another, to report the latest

level at the capital. These extremely good rowers, rowing with the current, were able to outpace the approaching peak flood and give sufficient advance warning to the people along the river of the forthcoming catastrophe.

No one would seriously consider instituting a flood prediction system today based on a fleet of rowboats. However, there is a need to forecast floods in a "real-time" setting in order for warnings to be broadcasted, people and property to be evacuated, and emergency resources to be mobilized in the area of the expected threat.

But how does one go about forecasting a flood? This thesis will present only one method. That method is the unit hydrograph approach for ungaged watersheds which was suggested by the U.S. Army Corps of Engineers in its manuals and which will be presented later in Chapter Three. The selection of this method was based in part on the observation of Dooge (1959) who wrote, "The unit hydrograph approach to stream flow has developed into one of the most powerful tools of applied hydrology."

#### GOAL AND SCOPE OF THESIS

The goal of this thesis is to use the U.S. Army
Corps of Engineer's HEC-1 computer program (discussed in
Chapter Three) in developing, (a) Regression relationships
between the unit hydrograph parameters (based upon the
Clark Method) and some of the more important physical

basin characteristics; and, (b) Regional relationships to determine the initial loss and constant loss rate parameters in the area of interest.

The scope of this thesis is to accomplish the goal by using historical storm and rainfall data from several gaged streams in the area of interest in solving the inverse problem. The inverse problem is defined as a problem in which the input (rainfall) and the output (stream runoff) are known and are used in determining the transfer functions, i.e. the Clark unit hydrograph and loss rate parameters.

The area of interest is that of the Chemung and Upper Susquehanna River basins in the state of New York and a small portion of Pennsylvania. Nine gaged stream basins in this area were selected for use in calibrating the computer model HEC-1. These basins which are unregulated were chosen on the basis of their hydrological similarity. The ungaged watersheds in the area around Corning and Elmira, New York, are the main targets of this modelling effort.

#### **OVERVIEW**

This work is organized into six chapters. Chapter
One is the introduction in which the consequences of floods,
the need for flood prediction, and the goal and scope of
this thesis were presented. Chapter Two is a literature
review primarily concerned with the origins of the HEC-1

approach selected for this study. Chapter Three describes the specific method employed to obtain the results. Chapter Four presents the data and calculations performed in the application of the method in the area of interest. Chapter Five is a discussion of the results. Chapter Six presents conclusions reached by this author and other authors. Several appendices are provided to keep the basic chapters as free as possible from "data overload" and are identified in the basic work.

#### CHAPTER TWO

#### LITERATURE REVIEW

#### PURPOSE AND SCOPE

The purpose of this chapter is to trace the development of the unit hydrograph theory and its application in flood prediction on ungaged basins. The scope of this chapter is confined to a brief discussion of the origins of the three components of the HEC-1 model with a limited presentation of other methods.

#### LOSSES

The term precipitation loss is defined in this thesis as the amount of precipitation not available during a storm for overland flow. There are basically two types of precipitation losses. One is called interception. It represents the amount of rainfall which is held by the trees, grass and local depressions such as potholes and cracks. Anything that traps and does not allow the rainfall to move about freely is classified under this type of loss. The other type of precipitation loss is infiltration. It represents the amount of rainfall absorbed into the earth.

In computing rainfall excess, the precipitation losses are subtracted from the total storm rainfall. Sokolov et al. (1976, p.175) listed five methods to determine rainfall losses. They are mentioned below:

- 1. Constant runoff coefficient, which is the ratio of surface-runoff depth to total rainfall depth.
- 2. Constant infiltration rate during a storm known as the  $\bar{\Phi}$ -index in the U.S.
- 3. Initial rainfall loss followed by a constant infiltration rate. In this method all rainfall is considered to be lost until an initial rainfall-loss demand has been satisfied. Thereafter, rainfall is lost at a constant rate.
- 4. Regional curves of infiltration capacity, that are variable with time, and are based on soil type, land use, and an antecedent-rainfall index.
- 5. A variable runoff coefficient  $Q_t$  method, which varies with time, even during storm periods, in response to the variation of a basin moisture index.

Any one of these methods may be included in a basin model. HEC-1 uses the methods described in 3. and 4. above and will be discussed further in Chapter Three.

#### UNIT HYDROGRAPH

#### ORIGINAL UNIT HYDROGRAPH THEORY

Sherman (1932, p. 501) first proposed the unit-graph (as it was originally called) method as a way to predict flood-peak discharges and discharge hydrographs from rainfall events. Sherman defined the unit-graph as follows:

If a given one-day rainfall produces a 1-inch depth of rainfall over the given drainage area, the hydrograph showing the rates at which the runoff occurred can be considered a unit-graph for the watershed.

The unit hydrograph (unit-graph) theory makes two
assumptions: 1) Runoff response to rainfall is time invariant. 2) Runoff response to rainfall is linear; this
implies the principle of superposition. These assumptions
are not rigorous and there is some flexibility.

By means of the second assumption, Sherman observed the possibility of computing the runoff history corresponding to a rainfall of any duration or degree of intensity for the same watershed from a single observed hydrograph. In applying his method, he noted, that the ordinate and time intervals of unit hydrographs for two similar watersheds of different sizes are proportional to the square root of the watershed areas, provided the difference in size is not great. He also observed that consistent results in the data were based upon the observations being confined to a single area or to closely similar areas and being segregated accoding to the seasons.

Unit hydrograph methods are usually based on the ability to separate the surface-runoff hydrograph from the hydrograph of the total rainfall. The total hydrograph is often assumed to consist of three types of flow defined below by Dooge (1959, p. 241): 1) Surface runoff, or the water reaching surface channels by the overland route; 2) Interflow, or the portion of infiltrated water that passes through the shallower horizons of the soil to reach defined stream channels within a relatively short time, without first

reaching the main ground-water table; and, 3) Base flow, or the water contributed as ground water outflow from an aquifer. It should be noted that interflow is usually combined with surface runoff and collectively called surface runoff.

#### LATER UNIT HYDROGRAPH DEVELOPMENTS

In 1934, 1936 and 1937 Zoch (Dooge, 1959, p.241) published several papers on the unit hydrograph method based on the assumption that at any time the rate of discharge is proportional to the amount of rainfall remaining with the soil at that time. The S-hydrograph method developed by Morgan and Hullinhors (1939) employed a unit hydrograph of a certain duration to form an S-hydrograph resulting from a continuous applied rainfall to construct unit hydrographs for other than the original duration.

Snyder (1938, p.447) and Taylor and Schwartz (1952, p. 235) applied unit hydrograph theory to basins by relating features of the unit hydrograph to watershed characteristics by strictly statistical considerations. Their attempts led to the "synthetic unit hydrograph method."

In contrast to this statistical approach, several workers attempted to give a more conceptual basis to the unit hydrograph. Clark (1945, p. 1419) developed the instantaneous unit hydrograph (IUH) method by suggesting the unit hydrograph for instantaneous rainfall could be derived by routing

the time-area-concentration curve through a single element of linear reservoir storage. O'Kelly (1955, p. 365) of the Irish Office of Public Works suggested replacement of the time-area-concentration curve by an isosceles triangle.

Nash (1957, p. 114) noted that the instantaneous unit hydrograph could be derived by routing the instantaneous rainfall through a series of successive linear reserviors of equal delay time. Dooge (1959, p. 241) attempted to remove many of the subjective elements from unit hydrograph analysis. The number of IUH investigations goes on and on. Chow (1964, p. 14-25) compiled an extensive list of past accomplishments in his handbook. He concluded that the use of the IUH rather than other unit hydrograph methods is better suited for investigations on the rainfall and runoff relationship in drainage basins.

#### CLARK'S METHOD OF UNIT HYDROGRAPH ANALYSIS

Clark (1945, p. 1422) was the first to adapt the idea of the IUH in hydrograph analysis to the unit hydrograph theory. In his original work he wrote:

"The range of possible unit-graph determinations can be reduced by correlating recognized discharge concepts with the physical limitations of valley storage and with the time elements which necessarily result from storage and discharge relations. The correlation can be utilized to develop, from time-area-concentration curves of specific drainage areas, unit hydrographs with a small range of determination variability, independent of assumptions regarding runoff distribution in the flood-producing storm and reflecting influences of drainage area shape and stream pattern."

He showed the relationship between the unit hydrograph and the methods of flood routing (means of modifying a hydrograph by the effects of valley storage). Clark stated that one advantage of using an "instantaneous hydrograph is that it can be derived from the fundamental characteristics of the basin and then used with any length of unit period to determine the unit hydrograph, ...". Other advantages stated by Clark which make his method appealing to "real-time" flood forecasting are: 1) The procedure is definable. Identical hydrographs will be obtained by different people. 2) Determination is independent of any knowledge of runoff distribution, except the time of ending.

3) The unit hydrograph quantities are instantaneous rates of discharge at the time specified.

#### FLOOD ESTIMATION: UNIT HYDROGRAPH

Richards (1955, p.10) presented the principal factors affecting floods. A number of flood formulas is discussed and a brief outline of the flood frequency, probability, and unit hydrograph methods used in the U.S. is included. Sokolov (1976, p.182) states that flood flow computations for unit hydrograph derivation are accomplished using three general methods: 1) Analysis of rainfall--runoff records for isolated unit storms, 2) Analysis of rainfall--runoff records for major storms, and 3) Computation of synthetic unit hydrographs from a) Direct analogy with

basins of similar characteristics, or b) Indirect analogy with a large number of other basins through the application of empirical relations.

BASIN CHARACTERISTICS USED IN HYDROGRAPH ANALYSIS
POSSIBLE PARAMETERS

Relating model parameters to physical basin characteristics is the major purpose of this study. The geomorphic and geophysical structures of a basin play an important part of the watershed's response to rainfall. Many basin characteristics have been used in the past to relate a response of precipitation to physical structures in a watershed. Linsley (1982. p.311-316) defines several characteristics of interest in this study. They are: drainage area, stream density, basin shape, channel slope, and channel length.

In a very large study of the Northeastern U.S.,
Langbein (1947, p.125) using U.S. Geological Survey (USGS)
topographic maps of 340 basins withdrainage areas of 1.64
to 7,797 square miles collected data on drainage, area,
length of streams, stream density, land slope, channel
slope, area-altitude distribution, and area of water bodies
of the basins. Characteristics were divided into geographic
such as water bodies, direction of stream flow, latitude
and longitude, and topographic such as basin area, stream

length, area-distance distribution, land slope, basin altitude, and tributary and principal stream slopes. A principal channel was defined as one that drains more than ten percent of the total area while a tributary channel drains less than ten percent of the total area. Their conclusions were that many characteristics have a direct relationship with other characteristics; for example, steep land slopes imply steep channel slopes and stream density tends to vary with the land slope. The study showed that no one element is unique in any one basin and that not all characteristics must be correlated to model parameters to get satisfactory regression relationships.

#### SOME UNIT HYDROGRAPH APPLICATIONS

Taylor and Schwarz (1952, p.235) published a study based on 65 rainfall excess periods over 20 basins in the North and Middle Atlantic States with varying drainage areas of 20 to 1600 square miles. Their results were that the drainage area, length of longest watercourse, length to center of area and the equivalent mainstream slope were the most significant basin characteristics.

Brater's (1940, p.1154) study indicated that the unit hydrograph method is one of the best practical devices for predicting flood flows. His work based on 22 small

watersheds in the Southern Appalachians of varying cover with areas of 4.24 to 1,876.7 acres showed that the unit hydrograph principle "permits the engineer to estimate not merely the peak discharge, but the entire hydrograph of runoff. The peak may be determined for any desired time interval."

Analyzing data on nine streams tributary to the Chemung River in New York, Morgan and Hullinghorst (1939, p. 1) found good correlation to the area of the watershed, mean length of travel and the mean height of the watershed above the outflow.

Unit hydrographs were used in a flood-frequency work by Kinnison and Colby (1945). On about 48 basins in Massachusetts, the characteristics, found to be most influential, were: median altitude (s), drainage area (M) and the average distance which water from runoff uniformly distributed over the basin must travel to the outlet (L). Thus M is a measure of the volume of water to be discharged, s is a measure of the fall and L is a measure of the distance that the water travels to the point of discharge.

More recent works by Rodriguez and Gonzalez (1982, p. 877) and Rodriguez et al.(1982, p. 887) have applied the geomorphoclimatic theory of the IUH. Their works have helped to explain much of the noise observed in relating IUH parameters to basin characteristics without considering the

coupling effects of the climate and geomorphology on a basin.

Hence, it seems that most authors use easily definable basin characteristics such as main channel length, main channel slope and drainage area to obtain their hydrograph response relationships to basin charactersitics.

#### BASE FLOW

The release of water from underground storage into the channel is called base flow. Dooge (1973, p. 89) has presented a thorough discussion of the subject as well as its separation from the total storm hydrograph. Although several methods exist to calculate base flow, the difference between the methods is not significant in relation to the total flow in the channel during a major storm.

## CHAPTER THREE METHOD

#### PURPOSE AND SCOPE

The purpose of this chapter is to describe the method used in determining HEC-1 parameters for the study region's gaged basins and in developing the relationships between unit hydrograph parameters and basin characteristics. The scope of this chapter is to present the HEC-1 options employed and the multiple regression scheme used in this thesis.

#### WHAT IN THE HECK IS HEC-1?

HEC-1 is the shortened name of HEC-1 Flood Hydrograph Package, Computer Program 723-x6-L2010, with its latest revision taking place in September 1981. The program was developed at the Hydrologic Engineering Center (HEC), Davis, California, which is a part of the U.S. Army Corps of Engineers' Water Resources Support Center. The original development of HEC-1 was made in 1967 by Leo Beard and others (HEC-1,'81), with the first version being published in October 1968. Major revisions have been made in 1969,1970, 1973, and 1981. The current version has the major additional capabilities of the dam-break (HEC-1DB), project optimization (HEC-1GS), and the kinematic wave (HEC-1KW) programs included in the package.

The HEC-1 model is designed to simulate the surface runoff response of a river basin to precipitation by representing the basin as an interconnected system of hydrologic and hydraulic components within a portion of the basin, commonly called a subbasin. A component represents a surface runoff entity, a stream channel, or a reservior. Representation of a component requires a set of parameters which specify the particular characteristics of the component and mathematical relations which describe the physical processes. The result of the modelling process is the computation of streamflow hydrographs at desired locations in the river basin.

As previously stated in the Introduction under Scope and Goals, this thesis will use HEC-1 to model gaged basins in order to determine values for parameters of a surface runoff entity component which will be used as a predicting tool for ungaged basins in the region of interest

#### ASSUMPTIONS AND LIMITATIONS

As pointed out in the HEC-1 Users Manual, HEC-1 makes a number of assumptions and has some limitations. The major assumptions are that hydrologic processes such as precipitation and interception/infiltration, may be represented by model parameters which reflect average conditions within a subbasin; model parameters represent temporal as well as spatial averages. Some of its limitations are that simu-

lation is limited to a single storm event (because HEC-1 does not have the ability to account for soil moisture recovery during periods of no precipitation) and that results are given in discharge rates rather than in stage heights.

## HYDROLOGIC ANALYSIS PROCEDURES OF UNGAGED WATERSHEDS USING HEC-1

Guidelines and methods for modelling ungaged basins are presented in a HEC report entitled "Hydrologic Analysis of Ungaged Watersheds Using HEC-1", Training Document No. 15, April 1982. This report formed the methodology used in this thesis. The report gave detailed techniques for estimating and calibration of HEC-1 model parameters including discussions of the runoff transformation and loss rate parameters.

Below is outlined the basic method used in the regional analysis of watershed characteristics of this thesis (which is described in greater detail in the Report);

- 1. Collect precipitation and runoff information for a range of major flood events (usually 5-10 for each gaged basin) in the region. Identify the watershed boundaries on a topographic map where the gaging stations are taken as the basin outlets.
- 2. Run a HEC-1 rainfall-runoff analysis to optimize

the unit hydrograph and loss rate parameters for the gaged drainage areas.

- 3. Correlate the unit hydrograph and loss rate parameters with basin characteristics and develop generalized relationships for these parameters.
- 4. Use the generalized relationships to compute parameters for the ungaged basins by means of measurable basin characteristics.
- 5. Perform a watershed simulation using HEC-1 for several historical storm events for verification of reconstituted runoff peaks and volumes at gaged locations. If unsatisfactory results are obtained, the analysis is repeated after adjusting the parameters.

#### RAINFALL-RUNOFF SIMULATION

The rainfall-runoff simulation involves five processes.

They are precipitation, interception/infiltration, transformation of precipitation excess to basin outflow, baseflow, and flood routing (not used in this study). Each of the first four processes will be discussed below.

#### PRECIPITATION

The precipitation distribution chosen for this study among the possibilities in HEC-1 was the method of weighted precipitation gages. This method determines a basin average rainfall amount for a historical storm and distributes it temporally over the entire basin. This process produces a precipitation hyetograph.

Each gaged basin in the study region was divided into Thiessen polygons (see Linsley (1982, p. 71) for the Thiessen polygon method). With one rain gage per polygon, the area of each polygon covering the basin could be determined using a planimeter. The area of a polygon divided by the total basin drainage area yielded the percent (expressed as a decimal) of the relative weight for each rain gage. Thiessen polygons were drawn for the recording (hourly measurements) and non-recording (daily measurements) rain gages. To obtain the total storm precipitation the relative weights were placed on the PT-PW input cards (input card data/format will be discussed later in Chapter Four) for the recording and non-recording gages which actually contributed rain

to the basin. However, in order to develop the temporal pattern on the basin, the relative weights placed on the PR-PW input cards only used recording rain gage polygons.

The total storm precipitation for a basin was computed as the weighted average of measurements from several gages using the formula:

$$PRCPA = \frac{\sum_{J=1}^{n} PRCPN(J) * WTN(J)}{\sum_{J=1}^{n} WTN(J)}$$

where PRCPA is the basin average total precipitation, PRCPN(J) the total precipitation for gage J, WTN(J) the relative weight for gage J, and n the number of gages.

The temporal pattern for distribution of the stormtotal precipitation is computed as a weighted average of temporal distributions from recording stations using the formula:

$$PRCP(I) = \frac{\sum_{j=1}^{n} PRCPR(I, J) * WTR(J)}{\sum_{j=1}^{n} WTR(J)}$$

where PRCP(I) is the basin-average precipitation for the Ith time interval, PRCPR(I,J) the recording station precipitation for the gage J.

The basin-average precipitation hyetograph is then computed using the temporal pattern, PRCP, to distribute the total rainfall amount, PRCPA, see Figure 3.1.

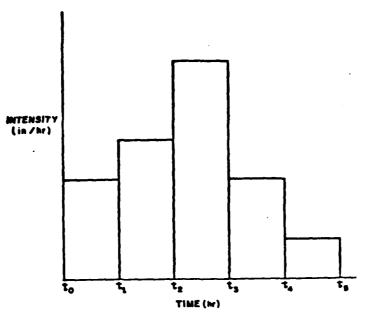


Figure 3.1. Precipitation Hyetograph (HEC-1, 1981). INTERCEPTION/INFILTRATION

Definitions. Interception and depression storage are intended to represent the surface storage of water by trees or grass, local depressions in the ground surface, in cracks and crevices in parking lots or roofs, or in a surface area where water is not free to move as overland flow. Infiltration represents the movement of water to areas beneath the land surface. Both of these are lumped together in HEC-1 and considered as precipitation losses.

Two important factors should be noted about this process: one, precipitation which does not contribute to the runoff process is considered to be lost from the system and two, there is no provision in the model for soil moisture or surface storage recovery. The second factor has already been mentioned under the Assumptions and Limitations

portion of this chapter.

As in the precipitation process, the interception/
infiltration process uniformly distributes the precipitation loss over the entire basin. This process yields
a basin-average for rainfall losses. There is a feature
in the model in which a percentage of the basin may be
labeled as impervious and no losses are calculated for this
portion; however, this feature was not used in this study.

Four different methods may be used in calculating the precipitation loss. By using any one of these methods, an average precipitation loss is determined for a computation interval and subtracted from the rainfall hyetograph. The resulting precipitation excess is used to compute an outflow hydrograph for a subbasin as is shown in Figure 3.2.

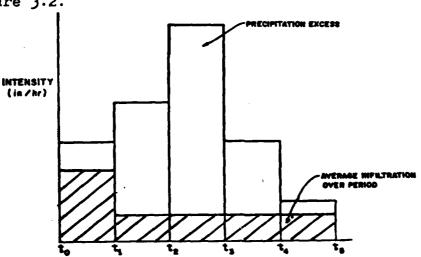


Figure 3.2. Input Hyetograph (HEC-1, 1981).

The four methods available to calculate rainfall

losses are initial and uniform loss rate, exponential loss rate, Soil Conservation Service (SCS) curve number, and the Holtan loss rate. This thesis only considered the initial and uniform loss rate. It will be the only method discussed here.

The initial and uniform loss rate method has two parameters; an initial loss, STRTL (units of depth), and a constant loss rate, CNSTL (units of depth/hour). All rainfall is lost until the volume of initial loss is satisfied. After the initial loss is satisfied, rainfall is lost at the constant rate, CNSTL. These parameter values are entered as input on the LU card.

## TRANSFORMATION OF PRECIPITATION EXCESS TO BASIN OUTFLOW

The transformation of rainfall excess into basin outflow for this thesis is accomplished through the unit hydrograph technique, more specifically, the technique that was proposed by Clark (1945) which has already been discussed in Chapter Two.

The basic methodology of any unit hydrograph component in HEC-1 is the same. The rainfall excess hyetograph is transformed to a subbasin outflow by utilizing the general equation:

N
i

$$Q(i) = \sum_{i=1}^{N} \sum_{j=1}^{i} U(j) * X(i-j+1)$$

where Q(i) is the subbasin outflow at the end of computation-

al interval i, U(j) the jth ordinate of the unit hydrograph, X(i) the average rainfall excess for the computational interval i, and N the number of rainfall ordinates.

Application of the unit hydrograph technique implies two assumptions: one, the unit hydrograph is characteristic of the subbasin and is not storm dependent and two, runoff due to excess from different periods of rainfall excess may be linearly superposed.

The Clark unit hydrograph requires three parameters to calculate a unit hydrograph; the time of concentration for the subbasin, TC (units of time), a storage coefficient, R (units of time), and a time-area curve.

The time of concentration, TC, may be defined conceptually as the amount of time required for a particle of water deposited at the furtherest point upstream in the watershed to travel through the watershed until it passes the subbasin outlet. TC is estimated by the lag time between the end of the runoff producing rainfall to the inflection point on the recession limb of the surface runoff hydrograph.

The storage coefficient, R, is used to account for the effect of subbasin storage on the hydrograph. This parameter is estimated by dividing the flow at the recession inflection point of the surface runoff hydrograph by the rate of change of discharge (slope) at this same time.

See Figure 3.3.

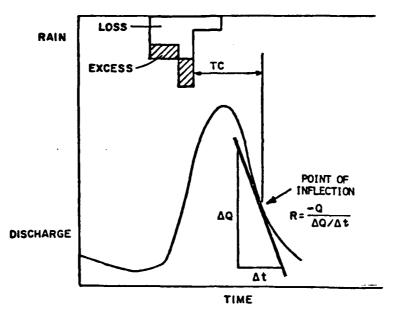


Figure 3.3. Clark UH Parameters (HEC-1, 1981).

The time-area curve defines the cumulative area of the watershed contributing runoff to the subbasin outlet as a function of time (expressed as a portion of TC).

No developed time-area curves for any of the gaged subbasins in the study region could be found, therefore, use of the HEC-1 synthetic dimensionless time-area curve was made. The equations of this curve are:

$$AI = 1.414 \text{ T}^{1.5}$$
  $0 \le T < .5$   
 $1 - AI = 1.414 (1-T)^{1.5}$   $.5 \le T \le 1$ 

where AI is the cumulative area as a fraction of the total subbasin area and T the fraction of the time of concentration. See figure 3.4.

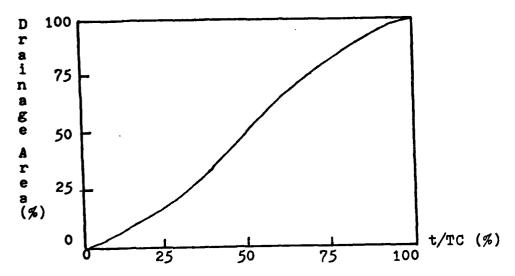


Figure 3.4. Synthetic Time-Area Curve

The ordinates of the time-area curve are converted to volume of runoff per second for unit excess and interpolated to the given time interval. This results in the development of a translation hydrograph which is then routed through a linear reservoir to simulate the storage effects of the subbasin. The routing produces a unit hydrograph for instantaneous excess which is averaged to yield the hydrograph for unit excess occurring in the given time interval (one hour in all cases of this thesis).

The linear reservoir routing is accomplished using the equation

$$Q(2) = CA * I + CB * Q(1)$$

where the routing coefficients are

$$CA = \Delta t/(R + .5\Delta t)$$

$$CB = 1 - CA$$

## QUNGR = .5(Q(1) + Q(2))

and where Q(2) is the instantaneous flow at the end of the period, Q(1) the instantaneous flow at the beginning of the period, I the ordinate of the translation hydrograph, at the computational time interval (one hour in all cases in this study), R as previously defined, and QUNGR the unit hydrograph at the end of the computation interval.

The parameters TC and R are entered as input on the UC card.

## BASEFLOW

Unlike direct surface runoff which is calculated as the total precipitation minus the losses, baseflow is defined as the release of water from subsurface storage. To include baseflow effects the HEC-1 model must be supplied with three input parameters, STRTQ, QRCSN, and RTIOR.

STRTQ is the initial discharge (starting flow) in the stream outlet at the beginning of the storm. It is the flow at the outlet at the beginning of the rainfall hyetograph. It is affected by the long term contribution of groundwater releases in the absence of precipitation and is a function of antecedent conditions.

QRCSN is the flow on the receding limb of the computed hydrograph at which an exponential recession begins.

RTIOR is a user specified exponential decay rate

which is assumed to be characteristic of the subbasin.

It is equal to the ratio of a recession limb flow occurring one hour later. See Figure 3.5.

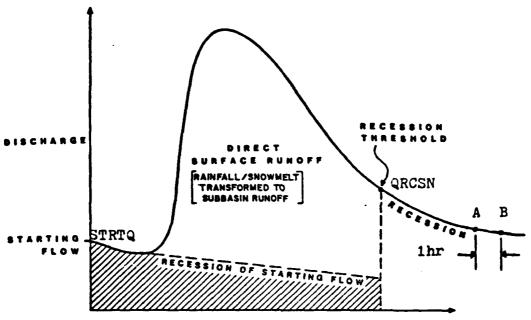


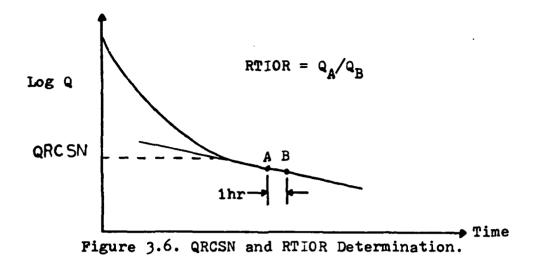
Figure 3.5. Base Flow Parameters (HEC-1, 1981).

Baseflow or recession flow is computed with the equation:

$$Q = Q_0 (RTIOR)^n$$

where Q is STRTQ or QRCSN and n the number of time intervals since recession was initiated.

The values of QRCSN and RTIOR are obtained by plotting the observed flows (starting at the peak) versus time on semi-log paper. The point at which the flows begin to become a straight line defines QRCSN. See Figure 3.6.



RTIOR is usually obtained by taking the average ratio of Q<sub>a</sub> and Q<sub>b</sub> (flows one hour apart) of several, usually about ten, along the straight line portions of Figure 3.6 In this study a subbasin average is determined by taking the average RTIOR for each storm on the basin and then averaging these values. The subbasin average was then held constant throughout all calculations for that subbasin while using HEC-1.

The computed flood hydrograph is adjusted to include the baseflow. At the beginning of a storm the baseflow is computed starting at STRTQ and decays according to the formula above until the computed flood flow on the falling limb equals the value of QRCSN. From then on, the flood flow is calculated using the base flow equation as the computed flood flow unless the computed flow rises above

the recession flow. This would occur in the case of a double-peaked hydrograph such as in Figure 3.7.

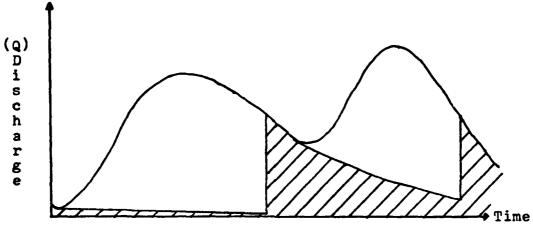


Figure 3.7. Double-Peaked Recession Flow

#### CALIBRATION OF MODEL PARAMETERS

As with all numerical models, the HEC-1 rainfallrunoff model for single event simulation requires data
for calibration in order to be a predictive tool. The
parameters of the numerical model are determined by using
observed precipitation and runoff data to solve the
inverse problem; in other words, given the system input
(precipitation) and system output (runoff), the inverse
problem consists of defining the characteristics of a
system, that produces the transformation from input to
output. HEC-1 has two approaches to transform rainfall
values to runoff. They are the unit hydrograph and
kinematic wave approaches. The unit hydrograph is the

is the most commonly used and was the approach selected for this study. The unit hydrograph approach assumes that a single unit hydrograph is appropriate for all magnitudes of rainfall excess. HEC-1 has three unit hydrograph methods available in its package, those of Clark, Snyder, and the Soil Conservation Service (SCS). The U.S. Army Corps of Engineers uses the Clark unit hydrograph technique most frequently and for this reason was the technique selected for this thesis.

HEC-1 has the capability of automatically producing its "best" estimate of selected model parameters through its optimization subroutine. This subroutine selects those values of the parameters, which yield the "best" reproduction of some measured runoff event with the available measured precipitation data and the selected modelling approach. This automatic calibration approach is accomplished through an objective function, defined as;

STDER = 
$$\sqrt{\sum_{i=1}^{N}} ((QOBS_i - QCOMP_i)^2 * WT_i)$$

where STDER is the error index, QOBS; the observed runoff hydrograph ordinate for period i, QCOMP; the computed runoff hydrograph ordinate for period i, computed by HEC-1 with current parameter estimates. N is the total number of hydrograph ordinates and WT; a weight for the hydrograph ordinate defined as;

 $WT_i = (QOBS_i + QAVE)/(2 * QAVE)$ 

where QAVE is the average computed discharge.

The STDER calculation for optimization may be viewed graphically as in Figure 3.8.

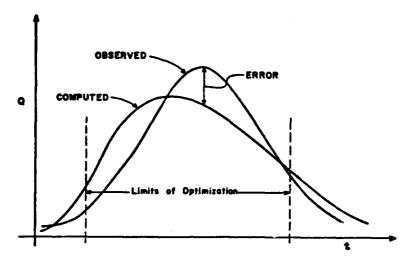


Figure 3.8. Optimization Calculation (HEC-1, 1981)

in which the limits of optimization may be user specified on the OU input card in order to get a better fit on the portion of the curve which is of interest.

The WT<sub>i</sub> term biases the objective function by emphasizing accurate reproduction of peak flow rather than low flow. The objective of this function is to minimize the value of STDER in order to get the "best" fit. The word "best" has been placed in quotation marks to indicate that the values determined through this optimization technique do not guarantee that the values are global solutions; indeed an example of this is shown in Chapter

Pive. A better solution may be found by varying the starting values of the parameters from their default values selected by HEC-1 in its first run. These default values are listed below:

Parameter Name	Initial Value
TC + R	(Drainage Area (mi <sup>2</sup> ))
R/(TC + R)	0.5
STRTL	1.00
CNSTL	0.1

The user also has the option of designating initial values for model optimization as well as designating which parameters are to be held constant during the optimization. Assignment of the initial values is discussed in the HEC-1 Programmer's Manual, October 1981. Constraints are placed on parameter values. These values are bounded by a range of feasible values because of the physical limitations on the values that the various unit hydrograph and loss-rate parameters may have. These constraints are:

TC + R 
$$\geq$$
 1.03  $\Delta$ t/(1-R/(TC+R))  
R  $\geq$  .52 $\Delta$ t

where t is the computation interval which is one hour in all cases in this study.

STRTL = 0

CNSTL 2 0

A complete discussion of the optimization technique as well as the algorithm is presented in the HEC-1 User's Manual. The technique is a univariate search technique in which each parameter value to be optimized is varied one at a time with all other parameters held constant. The "best" value of each parameter is estimated by Newton's method. After each parameter is estimated in turn through four complete cycles, the parameter which most improved the objective function in its last change is adjusted again. This adjustment continues until no one parameter yields a reduction in the objective function of more than one percent. After one more complete search of all parameters, a final adjustment to the computed hydrograph volume to within one percent of the observed hydrograph volume is made. The final objective function value is determined from this final adjustment. An example of the optimization output is shown in Figure 3.9.

Listed first are the initial values of the parameters. The asterisked (\*) values denote which variable was changed and its "optimum" value along with the objective function value with the other parameters.

After the objective function routine is printed, the optimization results are blocked out followed by a statistical summary. If the user specifies optimization limits on the OU card, two summaries are printed. The first contains statistics based on the optimization region

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Figure 3.9. Optimisation Output.

while the second summary is based upon the entire flood hydrograph.

The definition of the terms used in the summary are self-explanatory except for the following four terms:

Standard Error-the root mean squared sum of the difference between observed and computed hydrographs.

Objective Function--the weighted root mean squared sum of the difference between observed and computed hydrographs.

Average Absolute Error -- the average of the absolute value of the differences between observed and computed hydrographs.

Average Percent Absolute Error--the average of absolute value of percent difference between observed and computed hydrograph ordinates.

APPLICATION OF THE CALIBRATION CAPABILITY

The following steps represent the strategy recommended to determine the model parameters for HEC-1 in this study:

1. For each storm, the base flow and recession parameters (STRTQ, QRCSN, and RTIOR) are determined graphically. These parameters can not be estimated automatically and must be entered as input. These parameters are placed on the BF input card for each storm.

- 2. For each storm and each gage, the optimal estimates of all unknown unit hydrograph and loss rate parameters (TC, R, STRTL, and CNSTL) are determined using the automatic calibration feature.
- 3. A value of the constant loss rate, CNSTL, is estimated. A regional value of CNSTL based on all storms at all the gages is selected.
- 4. With CNSTL fixed at the selected value, the automatic calibration feature is again employed for all the storms at all the gages. A regional relationship for the initial loss, STRTL, is selected. (See Figure 5.1)
- 5. With both loss rate parameters fixed the automatic calibration of the model is repeated.
- 6. Values of TC + R are selected for each basin. An average value of R/(TC + R) is selected for the region.
- 7. After all parameters are selected, the values are verified by simulating the response of the gaged basins to other storms not used in the calibration process.

REGRESSION ANALYSIS APPROACH TO PARAMETER ESTIMATION
(HEC-1, 1982)

Multiple linear regression techniques were used to correlate the unit hydrograph parameters with several

basin characteristics. Nonlinear relationships were investigated by transforming values logarithmically. Multiple regression analysis requires a computer program. Herein, the computer package MINITAB was used in the analysis.

The parameters TC and R were considered dependent variables and the basin characteristics were the independent variables. TC, R, (TC + R), and R/(TC + R) were analyzed, first, by considering several basin characteristics and, later, by reducing the number of independent variables until the "best" relationships (equations) were found.

The statistics describing the "goodness-of-fit" of the regression equation to the data are used in evaluating the analysis. These statistics are the coefficients of determination (both the adjusted and unadjusted), the partial determination coefficient, and the standard error of estimate. Their definitions are given below:

- 1. The adjusted and unadjusted multiple-determination coefficients (R<sup>2</sup>) are a measure of the percent of variance in the dependent variable explained by the independent variable. The magnitude of these coefficients varies from 0 to 1. The closer to the value of 1, the greater the reliability is of the estimate.
- 2. The partial-determination coefficient  $(r^2)$  is a measure of the importance of an independent variable by determining the reduction in variance in the

dependent variable when the variable is included with the other independent variables.

3. The standard error of estimate ( $S_e$ ) is the standard deviation of the difference between the observed dependent values and the values computed from the regression equation in the units of the dependent variable; therefore, it must be compared with the mean and the standard deviation of that variable.

For a more detailed description of multiple-regression analysis the reader is referred to a statistics book, such as Benjamin (1970, p.419). The general rule is to use the values of  $\mathbb{R}^2$ ,  $\mathbb{R}^2$  adjusted, and  $\mathbb{S}_e$  computed for each regression equation as a guide and select the equation with the fewest independent variables and the best values of  $\mathbb{R}^2$  and  $\mathbb{S}_e$ .

# CHAPTER FOUR DATA

## PURPOSE AND SCOPE

The purpose of this chapter is to present the data used in this study. The scope of this chapter is to display the basic data obtained from various sources after some refinement into a usable form for this analysis.

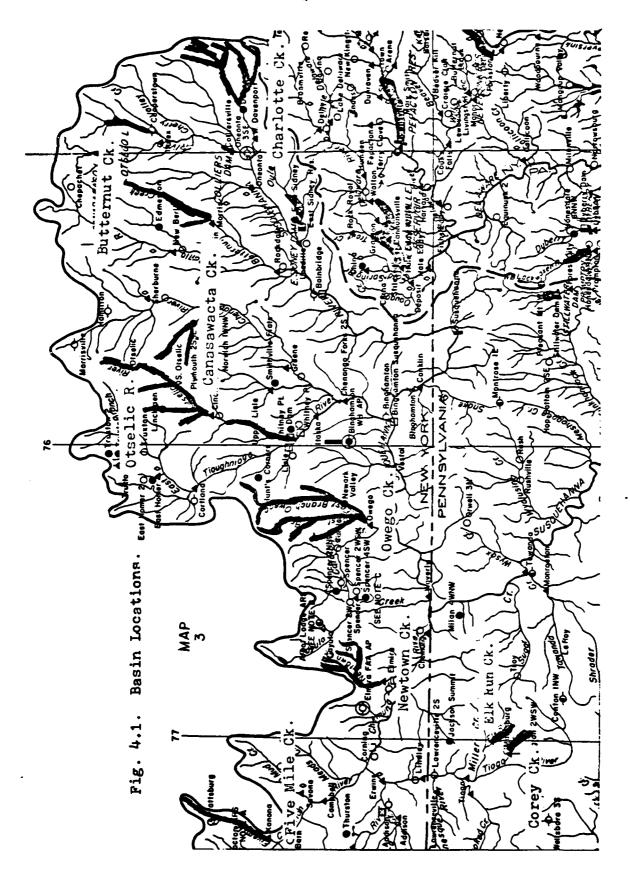
## BASIN CHARACTERISTICS

Nine watersheds in the Susquehanna and Chemung River basins were chosen for the present investigation. The watersheds were unregulated and possessed a U.S. Geological Survey (USGS) stream gage. The locations of these gages were taken as the outlets from the watersheds. The nine basins are henceforth referred to by the main channel watercourse. Their names are: Butternut Creek, Canasawacta Creek, Charlotte Creek, Corey Creek, Elk Run Creek, Five Mile Creek, Newtown Creek, Otselic River, and Owego Creek. All basins are in New York state except Corey Creek and Elk Run Creek, which are in Pennsylvania. Figure 4.1 depicts the basins used and their location with respect to one another.

The basin characteristics were measured from USGS

Topographic Quadrangle maps ( 1:2400 ). Table 4.1 contains

a tabular synopsis of the basin characteristics chosen



	ଧ	ΩI	<u>S</u> 2	SWI		Ы	2	SF1	SF2	ALT
Butternut Creek	59.7	34.12	34.25	80.26	1.740		20.83 8.56 7.27 .419	7.27	.419	780
Canasawacta Creek	57.9	46.55	50.80	49.89	1.620 12.42 8.48 2.66	12.42	8.48	2.66	.691	830
Charlotte Creek	167	34.60	36.83	34.60 36.83 67.58 1.633 24.77 15.15 3.67	1.633	24.77	15.15	3.67	.589	1110
Corey Creek	12.2	135.48	164.12	12.2 135.48 164.12 106.39 1.422 5.38 3.03 2.37	1.422	5.38	3.03	2.37	.733	1003
Elk Run Creek	10.2	105.23	115.52	10.2 105.23 115.52 123.89 1.188 6.74	1.188	46.9	4.43 4.45	4.45	.535	850
Five Mile Creek	8.99	29.75	38.10	66.8 29.75 38.10 67.06 1.790 19.51 10.42 5.70	1.790	19.51	10.42	5.70	.473	830
Newtown Creek	77.5	45.15	46.94	77.5 45.15 49.34 70.75 1.494 17.92 10.68 4.14	1.494	17.92	10.68	4.14	.554	865
Otselic River	147	23.72	25.19	25.19 51.74 1.473 29.51 14.24 5.92	1.473	29.51	14.24	5.95	494.	800
Owego Creek	185	29.22		32.34 54.38 1.809 30.57 15.95 5.05	1.809	30.57	15.95	5.05	.502 1105	1105

Table 4.1 Basin Characteristics.

for the present analysis on the basis of the review in Chapter 2. A definition of each term follows: (after Linsley (1955, p. 313))

DRAINAGE AREA (DA)--The area in square miles of the water-shed above the stream gage. Any excess rainfall falling in the watershed would at some time pass through the gaging station. Rainfall falling outside of this area would drain into another watershed. These data were obtained from USGS Water Supply Papers.

SLOPE (S)--This is the average main channel slope in vertical feet per horizontal channel mile. Slopes were obtained by plotting the elevation of the map contours (in feet) versus the distance of the map contours from the gaging stations (in miles) and taking the straight line slope fitted by the least squares method. The steeper the slope the faster the basin will drain.

SLOPE 2(S2)--This is similar to slope (S) except the main channel is divided into two segments. Two straight lines are drawn representing a closer fit to the actual main channel profile. Slope 2 is a weighted average of these two segments.

S2 =((slope of segment 1)(horizontal length of
 segment 1) + (slope of segment 2)(horizontal
 length of segment 2))/(total main channel length)

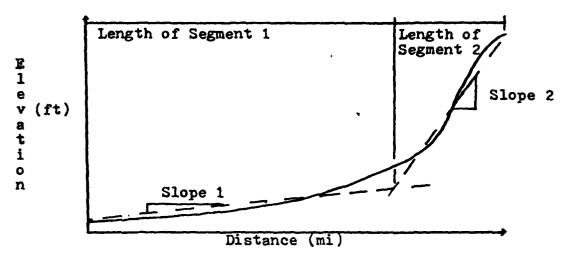


Figure 4.2. Slope2 Graphical Representation.

SLOPE WT (SWT)--This slope is a weighted slope intended to represent a better picture of the basin's slope characteristics by including the slope contribution of some of the tributary streams. Its feet is also listed in feet per mile.

DRAINAGE DENSITY (DD)--This is a measure of the amount of stream channel miles per basin square mile. It is obtained by measuring the cumulative length of every stream channel shown on the USGS topographic map. This cumulative length is divided by the drainage area. The higher the drainage density, the smaller distance any particle of excess rainfall must travel before entering a stream channel. This implies that the larger the DD value, the faster the basin will drain.

LENGTH OF THE MAIN CHANNEL (L) -- This is a measure of the longest watercourse measured from the gaging station to the furtherest point of stream channel origination upstream.

L is measured in miles.

LENGTH TO THE BASIN CENTER OF MASS (LCM)--The center of mass is found by hanging a paper cut-out diagram of the drainage basin and marking off the vertical lines obtained from three different suspension points. The value for LCM is obtained by measuring the distance from the gaging station to the point on the main channel opposite the center of mass. LCM is measured in miles.

SHAPE FACTOR ONE (SF1)--This is a measure of the basin's shape. It is equal to the length of the main channel squared divided by the drainage area, i.e.  $L^2/DA$ .

SHAPE FACTOR TWO (SF2)--It is similar to SF1. It is the ratio between the diameter of a circle, the size of the drainage area (DA), and the length of the main channel (L), i.e. SF2 = DIA/L where DIA =  $(4*DA/\pi)$ . This term was proposed by Cruff and Rantz (1965).

<u>ALTITUDE</u> (ALT)--This is a measure of the vertical drop between the elevation at the main channel headwaters and the gaging station in feet.

## STORM/FLOOD DATA SELECTION

An effort was made to obtain 5-10 of the major flood events occurring in each basin since the record flooding which was produced by Hurricane Agnes (June 1972). There were several seasons for selecting events since the summer of 1972. First, it is desired to minimize the influences

of basin changes which either naturally occur or are caused by man over time. It was felt that a time period of ten to eleven years would be a reasonable length of time to minimize basin change effects. Second, results of studies prior to, or as a result of, Hurricane Agnes may be compared with the results of this work. The storm events for each basin are presented in Table 4.2. It should be noted that not all storm dates were taken after June 1972, because not enough major storms had occurred to provide a sufficiently wide data base to conduct the analysis. In addition, only events were selected which were free from any indication of snowmelt, ice jam, or difficulties involving the stream gaging stations. These problems are beyond the scope of this work.

## PRECIPITATION DATA

Precipitation data for the storm events in each basin were obtained from documents published by the National Weather Service. Basins were traced on paper from small scale maps. Locations of the rain gages in the vicinity of each basin were plotted on the tracing paper. The Thiessen Polygon method was used to divide up the basin. Planimeter measurements were made for all recording and nonrecording gages influencing the basin. The fraction of the basin contained in each polygon was recorded along with the rainfall amounts obtained from the Climatological Data

Butternut Creek (NY) 10-18-75 * 10-09-76 10-21-76 9-26-77 10-17-77	Five Mile Creek 5-07-75 9-26-75 9-14-77 9-25-77 * Newtown Creek (NY)
Canasawacta Creek (NY) 6-26-68 11-18-68 6-24-69 7-31-71 * 6-22/23-72 7-03-74	9-26-75 10-09-76 * 10-21-76 9-25-77 10-17-77 11-04-77 11-08-77
Charlotte Creek (NY) 6-29-73 12-21-73	10-06-79 11-26-79 10-28-81 *
Corey Creek (PA)  11-14-72  9-26-75  10-09-76 *  5-14-78  10-05-79  11-26-79  6-05-82	Otselic River (NY) 7-03-74 9-26-75 10-09-76 * 10-17-77 4-05-78  Owego Creek (NY) 11-08-72
Elk Run Creek (NY)  10-22-70  11-14-72  9-26-75  10-09-76  10-20-76  5-14-78  8-06-78 *	9-26-75 10-09-76 * 9-25-77 10-17-77

# Total Storms

52

# Table 4.2

Storm Data (\*)--Verification storms not used in model calibration.

and Hourly Precipitation Data (recording stations), documents of the National Weather Service. Appendix A contains the watershed rain gage locations and the Thiessen polygons.

## STREAMFLOW DATA

The streamflow data for each flood event at a particular gage were obtained through the Regional Office of the U.S.G.S. located in Ithaca, New York and the State office of the U.S.G.S. located in Harrisburg, Pennsylvania for the Pennsylvania stations. Stage data were obtained as height of the stream at particular times. For this study therefore, the stage data (FT) had to be converted to flow rates (FT<sup>3</sup>/sec) by the use of the stream's rating curve. Flow rates were tabulated and retained for use as input for the HEC-1 calibration sequence.

The recession parameters, RTIOR and QRCSN were calculated as described in Chapter Three and retained as input for the computer model.

## INPUT FORMATTING AND DATA STRUCTURE FOR HEC-1

Input for the HEC-1 model is accomplished by creating a card deck organized in a certain structure and sequence for easy computer reading. Each card of eighty columns is set up in the following manner; columns 1 and 2 are used to identify the card type and purpose. Following columns 1 and 2 are ten data fields of eight columns each,

except the first field which is composed of only six columns. More detailed discussion of data organization may be found in HEC-1 manuals. A brief line-by-line explanation of a typical card deck as used in this study may be found in Appendix B. as well as the input card data information for all the storms used in this study.

## REGRESSION DATA FOR THE INITIAL LOSS PARAMETER (STRTL)

Because STRTL is storm dependent, a method of relating the STRTL values to the total amount of rain occurring in the 2, 3, 4, 5, 7, 8, and 10-day period prior to the storm was attempted. The statistical package MINITAB was used to perform regression analysis on the data. The data are displayed in Table 4.3. The reader may obtain further information on MINITAB by reading the MINITAB Reference Manual or by contacting MINITAB PROJECT, Statistics Department, 215 Pond Laboratory, Pennsylvania State University, University Park, PA 16802.

The values in columns two through eight were obtained by summing the products of the total rainfall recorded for each rain gage covering a portion of the basin and the fraction of the drainage area covered by that gage. The analysis of this data will be discussed in Chapter Five.

Day Period

Basin/Storm	STRTL	Sa	<b>D3</b>	<b>D</b> 4	<b>D</b> 5	D?	D8	D10
Butternut * 10-18-75 10-09-76 10-21-76 9-26-77 10-17-77	53 0 0 0	.31 0 0 .73 .80	0 .01 .73	.87 0 .01 .75 1.01	0. .40 1.14	0 1.29	0 1.29 3.19	1.25 .20 1.29 5.40 2.02
Canawacta 6-26-68 11-18-68 6-24-69 * 7-31-71 6-22/23-72 7-03-74	1.09 .26 2.26 .08 .98	.09 .36 .44 .03 0	.09 .36 .51 .88 0	.88	.78	.38 1.47 .81 1.6 1.67	1.67	.40 1.96 .85 2.67 1.67
Charlotte 6-29-73 12-21-73		.01 .05	.01	.14 1.54	.32 1.54	.41 1.71	.41 1.71	.56 1.8
Corey 11-14-72 9-26-75 10-09-76 5-11-78 * 10-05-79 11-26-79 6-05-82	.33 .32 1.33 .56 .14	.46 .37 0 0 .69 .76	1.14	.21	.33 1.18 .77	.33 2.02 .77	.42 2.02 .77	.99 .77 .06 .87 2.02 .77
Elk Run 10-22-70 11-14-72 9-26-75 12-09-76 10-20-76 5-14-78 * 8-06-78	1.01 .39 .01 1.45 .69	0 0 .45 0 0 0	0 0 .7 0 0 0	0 0 .7 0 0 0	1.18 .93 0 0	2.28 1.05 0 .29	.55 2.28 1.05 0 .29 .78 2.75	2.28 1.05

Table 4.3 Previous Rainfall Amounts

# Day Period

Basin/Storm	STRTL	D2	DЭ	D4	<b>D</b> 5	D7	<b>D</b> 8	D10
Five Mile 5-07-75 9-26-75 9-14-77 * 9-25-77	0 .92 1.34	.1 .6 .01	.9 .01	·9 ·01	ed * .9 .01 2.19	1.1	1.1	* 1.1 .27 5.34
Newtown 9-26-75 *10-09-76 10-21-76 9-25-77 11-04-77 11-11-77 10-17-77 10-06-79 11-26-79 *10-28-81	.01 .36 .66 .65 .13 .19 .01 .96	0 0 .04 0 .12 .54	.81 0 .09 0 1.64 1.09 .25	.84 0 .46 0 2.05 1.09 .26 .52	.94 .05 .80 0 2.05 1.85 .52	1.43 0.26 1.69 0 2.05 3.02 .76 1.44	1.43 0.27 2.48 0 2.05 3.02 .76	1.43 .1 .27 3.9 0 2.05 3.02 .93 1.44
Ostelic 7-03-74 9-26-75 *10-09-76 10-17-77 4-05-78 Owego 11-08-72 9-26-75 *10-09-76 9-25-77 10-17-77	2.33 1.3 .04 0	.11 .24 0 .63 .01 .46 .03 .03	.16 .24 0 .63 .1 .07 .46 .03	.49 0 .63 .26 .08 .63	.49 0 .85 .26 .93 .78 .03	.91 0 1.34 .28	.99 1.2 .03 4.0	1.56 .98 .13 2.02 1.26 1.6 1.2 .13 5.25 1.37

Table 4.3. (Continued)
Previous Rainfall Amounts

## CHAPTER FIVE

## RESULTS AND DISCUSSION

## PURPOSE AND SCOPE

The purpose of this chapter is to present the results of the calibration effort for the basins and storms selected as the control group; these comprise selection of the unit hydrograph and regional loss rate parameters, determination of regression relationships, and verification of the results using uncontaminated flood events. The scope of this chapter is to display the results and to discuss the findings in relation to an earlier study.

## MODEL CALIBRATION FOR EACH BASIN

The 'best' fit values of the four model parameters for each event are presented in Table 5.1. Each event was initially optimized using the model's default values. From the initial results the starting values of the parameters were manipulated until the 'best' fit choices produced the 'best' statistical summary results. These summaries as well as other 'best' fit results may be found in Appendix C.

Once the optimized parameters were determined for the control storms, the parameter for the constant loss rate, CNSTL, was averaged for each basin. For certain

Table 5.1.

\*Best-Fit\* Values Optimization Results.

Basin/Storm	<u>TC</u>	<u>R</u>	STRTL	CNSTL
Butternut Creek 10-10-76 10-21-76 9-26-77 10-17-77	13.83 10.98 9.93 14.77	6.24 14.06 10.52 4.53	.53 0 0	.05 0 0 0
Canasawacta Creek 6-26-68 11-18-68 6-24-69 6-22/23-72 7-03-74	4.25 2.25 1.83 3.33 6.13	9.69 10.61 12.37 7.94 4.10	1.23 0 2.0 .99	.06 0 .20 01 .07
Charlotte Creek 6-29-73 12-21-73	14.61 4.21	20.90 15.21	. 92 . 50	.09 .04
Corey Creek 11-14-72 9-26-75 10-09-76 5=14-78 11-26-79 6-05-82	1.62 2.27 3.97 1.03 2.60	10.39 2.71 4.73 6.94 3.73 11.96	.30 .39 1.62 .28 .22	.12 .08 .15 .15 .09
Elk Run Creek 10-22-70 11-14-72 9-26-75 10-04-76 10-20-76 5-14-78	1.07 1.50 1.03 4.83 1.92 1.06	5.95 8.54 4.12 1.69 3.69 2.36		.16 .14 .10 .07 .07

Table 5.1 Continued 
"Best-Fit" Values Optimization Results.

Basin/Storm	TC	R	STRTL	CNSTL
Five Mile Creek 5-07-75 9-26-75 9-14-77	14.73 4.27 9.65		.2 .89 1.09	.02 .02 .11
Newtown Creek 9-26-75 10-21-76 9-25-77 10-17-77 11-04-77 11-08-77 11-11-77 10-06-79 11-26-79	12.48 11.01 10.42 8.31 9.18 8.41 7.86 2.64 5.79	11.90 14.53 13.48	.03 .36 .66 .01 .65 .13 .19 .72	.05 .07 .07 .06 .07 .02 .02 .14
Ostelic River 7-03-74 9-26-75 10-17-77 4-05-78  Owego Creek 11-08-72 9-26-75 9-25-77 10-17-77	10.37 6.14 15.46 6.32 8.35 7.09 6.15 10.12	29.53 26.52 18.94 15.23 17.78	2.0 1.27 .01 .05	.15 .06 .05 .01 .06 .03 .03

values within the observed range of this parameter (around its average) and depending on its performance the basin value of CNSTL selected for each basin is displayed in Table 5.2.

Because of the unique structures of the Butternut and Five Mile Creeks, their values do not appear to fit into the general pattern of increasing CNSTL values as one proceeds southwestwardly from the Upper Susquehanna region. Butternut Creek has a rather large value for SWT (80.26 ft/mi) and the largest value for SF1 (7.27) implying a long,narrow, steep-sided basin. Such a basin would not allow as much water to be intercepted/infiltrated as the other basins. On the other hand, a casual observation of the Five Mile Creek basin on a topographic map shows a great deal of swamp area. This condition would retard the flood wave; however, less rainfall would be lost due to the already wet soil condition.

Holding the parameter CNSTL constant for each basin, the other three parameters were optimized. The values of the initial loss parameter, STRTL, were regressed against the amount of total precipitation falling at different time periods preceeding the storm. The best derived regression equation was STRTL = 0.861 - 0.504 \* D5, where D5 is the total amount of rainfall occurring five days prior to the beginning of the storm. See Figure 5.1 for a plot of STRTL vs D5 for all of the control storms. This equation has an unadjusted multiple-determination coefficient,

Table 5.2.

Basin Averaged Constant Loss Rates (CNSTL).

	CNSTL (inches/hour)
Butternut Creek	0.00
Canasawacta Creek	0.05
Charlotte Creek	0.05
Corey Creek	0.10
Elk Run Creek	0.10
Pive Mile Creek	0.03
Newtown Creek	0.07
Ostelic River	0.05
Owego Creek	0.05

 $R^2$ , equal to .192.

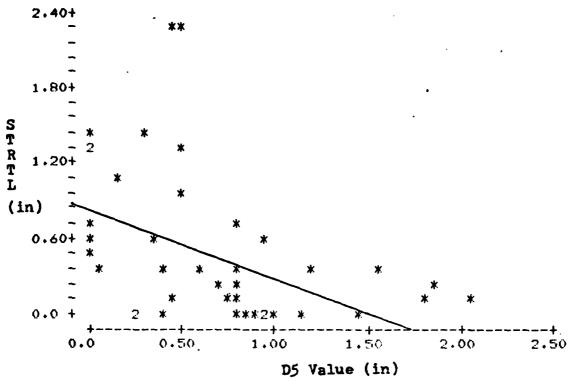


Figure 5.1. Regression Equation for STRTL Plot.

Although the coefficient is not as high as desired, the D5 term gave the best correlation. The control storms plotted below the line in Figure 5.1 will give conservative forecasts. This means higher computed peaks than the observed peaks. The STRTL term is rather sensitive (see Sensitivity Analysis section, later), however, in the absence of any other method, this approach was used.

The values for STRTL were predicted by means of this regression equation. Each storm was then optimized for the TC and R parameters. The final parameter values for

the control storms are listed in Table 5.3.

These values were used in computing and comparing the verification storm group hydrographs with the observed flood hydrographs. The statistical results are presented in Appendix D for the verification storms.

# COMPARISON WITH PARAMETERS DETERMINED PRIOR TO HURRICANE AGNES (1972)

In June 1970 the Susquehanna River Basin Study Coordinating Committee published an extensive work on the Susquehanna basin. Appendix D (Hydrology) of that study contained the Clark unit hydrograph parameters for several gaged basins, obtained on the basis of a similar computer program developed by the Hydrologic Engineering Center, U.S. Army Engineer District, Sacramento, CA, (now HEC is located in Davis, CA).

The data base used in the study consisted of four major flood events. Basin average rainfall was determined by Thiessen polygons. Two or four-hour unit hydrographs were developed depending on the drainage area of the basin. The TC and R parameter values are presented in Table 5.4. Not all of the basins could be compared with the current results because some stream gages have been relocated since the publication of the study. Table 5.4 shows a very good comparison between most of the terms. The values of R for Charlotte and Five Mile Creeks are not in good

Table 5.3.

Derived Parameters.

Basin	TC	R (hours)
Butternut Creek	11.66	6.76
Canasawacta Creek	8.65	8.59
Charlotte Creek	11.68	27.70
Corey Creek	1.90	8.27
Elk Run Creek	2.02	8.13
Five Mile Creek	7.47	46.12
Newtown Creek	9.00	14.46
Ostelic River	9.80	36.13
Owego Creek	9.35	13.83

agreement. A possible explanation for the Charlotte Creek discrepancy is the fact that only two storms meeting the established criteria in the past fifteen years could be found for this basin. The expected value variance of any derived parameter based on only two data points is very great. Therefore, one would not expect the values of the two works to be very close.

The discrepancy of the R values for Five Mile Creek could be related to the goodness-of-fit of the storms selected for this study. Looking at the plots of the 'best' fit computed flow with the observed flow, the reader will notice that the computed hydrograph had a rather difficult time matching the observed hydrograph. Whether the rain gage in the middle of the basin failed to provide accurate data or the U.S.G.S. gaging station measurements were fouled by mechanical defects is not clear. The author is inclined to doubt the observed hydrograph measurement because of the peculiar combination of spike and round portions.

	Study(Jı	ın 70)	Pres	ent Results
	TC	R	TC	R
Canasawacta Ck. Charlotte Ck. Five Mile Creek Newtown Creek Owego Creek	2.03 11.83 5.50 9.71 10.10	11.04 14.79 19.00 18.39 11.82	8.65 11.68 7.47 9.00 9.35	8.59 27.70 46.12 14.46 13.83

Table 5.4.
Parameter Comparison with Previous Study

#### SENSITIVITY ANALYSIS OF PARAMETERS

A sensitivity analysis was performed on a selected storm whose 'best' fit results initially yielded a very good statistical summary. The storm selected from the control group was the Newtown Creek 11-04-77 storm. Holding the other three parameters constant, the parameter TC was decreased and increased 5 and 10%. The same was done for the parameter R. The loss rate parameters are not normally known to such a high degree of accuracy. Therefore, the loss rate parameters were not varied by 5 and 10% because the results were not significantly different. For the initial loss rate parameter, STRTL, the value was decreased and increased .1 and .2 inches to show a better variability. The parameter for constant loss rate, CNSTL, was varied by .02 and .04 inches/hour each way. These variations are reasonable for any basin under investigation. The results are displayed in Table 5.5.

The reader is directed to the Peak Flow column. The Peak Flow percent was not very sensitive to variations in the TC and R terms. The most change was about 6% between the computed and observed peaks. Variations in the normal range of expected values of the loss parameters are a different story. A decrease of STRTL of .2 inch caused an overestimation of the peak by 14%. While a .04 inch/hour decrease or increase in the CNSTL produced approximately a 40% change in both directions.

Adjustments	Sum of Flows Diff	Equiv Depth Diff.	Mean Flow Diff	Time To C.M. Diff	C.M. to	Peak Flow	Time of Peak	AVG A ABS Error	el el
Control TC -10% -5% +10%	5.54 5.45 1.86 1.86 1.86 1.86 1.86 1.86 1.86 1.86	031 .018 .004 .004	1,884		-5.37 -1.84 -39 -39 -39 -39	2.87 1.33 -1.8 -1.8	01400	12.61 10.77 13.01 9.36	63 88 80 80 80
H 1-1+1+ 1001-1+1+ 1001-1+1+	6.81 4.21 -4.64 -2.95	.035 039 015	67 42 -28 -29		- 4.33 -4.02 2.22	6.78 3.59 -2.29 -4.98	••••	11.74 10.46 11.80 9.62	104 74 60 77
STRFL2 1 +.1	27.45 14.60 -11.10 -17.61	.141 .075 057 09	271 144 -110 -174	75 37 1.02	-7.09 -3.72 8.12 12.52	14.23 7.34 -6.18 -9.93	4000	46.17 24.84 22.21 31.56	299 163 139 208
CNSTL04 02 +.02 +.04	42.20 21.97 -18.48 -38.70	.217 .113 095	417 217 -182 -387	.06 .07 .13	.62 .78 1.37 1.97	38.5 19.55 -18.35	0000	42.77 24.4 19.05 38.10	536 282 243 496
Initial Par	Parameter '	Values	were	R = 11	.18	STRTL = CNST =	65		
(-) 0BS >cc (+) 0BS <cc< td=""><td>OBS ≯computed OBS ≮computed</td><td></td><td></td><td>Z</td><td>Newtown Greek 11-04-77</td><td>Greek -77</td><td></td><td></td><td></td></cc<>	OBS ≯computed OBS ≮computed			Z	Newtown Greek 11-04-77	Greek -77			

Table 5.5.

Sensitivity Results

The results of the analysis show that the loss parameters are the most sensitive and must be estimated accurately to give reasonable results.

The reader may recall the discussion in Chapter Three concerning the optimized solutions of HEC-1 not being the global solution. This point is reflected in this analysis where a 5% increase in TC produced a better fit than the original controlled parameter storm.

# REGRESSION ANALYSIS OF THE PARAMETERS TO AND R

Various combinations of the basin characteristics
listed in Table 4. were tried. The only restriction placed
on the combinations was that only one slope value and
only one shape factor were permitted in any one combination. Combinations of the characteristics as well as
their log values were evaluated. Table 5.6 lists the
top combinations of each type along with the statistical
coefficients. Table 5.7 lists some of the characteristic
combinations used.

	R <sup>2</sup>	R <sup>2</sup> -ADJ	Standard <u>Deviation</u>
$TC = 7.942.63(DA)^{.66}/(ALT)^{1.44}$	.910	.879	0.2474
$TC = 450.34/(S2)^{1.08}$	.878	.861	0.2659
$TC = 5.93 (DA/S)^{.412}$	.871	.853	0.2736
$TC = 1.73(L\sqrt{DA/S})$	.861	.842	0.2836
$TC = 3.60 (L/VS)^{.716}$	.847	.825	0.2980

Table 56. Highest Correlated Regression Equations.

	R <sup>2</sup>	R <sup>2</sup> -ADJ	Standard Deviation
TC+R = 626.41/(S).87	.731	.693	0.3365
TC+R=14.15(L/VS)·549	.686	.641	0.3638
TC+R=20.91(DA/S).312	.686	.641	0.3639
TC+R=8.25(L \(\overline{DA/S}\)\.35	.677	.630	0.3693
TC+R=15.18(L/\swT).653	.667	.619	0.3748
TC+R=507.76/(S2)·794	.658	.610	0.3795

Table 5.6.

Highest Correlated Regression Equations (Continued)

	<u>DA/S</u>	DA/SWT	L LCM/VS	LVDA/S	<u>l/V</u> s	L/VSWT
Butternut Ck.	1.75	.74	30.53	27.55	3.57	2.33
Canasawacta Ck.	1.24	.84	15.44	13.85	1.82	1.50
Charlotte Ck.	4.83	2.47	63.80	54.42	4.21	3.01
Corey Ck.	.09	.11	1.4	1.61	.46	.52
Elk Run Ck.	.10	.08	2.91	2.10	.66	•59
Pive Mile Ck.	2.25	1.00	37.27	29.23	3.58	2.38
Newtown Ck.	1.72	1.10	28.48	23.48	2.67	2.13
Ostelic Ck.	6.2	2.84	17.72	73.46	6.06	4.10
Owego Ck.	6.33	3.40	90.20	76.92	5.66	4.15

Table 5.7 .

Basin Parameter Combinations

The first equation in each group in the table has the highest correlation coefficients, lowest variance, and least number of parameters. The parameters relating best to the Clark unit hydrograph parameters TC and R were found to be the drainage area (DA), altitude difference (ALT), and the main channel slope (S). The shape factors were found to be highly related to the other terms. They were eliminated without exception in every stepwise regression operation. The term S2 faired better in the TC group than the S term. The opposite is true in the TC + R group. However, the difference between them is not great and does not warrant the additional effort of deriving S2.

#### AUTOMATIC CALIBRATION RESULTS

In Chapter Three under the section "Calibration of Model
Parameters", it was mentioned that the optimization technique
employed by HEC-1 does not guarantee that the parameters derived
are global solutions. An example of this is displayed in Figures
5.2 and 5.3. Figure 5.2 is the graphical results obtained using
the HEC-1 initial parameters as starting points. Figures 5.3
is the final results used in this analysis for the Butternut
Creek (9-26-77) storm.

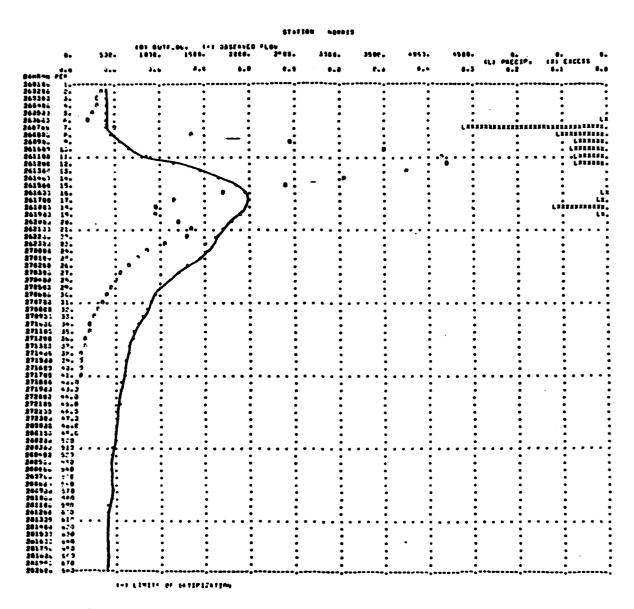


Figure 5.2. Initial Optimization Results.

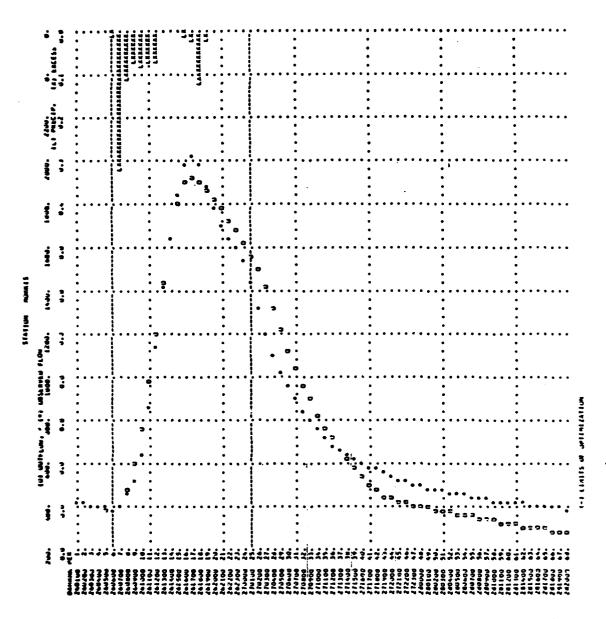


Figure 5.3. Finial Optimization Results.

#### CHAPTER SIX

# SUMMARY AND CONCLUSIONS

McSparran (1968, p.937) found in his study of 32 watersheds varying in drainage area from 2 to 210 square miles in Pennsylvania that relating hydrograph parameters to watershed characteristics is not an easy task. He found nonlinearity factors present in every storm, adjusting and altering the parameters. As a result he felt that the unit hydrograph method was not as accurate as it should be. He stated that major floods would be underpredicted and on the average the observed peaks would be underpredicted by 22% using the Snyder unit hydrograph method. Yet by using the Soil Conservation Service (SCS) method the average peaks were predicted 51% too high.

On the other hand, Heerdegen (1974, p.1143) concluded with a resounding "affirmative" to his article's question-title, "Unit Hydrograph: A Satisfactory Model of Watershed Response?" Sixty measured parameters for each of the sixteen watersheds in Pennsylvania were used in analyzing ninety-six flood events. Five parameters were eventually selected as the most influential for his area. However, they were not constant from area to area. In his study he concluded that the storm variable of precipitation duration and volume of precipitation excess are major factors in flood hydrograph responses.

# SOURCES OF ERROR

Whenever one attempts to measure a physical property in nature, there is some error. Rain gages and streamflow gages leave a lot to be desired when it comes to accurate measurements. Rain gages are point measuring instruments and they hardly ever represent the amount of rainfall deposited on a basin accurately. In some cases rain gages at a great distance from the basins had to be used, because the local data were fouled or not reported.

model makes some simplifying assumptions for ease of computation, cost, or speed. HEC-1 is a lumped or bulk parameter model. It groups all of the irregularities and unique features of a basin into one big box. Richards (1955, p.110) presented nine factors affecting runoff.

Some of these factors, such as moving storms and intensity variations, are ignored by HEC-1.

In the absence of specific information, a synthetic time area curve (supplied by HEC-1) was made to create a translation hydrograph for each basin. The basins varied in many sizes and shapes and forcing the time-area curve to fit all the basins seems unrealistic.

The type of optimization routine, i.e. univariant search technique, does not guarantee the results to be a global solution and indeed the user of the model must

be aware of its hazards. Picking new starting points and adjusting the parameters to get a better fit is sometimes more luck than skill.

The loss rate selection is a major source of error. The inability to obtain a consistent and reliable STRTL predicting device greatly influenced the final results and future applications of this study. Bulk loss rates are perhaps better suited for large areas.

Selection of the Clark method to derive flood hydrographs affected the results. The application of a linear numerical model, although frequently used in government and private studies, to a nonlinear problem induces error.

Are the areas of interest, i.e. the Upper Susquehanna and Chemung River Basins, compatable? The nine stream basins are supposed to be hydrologically similar. Are they? What does hydrologically similar mean? This author concluded that the basins are fairly hydrologically similar because the slopes and valley structures appear to be consistent among the nine basins and most major storms track across both river basins. What basin or climatic factors were not included which would have improved the correlation between the model parameters and the basin characteristics? Chow (1964, p.27-104) states that some areas are just not adequate for unit hydrographs.

The unit hydrograph assumptions of time invariance and linear runoff responses are not strictly valid in

real life. Basin characteristics do change with time, e.g. land cover and seasonal changes influence loss rates.

Although the storms are the largest occurring in the last ten to fifteen years (unaffected by snow), most are not to be considered major storms. Sokolov (1976, p.199) quoted U.S. Army Corps of Engineers' studies when they said that in the majority of the basins studied the peak ordinates of the unit hydrographs derived from a major flood are generally significantly greater than those derived from data from a minor flood. Those studies stated that unit hydrograph peak ordinates from major floods were twenty-five to fifty percent greater than those derived from minor floods. Therefore, one would expect the results obtained in the application of the regression relationships to be somewhat underpredictive, i.e. on the conservative side.

The storms selected cover several seasons from early April until December. Although Heerdegen (1968, p.1143) found no significant seasonal differentiation in the unit hydrograph at the five percent level, Sherman (1932, p.504) felt that unit hydrographs should be seasonally segregated.

# LIMITS OF APPLICATION

The data were obtained from streams and rain gages in south central New York and north central Pennsylvania and they should be applied in that area. With the exception

of Five Mile Creek, the streams appear to be free from a great deal of valley storage and swampy areas. The results were obtained without snowmelt or ice jams affecting the outlet flow rate. Therefore, the results should apply primarily to summer and fall seasons.

### AREAS IN NEED OF FUTURE INVESTIGATION

The lack of adequate soil moisture data hindered the proper calibration of the parameter STRTL. Without adequate or reasonable loss rate parameter values, the model will not serve as a reliable predictive tool.

#### SUMMARY

By using the Flood Hydrograph, HEC-1, computer package, the Clark unit hydrograph, and the initial/uniform loss rate, parameters were determined for nine watersheds in the Chemung and Upper Susquehanna River basins. Fifty-two storms were used in the analysis on the watersheds which range in drainage areas from 10.2 to 185 square miles. Regression equations predicting the time of concentration, TC, and the term TC + R were derived.

# CONCLUSIONS

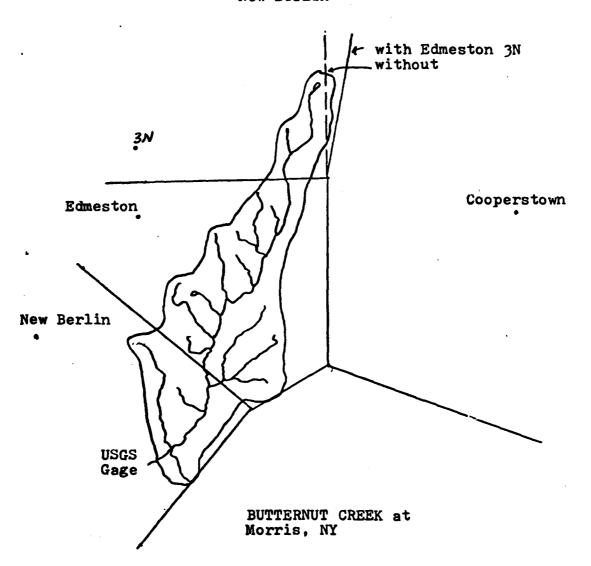
Specifically, the conclusions of this thesis are:

- 1. HEC-1 may be applied successfully in the study area.
- 2. Sensitivity of the parameters TC, R, STRTL, and CNSTL in HEC-1 are established.
- 3. The HEC-1 optimization search technique should be improved to give a global solution.
- 4. The relationship to determine the term STRTL needs to be improved.
- 5. Regional values of the term CNSTL are now established.
- 6. The initial and constant loss rate method is not the best method but it works fairly well.
- 7. The results should be applied seasonally.
- 8. Good correlation between the parameters TC and R with basin characteristics were obtained. The recommended equations are the first in each group in Figure 5.
- 9. A very good comparison with a 1970 study was obtained for the terms TC and R.

Appendix A Watershed Rain Gage Locations and Thiessen Polygons.

Recording Stations: Edmeston Edmeston 3N

Non-recording Stations: Cooperstown New Berlin



Recording Stations: Plymouth Plymouth 2SSE

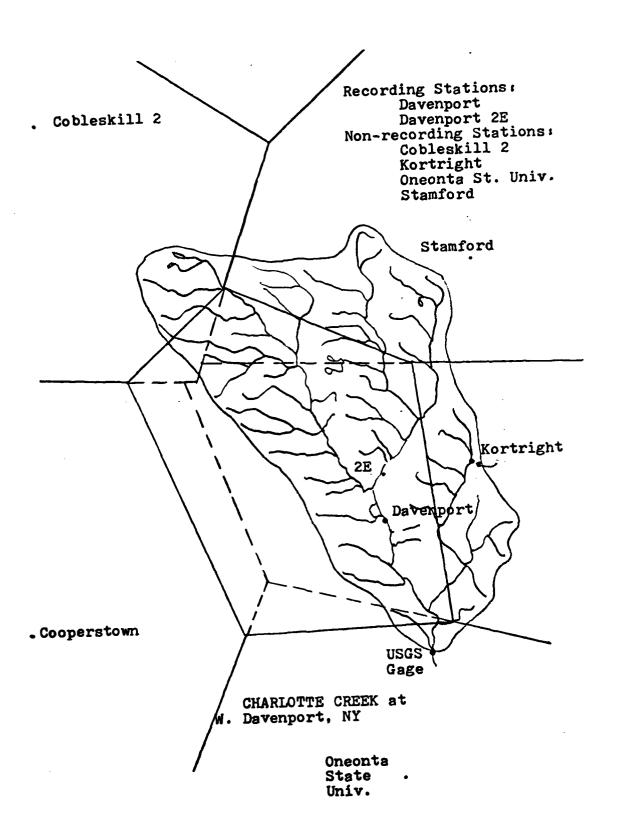
Non-recording Stations: Cincinatus Lincklaen

Norwich Norwich 1NE Sherburne

4N Sherburne Lincklaen . 22 • Cincinatus

Norwich

CANASAWACTA CREEK near Plymouth, NY



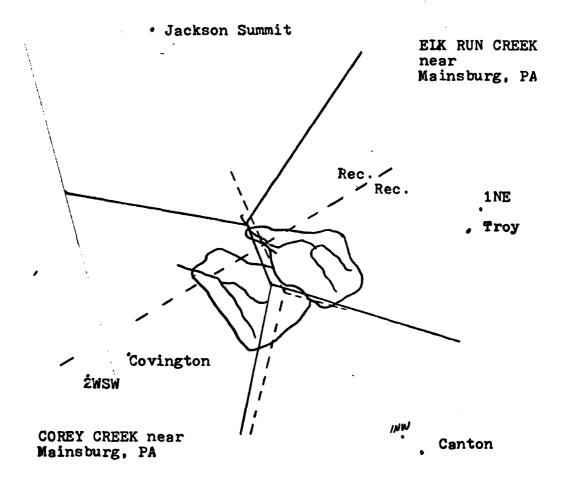
Recording Stations: Canton

Jackson Summit

Non-recording Stations: Canton

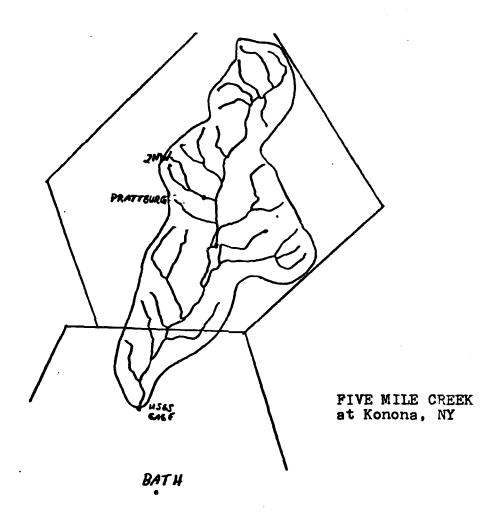
Covington

Troy

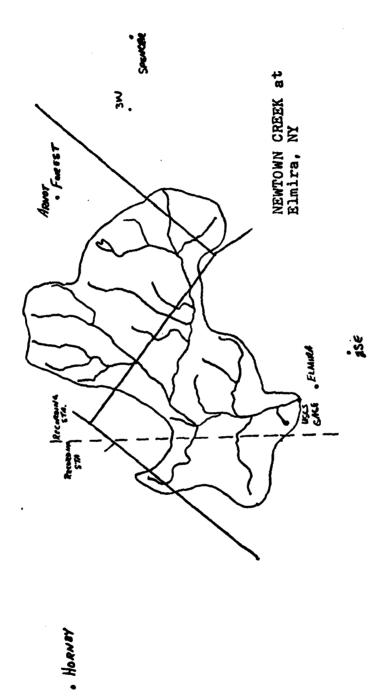


Recording Stations: None

Non-recording Stations: Bath Prattburg



Recording Stations: Arnot Forest Hornby Non-recording Stations: Elmira Spencer

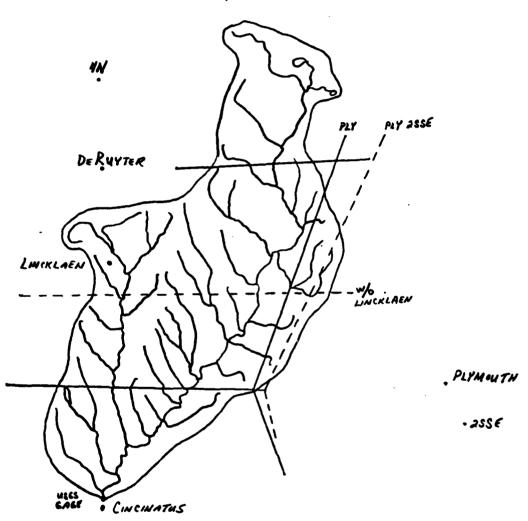


Recording Stations: Plym

Plymouth Plymouth 2SSE

Non-recording Sta.s: Cincinati

Cincinatus DeRuyter Lincklaen

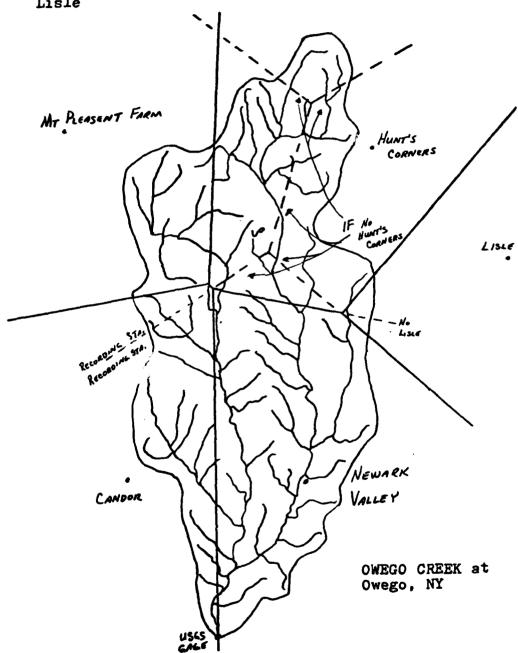


OTSELIC RIVER at Cincinatus, NY

# . Cortland

Recording Stations:
Hunt's Corners
Mt. Pleasent Farm
Newark Valley

Non-recording Stations: Candor Cortland Lisle



# Appendix B

# Input Card Deck Description

A brief line-by-line explanation of a typical card deck used in this study follows:

ID -- (Identification card) Serves to identify job by providing an opportunity to make comments or to describe an operation to be performed.

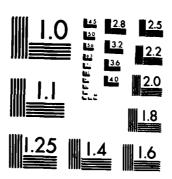
IT -- (Identify Time periods card) Serves to establish hydrograph time axis (60 minutes), start date of storm hydrograph (24 September 1977), start time of storm hydrograph (1300 hours), and number of hydrograph ordinates which must match the total number of flowrate entries on the QO cards.

10 -- (Identify Output formats) Serves to request supression or printing of selected output, calculations, or plots.

OU -- (Optimization limits for Unit hydrographs) The first number tells the program at which flood hydrograph ordinates optimization must be started. The second establishes the optimization ending ordinate.

PG -- (Precipitation Gage) Serves to identify rainfall gages. If a number follows the gage name, the gage is non-recording and the number is the total storm rainfall in inches. If no amount of rainfall is given, then the gage is a recording gage and must be followed by a PI card.

WATERSHED MODELLING IN THE CHEMUNG AND UPPER SUSQUEHANNA RIVER BASINS USI. (U) ARMY MILITARY PERSONNEL CENTER ALEXANDRIA VA J C STYRON 20 MAY 83 F/G 8/8 AD-A129 964 2/4 UNCLASSIFIED ΝL



MICROCOPY RESOLUTION TEST CHART NATIONAL BUREAU OF STANDARDS-1963-A

PI -- (Precipitation, Incremental) Serves to follow a recording rain gage PG card to describe the temporal rainfall distribution of the recording gage. The values are spaced at the time interval established on the IT card. The first value must represent the amount of rain falling in the first time period indicated on the IT card. The start of the rain must coincide with the start of the flood hydrograph flowrate on the QO cards.

KK -- (Kontrol Kard) Serves to identify the name of the stream gaging station at the basin outlet and to separate the precipitation data cards from the outflow data.

QO -- (Discharge (Q) Ordinate) Serves to input the observed hydrograph ordinates. The first flow value must match the first PI card entry as well as the IT card, i.e. all occur at the same instant.

PT -- (Precipitation Gage Total) Serves to identify those precipitation gages (either recording or nonrecording) which cover some fraction of the basin and contribute to the basin runoff.

PW -- (Precipitation Gage Weightings) Serves to provide the fraction of the basin covered by the gage above it. PR -- (Precipitation Recording) Serves to identify the recording stations which are used in deriving the temporal rainfall distribution pattern.

PW -- Same as the PW card above except it represents the

fraction of the basin covered only by recording stations.

BA -- (Basin Area) The first value is the drainage area

while the second is the mean annual rainfall for the

basin, if known.

BF -- (Base Flow) The values of STRTQ, QRCSN, and RTIOR respectively.

UC -- (Unit Hydrograph, Clark Method) The two parameters are the time of concentration (TC) and storage (R), respectively. If the numbers are positive, they remain constant throughout the calculations. If negative, the optimization routine begins with these values and then adjusts their values after removing the negative sign. If one or both are -1, then the optimization routine supplies an initial default value and then optimizes. In this model both parameters must be either positive or negative because of their interrelationship. LU -- (Loss Rate, Initial and Uniform) The two parameters are the initial loss (STRTL) and the constant loss rate (CNSTL), respectively. The same definitions for positive and negative values apply as above, except that both ... do not have to be positive or negative at the same time. ZZ -- (End of Data) HEC-1 begins calculations after all

data is read into its scratch files.

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		STATISTICS BASED (ORDINATES	ED ON UPTIMIZATION KEGIUN ES 35 THROUGH 60)	1710N REG	STATISTICS BASED ON UPTIMIZATION REGIUN (ORDINATES 35 THROUGH 60)		
NOS .		EQUIV DEPTH	MEAN FLOOR	TIME TO CENTER OF MASS	OF EQUIV MEAN CENTER C.M. TO PEAK TIME OF MS DEPTH FLOW OF MASS C.M. FLOW PEAK	PEAK	TIME OF PEAK
PRECIPITATION EXCESS		1.276		41.94			
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		STATISTICS BASED URDINATES	ASED ON FULL HYDRUGRAPH ES 1 THROUGH 1051	IYUKUGRAP (* 105)	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THROUGH 105)		:
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PRECIPITATION EXCESS	7L 0#3	1.276		46.14			YEAR
COMPUTED HYDROGRAPH DESERVED HYDROGRAPH	58674. 60239.	1.523	559.	58.64	16.70	3268.	52.00
DIFFERENCE PERCENT DIFFERENCE	-1565.	140*0-	-15.	-0.20	-0.20	-260.	00.0
STANDARD ERR	STANDARD ERROR	92.		AVERAGE	ABSULUTE ERROR	47.	

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	0100	17	0.30	2.00	3.40	25.	41.	•			1400	70	0.00	0.00	0.00	390.	471.
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	1500	Sa	3.01	0.01	J.03	24.	51.	•			1700	73	0.30	0.00	3.00	368-	414.
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# UC1	1 800	20	0.00	0.00	0.00	21.	59.	•			2300	/9	3.00	0.00	0.00	329.	343.
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9 UCT		53	3.02	3.32	0.00	26.	40.	•			9000	14	0.00	0.00	0.00	280.	244.
4 OCT	J200	34	0.01	10.0	0.00	29.	41.	•			0110	47	3.00	0.30	J. 90	265.	290.
Y UC!		35	3.03	U. 03	0.00	13.	43.	•			9400	de	0.40	0.00	0.00	277.	284.
y OCT		36	0.13	0.05	3.46	43.	47.	•			9900	84	0.00	0.40	9.00	272.	289.
9 001		3/	0.13	0.05	7.44	70.	74.	•			1000	90	3.00	J. VO	0.00	267-	274.
9 00.1		38	9.10	0.05	0.05	117.	45.	:			1100	91	3.00	U-00	0.00	262.	272.
9 02.1		40	0.10	0.05	U.05	185. 275.	104. 132.	:			1300	92 93	3.9J 3.93	0.UJ	4.00	257. 252.	266. 262.
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9 001	1000	42	U.14	U.05	0.15	532.	266.	•			1500	95	0.00	3.41	4.00	20.3.	255.
9 061		43	0.14	0.45	U-II	/12.	378.	•			1000	96	0.00	0.00	9.00	218.	251.
9 UC1		44	0.32	3.05	0.27	435.	544.	•			1300	47	0.00	9.00	4.00	234.	247.
9 00.1		•>	0.10	0.05	0.25	1212.	921.	•			L 400	94	0.00	Ø. JU	8.03	233.	244.
9 ULT		46	0.04	9.05	U-05	1733.	1538.	•			1400	44	3.00	0.00	9.00	225.	242.
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4 01.1		51	0.02	0.02	J. 00	2445.	3106.	•			0000	104	4.00	0.00	0.00	205.	226.
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		COMPARISON OF COM	CUMPUTED AND UBSERVED HYDRUGRAPHS	RVED HYDR	UGRAPHS		
		STATISTICS BASED	SED ON OPTIMIZATION REGION	ATION REG	STATISTICS BASED ON OPTIMIZATION REGION (ORDINATES 15 THROUGH 45)	•	
S		M OF EQUIV	MEAN	TIME TO CENTER OF MASS	TIME TO LAG MEAN CENTEK C.M. TO PEAK TIME OF FLOW OF MASS C.M. FLOW PEAK	PEAK	TIME OF
PRECIPITATION EXCESS		2.245		22.51			
CUMPUTED HYDRUGRAPH OBSERVED HYDROGRAPH	63347. 58575.	1.644	2043.	33.12	11.21	3511.	34.00
DIFFERENCE Percent difference	4772. 8.15	0.124	154.	-0.56	-0.56	-181-	1.00
STANDAKU OBJECTIVE FUI	4	281. 254.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERRÜR ABSOLUTE ERROR	229. 32.81	
STATISTICS BASED ON FULL HYDKUGRAPH (OKDÍNATES 1 THROUGH 101)		STATISTICS BASED LOKDINATES	3ASED ON FULL HYDKUG FES I THROUGH 101)	ON FULL HYDKUGRAPH 1 THROUGH 101)	STATISTICS BASED ON FULL HYDRUGRAPH I ORDINATES I THROUGH 101)		
	PO MUS	V1008	Z	TIME TO	LAG C.M. TO	PEAK	TINE OF
	FLOWS	ОЕРТН	FLOM	OF MASS	. W.	FLOW	PEAK
PRECIPITATION EXCESS		2.245		22.51			
CUMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	95982. 87000.	2.491 2.258	950. 861.	42.39	19.88	3511.	34.00
DIFFERENCE Percent difference	8982. 10.32	0.233	.68	-1.21	-1.21	-181.	1.00
STANDARD	STANDARD ERROR	238.		AVERAGE	ABSOLUTE ERROR	160.	

				- uni	1 MADEOFET	PH			
				#1 EMD-0	F-PERIUD D	AULMATES			
St.	141.	344.	414.	678.	4477.	1+30.	1044.	1740.	1468.
1901.	1041.	1676.	1579.	1471.	1370.	1270.		1107.	1431.
W. J.	444.	432.	115.	122.	ale.		543.	543.	sue.
471.	419.	w.	361.	154.	IN.	107.	266.	267.	248.
231.	21>.	201.	lef.	174.	104.	151.	141.	111-	122.
114.	104.	46.	92.	45.	eu.	74.	64.	64.	aJ.
30.	<b>52</b> ·	44.	45.	42.	17.	50.	14.	12.	24.
21.	25.	24.	22.	21.	19.	10.	17.	15.	14.

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NA DE	f 0130		۵.03	3.00	0.30	144.	144.	•	22		8438	52	0.00	ىد.د	J.00	1211.	735
	1 0230	ž	0.00	3.00	u.J0	145.	1.8.		22	120	3503	53	0.00	0.00	0.00	1131.	703
	4300	š	U. JO	9.40	J. JO	143.	144.				4400	54	J. 00	0.33	0.00	10>0.	- 14
N OC	CO+C 1	•	3.00	4.00	U.00	140.	140.	•	22	úĊΓ	9/00	55	3.04	٠. ٠٠	u.00	480.	641
10 OC	1 4500	5	0.04	0.00	9.30	157.	140.	• ,	22	OC 1	9800	54	0.40	J. 0J	9.00	921.	624
	7 0600	•	J. 04	0.00	9.33	135.	145.				4403	5/	J.90	0.00	0.00	<b>86</b> 6.	+01
	9700		J.34	0.00	<b>u. 6</b> 0	132.	145.				1000	34	0.00	0.00	0.00	804.	581
	1 2900		J.00	J. 10	0.00	130.	145.				1100	59	J. 00	4.60	J. 04	751.	564
	1 0400	•	u.00	J. 00	4.43	127.	143.				1240	••	0.00	0.00	٠.٠٠	702.	541
	1 1000	IJ	3.00	3.00	<b>4.0</b> 3	125.	Į45.				1 107	*1	0.00	3.30	0.00	654.	534
	1100	11	0.01	0.00	J.01	123.	1.44.				1400	•4	0.00	0.00	0.00	413.	513
	1500	1.5	3.09	3.00	0.09	127.	148.				1 200	•3	0.00	0.00	U. UO	574.	501
	1 100	13	0.03	0.00	3.01	140.	153.				1000	•	0.00	0.00	٠.٥٠	\$37.	481
	1+00	14	J. 04	0.00	3.04	165.	158.				1700	•5	0.00	0.00	0.00	502.	479
	1500	15	0.03	0.00	0.03	194.	144.				1 400	••	0.00	u.00	0.00	• 70 -	461
	1600	l e	9.04	0.00	0.04	245.	177.				1400	67	0.00	0.00	0.00	****	450
	1 100	17	3.00	0.00	3. 16	100.	190.				1000	44	0.00	0.00	<b>v.o</b> o	411-	434 432
	1 1900	10	0.12	0.00	0.12	304.	200.				\$100	• 4	0.00	0	0.00	345. 341.	420
	1 1400	14	0.14	0.00	9.14	484.	237.				5500	70	J. 04	J. UJ	0.00	339.	41
	2000	20	0.20	0.00	0.20	415.	284.				2300	71 72	0.00	0.00	0.00	333.	40
	5100	<b>21</b>	0.09	0.00	3.09	776.	355.				4000					347.	393
	2200	22	0.04	0.00	0.08	957. 1149.	475.				3100	73 74	0.00	0.00 J.03	U.00	321.	387
	2300	23		3.33	0.04	1347.	636. 608.				0300	15	J. 60	3.00	0.00	315.	340
		25	0.04 0.30		J. 30	1>5.	1034.				U40U	76	0.00	0.00	0.00	307.	37
	1 1100		3.36	0.00		1744.	1347.				0503	77	0.00	0.00	J.00	303.	360
	u300	26 21	0.20	0.00	3.3s	2041.	1754.				0900	78	3.00	3.33	0.00	297.	34.
	9-00	2.	0.10	0.00	J. 10	2312.	2000.				9100	79	0.00	0.00	4.00	245.	354
	U500	29	0.09	0.00	J.09	2562.	2522.				U800	60	0.00	0.00	3.00	266.	350
	Jedd	30	0.04	0.00	J.44	₹636.	2950.				9490	91	0.00	0.00	0.00	201.	345
	9700	31	U.00	0.00	J.JJ	3070.	3448.				1000	44	0.00	0.00	0.00	216.	340
	9400	52	0.33	3.00	4.00	3203.	3502.				1100	- 63	3.00	0.00	8.00	271.	334
	9900	33	3.00	0.00	0.00	3406.	3622.				1200	*	0.00	4.55	0.00	205.	332
	1000	34	0.00	0.00	0.00	3491.	3692.				1300	45	U-00	0.00	J. 00	200.	320
	1100	35	0.00	0.40	J.00	3511.	1650.				1400	**	0.00	0.30	0.00	256.	324
	1200	34	0.00	J. JJ	0.00	3454.	3515.				1500		0.00	0.00	0.00	251.	320
	1100	57	0.30	3.03	0.00	3327.	3294.				1000	44	y.00	0.03	U. UQ	246.	114
	1400	36	U.30	3.30	0.00	3157.	3100.				1700	84	0.00	U. 0U	0.00	241.	312
	1500	34	3.00	0.00	0.33	2971.	2647.				1600	90	0.00	9.00	0.00	217.	310
	1600	43	0.00	0.00	3.03	etet.	2496.				1900	*1	0.00	0.00	U.00	213.	300
	1700	41	u.00	0.00	v.00	2595.	2522.				2000	ŸŽ.	J. 0U	0.00	0.00	226.	302
	1800	42	3.00	0.00	0.00	2420.	2313.				2100	93	J. JO	4.00	0.00	224.	290
LUCI		•5	J.30	0.00	3.30	2251.	2050.				2200	94	0.00	0.00	0.00	220.	244
i uci		**	v.00	0.00	J.00	2104.	1745.				2300	95	0.00	0.00	0.00	216-	240
LOCI		45	U.UU	0.00	0.00	1764.	1461.				8000	90	J.00	0.00	J. JO	212.	284
	2230	40	U.00	0.00	J.0J	1033.	1233.				9100	47	0.00	9.30	0.00	204.	204
	2300	•1	0.00	0.00	J.J0	1710.	toed.				9200	48	6.00	y. 00	0.00	204.	282
	3000	44	J.00	0.00	0.00	1596.	957.				0300	94	J. JJ	9.00	0.30	200.	270
	3100	44	3.00	J. JO	3.00	1444.	677.				0440	103	0.00	0.00	0.00	190.	276
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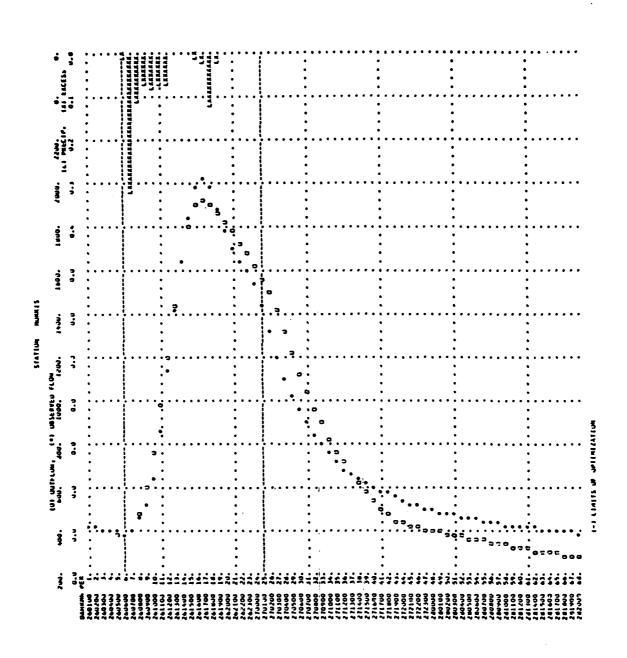
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CUMPAKISUN UF COMPUTED HYDRUGRAPHS	**********	CUMPAKISUN OF COM	COMPUTED AND DBSERVED HYDRUGRAPHS	ERVED HYDR	UGRAPHS ***********	• • • • • • • • • • • • • • • • • • • •	•
		STATISTICS BASED LORDINATES	SED UN OPTIMIZAT	OPTIMIZATION REGION 6 THROUGH 25)	NOI		
:		EQUIV	MEAN	TIME TO CENTER	TIME TO LAG CAM TIME DE	9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	TEME OF
	FLOWS	ОЕРТН	FLOW	OF MASS		FLOW	
PRECIPITATION EXCESS		0.912		10.34			
CUMPUTEU HYOROGRAPH UBSERVED HYDRUGRAPH	27459.	0.713	1373.	17.36	7.02	1928.	16.00
DIFFERENCE PERCENT DIFFERENCE	579. 2.15	0.015	29.	-0.00	00.0-	-42.	00.0
STAND	STANDARU ERROR UBJECTIVE FUNCTION	70.	AVERAGE	AV ERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE EKROR	59• 5•22	
	•	STATISTICS BASI (URDINATES	BASED ON FULL HY TES 1 THROUGH	ON FULL HYDROGRAPH I THROUGH 681	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THROUGH 68)	**************************************	
	SUN OF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK FLOM	TIME OF PEAK
PRECIPITATION EXCESS		0.912		10.34		:	:
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	53767.	1.396	791. 803.	27.33	16.99 18.45	1928.	16.00
OIFFERENCE PERCENT DIFFERENCE	-841.	-0.022	-12.	-1.46	-1.46	-92	00.0
STANDARD ERROR	ARD ERROR	95.	24 60 00 00 00 00 00 00 00 00 00 00 00 00	AVERAGE	ABSOLUTE ERROR	.18	

	126.		2023- 781- 302- 117- 45-	210.		1221 (134) (125) 587 534 686. 227 206. 186. 44 80. 125.	200.	137-	22. 22. 22.	25. 141. 170. 170. 170. 170. 170. 170. 170. 17	272. 1009. 1009. 1009. 200. 200.	•			
					NAH	HYDAUGKAPH AI STATION	**************************************	I	MORRES						
A MUN HRM	045	KAIN	LUSS	A MUN HAPIN UND MAIN LUSS EALESS	7 4403	7 595		* 43	ACK HAN	MUN MRAN GRO		\$67	EKLESS	RAIN LOSS EKLESS LOMP & COS	580
26 SEP UI	3	0.00	0.00	00.0	420.	.20.	••	3 27		55	9.0	9.0	90.9	7.57.	1
26 SEP 02	3	0.00	20.0	0.00	412.		•	27		*	9.0	3.0	00.0	713.	3
26 SFF 0300 28 SFF 0400	23	90.0	300	200	*04 ***********************************		• •	∑ 27	SEP 1300		900	90.0	36	•	
26 SEP UN	<b>.</b>	3	9	3	389.	340.	•				8	3	3	2/8	3
26 SEP 06	9	20.0	00.0	7.0	363.	387	•				9	9	8	\$40.	3
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26 SEP 20			9.0	3	1824.	1781.	•	 			00.0	3	3	359.	**
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		CUMPAKISUN UF CUMPUIED AND UBSEKVED HYDKUGKAPHS	UIED AND UBSE	RVED HYDKOG	CDMPARISON OF COMPUTED AND OBSERVED HYDROGRAPHS	*********	*******
		STATISTICS BASED (ORDINATES	ED UN UPTIMIZATION TES 20 THROUGH 50)	UN UPTIMIZATION REGION 20 THROUGH 50)	STATISTICS BASED UN UPTIMIZATION REGION (ORDINATES 20 THROUGH 50)		
		EQU [V DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	C.H. TO	PEAK FLOW	TINE OF PEAK
PRECIPITATION EXCESS		2.816		28.44			
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DIFFERENCE PERCENT DIFFERENCE	-5918.	-0-154	-161-	15.0-	-0.57	-1.05	1.00
STAND OBJECTIVE	STANDARD ERROR OBJECTIVE FUNCTION	472.	AVERAGE	AVERAGE AB	ABSOLUTE ERROR ABSOLUTE ERROR	354.	
STATISTICS BASED UN FULL HYDRUGRAPH (URDINATES 1 THROUGH 86)		STATISTICS BASEU (URDINATES	IASEU UN FULL HYG TES 1 THROUGH	UN FULL HYDRUGRAPH 1 THROUGH 86)			
	SUM OF FLOWS	EQUIV	MEAN	TIME TO CENTER OF MASS		PEAK	TINE OF
PRECIPITATION EXCESS	:	2.816		28.44			
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DIFFERENCE PERCENT DIFFERENCE	-34041.	-0-884	-396.	-3.70	-3.70	-63.	00.1
STAND	STANDARD ERROR	662.		AVERAGE	ABSOLUTE ERROR	460.	

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to UC			0.00	3.00	0.30	249.	282.	•			2300	47	0.00	0.00	0.00	3360.	4105-
to UC		,	0.00	0.00	0.00	284.	274.	•			0000	4.6	0.00	0.00	0.00	2991.	1902.
In UC		•	0.00	3.30	0.00	274.	270.	•	Lo	OC T	4100	44	U. DO	0.00	U.00	2000.	3717-
IS OL		7	0.00	9.30	J. 90	273.	244.		10	UCT	U20U	50	0.00	0.00	U.00	2243.	3531.
In UL	0600		0.00	0.00	0.00	244.	202.	•	10	OC T	J3UQ	51	V. 00	0.00	0.00	1713.	3348.
16 UC1	U900		0.00	9.00	0.00	263.	254.	•			0440	54	0.00	0.03	0.00	1415-	\$172-
to OCI	1,000	LO	0.30	3.00	0.00	254.	257.	•			0500	5.5	U.00	0.00	0.30	1351.	2989.
16 JL		11	3.02	9.00	0.42	255.	253.	•			9400	54	0.00	0.00	0.00	1122.	2813.
LA UCI		12	0.01	0.00	0.01	255.	252.	•			0700	55	0.00	0.04	0.00	450.	2443.
to OUT		1.5	0.02	9.04	4.42	201.	2>2.	•			0400	56	0.04	0	0.00	841.	2483.
14 OC		14	0-01	<b>J.3</b> 0	0.01	271.	250.	•			0400	51	0.00	0.00	0.00	825.	4319.
14 (16)		15	9-00	ەن.بە	0.90	245.	248.	•			1000	54	0.00	0.00	0.00	aro.	2179.
14 OC		10	3.00	0.00	0.00	300.	244.	•			1100	54	0.00	0.00	0.00	195.	2034-
16 UCI		1.7	4.00	0.00	0.00	310.	246.	•			1200	•0	0.00	0.00	0.00	/80.	1104.
14 IX		18	0.00	3.30	0.00	331.	245.	•			1300	44	0.00	0.00	0.00	765.	1491.
16 OL		(9	4.31	0.00	0.01	345.	245.	•			1400	62	U.00	0.00	9.60	751.	1563.
14 OC1		50	3.00	0.00	0.30	356.	243.	:			1500	43	J-00	0.00	0.00	737. 723.	1501.
14 OC		21	0.00	0.00	0.00	36/-	2+3.				1000	• •	J.00		0.00	710.	1440.
14 OCT		5,5	0.10	0.00	0.10	341.	243. 250.	:			1700	65	0.00	0.00	9.00	474.	1377.
LO UCT		23	U. 30	0.00	0.10	•33.	242.	:			1400	<b>66</b>	0.00	0.00	0.00	603.	1325.
		24 25	0.33	0.00	0.33	561. 784.	242. 244.				2000		0.00	0.00	9.00	670.	1277.
17 001		26	0.23	0.00	0.23	1071.	584.	:			2100	69	0.00	0.00	0.00	458.	1237.
17 901		21	3.19	0.00	3.19	1+24.	1041.				2200	70	0.00	0.00	0.00	445.	1194.
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ir out		11	0.00	3.00	0.05	3477.	. 222 د	•			0200	74	0.00	0.00	0.00	598.	1058.
17 OC		32	3.19	0.00	0.19	4018.	4465.	·			0300	75	0.00	9.00	0.00	56 7 .	Lez).
17 UC		ii	0.17	0.00	0.15	4527.	5100.	·			3400	76	0.00	0.00	0.00	570.	995.
LF OCT		34	3.12	J.50	0.12	4990.	5020.	·			J>00	17	0.00	0.00	0.00	565.	768.
IT UC		35	0.11	0.00	0.11	5384.	5920.	•			4600	74	U. 00	0.03	U-00	555.	942.
LT UCI		36	0.00	0.00	J.08	5097.	> 940.	•			J/03	77	0.00	0.00	9.00	544.	ATO.
IT OCI		17	J. U4	0.30	0.04	5877.	5480.	•			6400	80	J.00	0.00	0.00	534.	490.
L7 UC		14	0.08	0.00	U-05	5917.	5800.	•			J900	Ø L	0.00	0.00	0.00	524.	<b>465.</b>
LT UCI	1 500	34	v. 35	3.00	0.05	5452.	5/40.	•	19	UC T	1000	82	0.00	0.00	0.00	514.	847.
17 IL1		40	0.02	0.00	U.02	5714.	1000.	•			1100	83	0.00	0.03	v.01	505.	423.
ET UC		41	0.01	3.30	J. 01	>>00.	5400.	•	19	JCT	1200	84	0.00	0.00	0.00	495.	505.
17 JL1		42	4.30	9.00	J.30	5231.	5220.	•			1 100	85	0.00	J.00	U. 00	444.	793.
II OLI	5 900	4.5	0.00	0.00	0.00	4400.	4 440.	•	19	JL I	1-40	86	0.00	0.30	0.40	417.	761.
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15	7	25	57	67	\$7	30	30	15	13	34	42	
91	3	96	9 <b>9</b>	40	740	55	100	791	707	552	37.1	
11	2	526	980	1632	1506	1344	1210	1000	740	162	717	
87	3	1+9	404	564	540	715	164	404	470	754	740	
<u>6</u> 1	3	452	41°	366	374	394	C14	478	452	470	670	
22	o o	404	450	446	434	414	*0*	374	429	450	714	
77	3	538	131	068	1000	1120	7811	1072	980	006	430	
77	9	104	710	65°	620	500	550	250	215	164	411	
23	ဂ္ဂ	45 A	440	434	458	975	014	7.47	393	383	1115	
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52	P T	PLYA										
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32	3;	<del>;</del>	05									
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STATISTICS BASED ON UPTIMIZATIUN REGIUN (OKDINATES 14 THKUUGH 30)		STATISTICS BASED (OKDINATES	SED ON UPTINIZATION KEGIUN TES 14 THKOUGH 301	A11UN KEL H 301	ומא		
	SUM OF FLOWS	EqUIV UEPTH	MEAN	TIME TO CENTER UF MASS	C.H. TC	PEAK FLOW	TIME OF
PRECIPITATION EXCESS		0.932		30.78			
COMPUTEU HYDRUGRAPH GBSERVED HYDRUGRAPH	11992.	0.321 0.320	705. 702.	25.09	-11.69	1532.	23.00
DIFFERENCE PERCENT DIFFERENCE	53.	100.0	*°	0.51	0.51	-100-	1.30
STANDAR UBJECTIVE F	STANDARD ERROR UBJECTIVE FUNCTION	126.	AVERAGE	AVEKAĞE PEKCEN I	ABSULUTE ERRUR ABSULUTE ERRUR	104.	
		STATISTICS BASEU (ORDINATES	JASEU ON FULL HY TES I THROUGH	ON FULL HYDKUGKAPH I THWÜUGH 93)	STATISTICS BASED ON FULL HYDROGKAPH (ORDINATES 1 THROUGH 93)		
		ELUIV	MEAN	TIME TO CENTER OF MASS	C.H. TO	PEAR FLUH	TIME OF PEAK
PRECIPITATION EXCESS		0.932		36.78			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	40369. 46614.	1.080	434. 501.	49.62 52.00	13.05	1532.	43.00
DIFFERENCE PERCENT DIFFERENCE	-6245. -13.40	-0.167	-67.	-2.17	-2-17	-100-	1.00
STANG	STANDARD ERROR	112.		AVERAGE	ABSULUTE ERKUR	92.	

					I HYDRUGAL				
				S4 ENJ-0	F-PEKELL U	HUINATES			
244.	1114-	2115.	204).	3305.	2142.	2473.	44 30 .	2011.	1010.
1 - 54 -	1075.	1310.	1233.	1042-	7/6.	443.	179.	716.	•45.
542.	525.	413.	+27.	363.	147.	31.1.	404.	2>>.	210.
201.	Lu7.	ite.	152.	111.	144.	111.	1.0.	41.	82.
14.	44.	40.	54.	44.	•••	•3.		32.	27.
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104 1407	A.	1.00	9.00	0.00	2>.	25.	•	21	3.3%	1333	**	J. 60	3.00	4.30	420.	454
TON 1200	2	3.02	3.02	9.39	24.	25.	•	21	334	1.03	4.	3.30	0.00	4.00	321.	410
JUN 1003	,	3- 95	0.02	0.33	24.	54.	•	21	JUH	1503	50	0.41	0.01	3.00	313.	47
JUN 1/03	•	3.34	9.04	0.00	23.	29.	•			Levi	21	3.01	J.ul	4.30	300.	**
104 1403	•	3. 32	3.62	0.55	23.	30.	•			1197	25	3.03	٥. ٥٥	ددرد	244.	4>
נניון הטנ	•	3.02	0.02	3.00	24.	30.	•			1400	>3	0.00	3.00	0.00	242.	44
TAM SOOD		3.31	3.01	0.01	22.	31.	•			ミャンリ	>4	J. JJ	0.00	2.30	245.	• 1
TON STAD	•	3.41	9.21	9.53	Łį.	33.	•			2000	>>	3.00	0.33	0.30	216.	• 1
104 2233		3.32	0.02	9.00	41.	34.	•			2200	>=	2.02	0.33	3.33	271.	•0
104 5 103	ro	3.01	0.01	0.00	23.	35.	•			2200	51	0.04	0.34	4-30	205.	
100 TO13	11	7. Of	3.31	4.00	20.	34.	•	21	300	5370	74	2.24	0.00	0.14	254.	• •
Jui: 3133	15	3.63	3.33	3.30	14.	30.	•			<b>0</b> 000	3.4	U. U.	3.00	4.10	302.	• :
704 0533	1.1	3.42	3.02	9.33	Įy.	43.	•			21 20	نان	9.3-	J.U4	4	411.	5
JUN (133)	14	3.18	3. 18	3.33	id.	40.	•			ひとりこ		J. US	U.U5	3.33		50
JUN 3403	15	3.39	3, 39	4.33	i.e.	53.	•			2700	44	J.24	7.00	4.16	673.	7.
104 0200	16	3.12	9.12	4.00	Lø.	100.	•			~))		U.U-	3.04	9.10	lot.	•
Jun 3633	1.7	3.00	0.03	3.33	17.	162.	•	2.	JJN	7237		3.34	3.02	4.03	#¥4.	13
COIC MUL	1.	ē. 01	2.03	1.33	17.	202-	•	2.	ようか	3603	• >	3.00	0.00	3.33	982.	11.
104 1003	15	3.20	3.15	9.1-	>7.	2>>-	•	20	JJN	0/33		3.30	÷.00	0.30	429.	41
104 2477	13	3. +2	0.06	3.36	275.	<i>317</i> .	•	20	304	3430	. 1	3.30	0.33	3.33	e71.	10
170 1033	21	2.04	J. Je	0.32	713.	>24.	•	28	334	رن بحق		0.00	U.36	0.00	740.	94
IN IIIO	22	3.02	0.02	Q.00	1200.	740.	•	4.	JJie	1000	. 7	0.33	4.00	3.03	739.	4
100 1403	43	3.02	9.02	3.39	1504.	1032.	•	20	3.0%	1100	13	0.63	3.00	3.33	•4D.	
ini i jos	24	3.32	9.02	3.00	1532.	L500.	•	20	J.3%	1200	71	4.00	J.03	3. 30	574.	- 7
ine tim	25	J. J4	3.04	3.30	1433.	1394.	•	24	JUN	1200	12	0.00	3.00	3.33	541.	7
1303	24	2.31	0. 31	3.30	1267.	1216.	•	45	JUL	LAJJ	13	J. J.	3.03	3-30	446.	
14% 1473	21	3.02	0.32	a. 33	1144.	1360.	•	24	JUM	LáJú	1-	u. 33	J.00	0.33	421.	•
coll me	24	3.61	7.01	0.30	1932.	423.	•	24	JJN.	Level	75	0.00	J.JJ	0.00	467.	5
104 [47]	2+	J. JE	J. 01	3.33	932.	741.	•	2.	100	1730	10	3.00	3.00	3	404.	5
CEF1 NO	30	J. JL	0.01	3.30	841.	719.	•			1600	11	0.00	0.00	0.30	394.	5.
104 5303	31	3.33	2.03	30	leu.	647.	•			1477	16	3.00	0.00	3.33	109.	Ś
IN 5103	25	3.31	J. 01	3.33	484.	604.	•			2300	14	3.00	1.00	3.33	J#2.	
CC55 #W	33	3.31	0.01	0.30	e23.	367.	•			2130	40	3.30	3.00	0.33	371.	·
IUM 2300	34	3.00	0.00	J.00	560.	540.	•			2230	• i	3.33	3.00	3.33	Je2.	•
Je 8303	35	J. 31	0.01	0.30	504.	512.	•			2300	• 2	0.00	0.00	0.33	353.	- 7
OF 101	36	3.33	3.00	3.93	4>7.	441.	•			0333	43	J. JU	4.65	4-03	342.	
COSC MU	31	J.00	9.00	0.60	+24.	404.	•			0103		3.00	3.30	3.33	337.	
COLE MOI	38	3.33	0.00	0.00	·le.	410.	•			1433	45	3.03	3.33	3.33	327.	•
ium Jado	37	3.00	4.30	3.30	400.	452.				3300	8 11	0.30	4.00	3.33	321.	- 3
JM 0500	• 3	3. 33	0.00	0.00	341.	4.5.	•			444	•/	3.30	رن.د	3.33	314.	31
104 3603	*1	3.00	9.30	0.33	34 T.	422.	•			3503	•	3.30	1.03	3.43	307.	35
tote au	42	3.33	0.00	0.00	374.	+10.				J-00		3.00	9.60	0.33	299.	31
0440	+3	3.40	6.00	0.00	310.	349.	·			3/33	7.	1.33	4.00	2-47	292.	31
UK 3933	44	3.13	0.30	3.04	301.	344.	•			3430	7.	3.03			200.	31
W 13.J	*>	3.0/	0.06	0.01	352.	394.				3900	*;		0.33	3.03		
UN 1135	10	3.63	0.03	J. 00	344.	413.	-			1003	77	2-02	4.00	1.33	214.	34 34
m 1500	47	3.63	3.30	3.20	336.	424.	Ξ			.400	7,	. J.	2.03	3.33	212.	

PEAR FLUE	1140			HARINUM AVE	RAGE FLUG	
16121	( etec )		6 7 9 SEL	24-1M	12-m4	42.JU-14
4312.	21.03	16121	1120.	701.	547.	-31.
		CENCHESE	3.213	0.449	1.354	1.0/0
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I-I LIMITS OF COTINIZATION

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7	3	Š	IT GRAPH	(CLARK)	AND LOSS	KATE	CONTRUKA	OPT IMI Z	MILON				
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	و. د	PLYM											
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12	3	255	260	285	315	199	505	550	101	155	910		
1	9	018	950	930	850	1000	1192	1408	1664	1792	1949		
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91	3	იოგ	900	850	900	164	611	769	609	656	639		
11	3	612	580	564	543	540	533	975	215	27.5	114		
19	3	27.4	466	424	755	430	774	410	404	399	388		
67	3	311	360	372	372	366	360	၁	၁	0	၁		
22	P	PLYM											
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		(UKDINATES	SED ON UPTIMIZATION REGIUN TES 1 THROUGH 31)	31) H	•		
	\$ 2 K	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	TIME TO LAG  DEPTH FLUM UF MASS C.M. FLOW PEAK TIME OF	PEAK	TIME OF
PRECIPITATION EXCESS		1.130		15.97			
COMPUTED HYDRUGRAPH UBSERVED HYDRUGRAPH	35348.	0.946	1140.	18.49	2.52	2010-	21.00
DIFFERENCE PERCENT DIFFERENCE	1521.	1,0.0	.64	-0.85	-0.85	-46.	0.00
STANDARO UBJECTÍVE FUI		183. 168.	AVERAGE	AVERAGE AB	ABSULUTE ERROR ABSULUTE ERROR	132. 18.97	
		STATISTICS BASS (URDINATES	STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES I THROUGH 66)	HYÖRUGRAPH ih 66.1	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THROUGH 66)		;
:		EQUIV UEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.M. TO C.M.	PEAK	TIME OF PEAK
PRECIPITATION EXCESS		1.130		16.91			
COMPUTED HYDROGRAPH DBSERVED HYDROGRAPH	53863.	1.442	816.	28.29	12.31	2010.	21.00
DIFFERENCE PERCENT DIFFERENCE	1452.	0.039	22•	-0.63	-0.63	-46.	0.00
STANDARD ERROR	ARD ERROR	131.		AVERAGE AB	ABSOLUTE ENRUR	.70	1

44. 1914 1 113. 1074 1 103. 103. 153. 153. 153. 153. 153. 153. 153. 15		;	7574		2457	54 END-UF-F	ERI 10	7 4 10 M	2	-	774		404		
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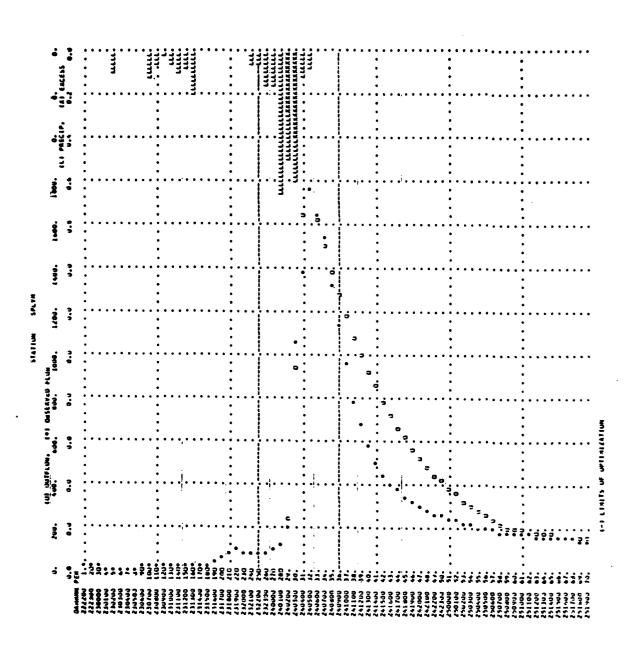
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<b>~</b> :	<b>3</b> 3	72.	171	177	164	191	158	155	751	146	143
2 5	)  -  -	NA Id			, ,						
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	U .	ARISON UF CUMPUTE  ***********************************	COMPARISON OF CUMPUTED AND OBSERVED HYDROCKAPHS  ***********************************	WED HTDR Gessess Tion Reg 1 36)			
SUM OI	•	EQUIV DEPTH	TIME TO EQUIV MEAN CENTER DEPTH FLOW OF MASS	TIME TO CENTER OF MASS	*	LAG C.M. TU PEAK TIME OF C.M. FLOM PEAK	TIME OF
PRECIPITATION EXCESS		0.679		29.59			
CUMPUTED HYDRUGAAPH OBŠERVED HYDROGRAPH	10366.	0.277	864.	32.92	3.32 3.10	1759.	31.00
DIFFERENCE PERCENT DIFFERENCE	-55.	-0.001	-5-	0.23	0.23	-1.	00.0
STAN! OBJECTIVI	STANDARD ERROR OBJECTIVE FUNCTION	105.	AVERAGE	AVERAGE Percent		84. 37.06	
		STATISTICS BASED (URDINATES	STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES 1 THROUGH 70)	IYDRUGRAP   701			
	SUM OF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.M. TU C.M.	PEAK FLOM	TIME OF PEAK
PRECIPITATION EXCESS		0.679		29.59			
COMPUTEU HYDRUGRAPH UBSERVED HYDRUGRAPH	25268. 21214.	0.676	361.	41.11	11.52	1759.	31.00
DIFFERENCE PERCENT DIFFERENCE	4054. 19.11	801.0	58•	1.40	14.54	-1-	00.00
STANE	STANDARD ERROR	135.		AVERAGE	ABSOLUTE ERROR	80.	

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10						HEC-1 INPUT	4PUT						PAGE
10   CANASANACTA CREEK   f-31-71   10   UNIT GRAPH (CLARK) AND LUSS KATE (UNIFORM) GPTIMIZATION   10   1   2   2   2   2   2   2   2   2   2	LINF	. e.				******			/	20		10	
10	-	01	3	ANASAWAC TI	A CREEK	1-31-7	•						
1	~	2	ā	NIT CRAPH	(CLAKK)	AND LUSS	KATE C	CNIFORMS	77 INILAD	NOTE			
10	•	_	3	19 JUL 91	0000	9							
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PI .01 .52 .05 .05 .12 .06 .01 .01 .01 .01 .00 .00 .00 .00 .00 .00	•	S	PLYM								;	:	
1	_	<u>a</u>	10.	.52	50.	50.	-12	ş.	70.	.0	ē.	0	
No.	10	<u>.</u>	30.	ۍ. د	00.	90.	00.	3.	90.	3	3.	00.	
KK SPLYM  VI .13 .59 .48 .12 .01 .00 .00 .00 .00 .00 .00  FI .13 .59 .48 .12 .01 .00 .00 .00 .00 .00 .00  KK SPLYM  VI .14 .15 .15 .15 .15 .15 .15 .15 .15 .10 .10 .00  VI .15 .15 .15 .15 .15 .15 .15 .15 .15 .15	~	ā	· .	٥٥.	90.	00.	3.	00.	90.	00.	8	00.	
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40     266     198     173     149     137     122     112     102     94       40     48     40     41     39     36     36     35     35       40     48     42     41     39     34     36     35     35       40     43     35     60     92     376     1000     1345     1075       40     131     123     612     516     430     162     153     149     143     137       40     131     128     122     120     117     115     107     107     104       PA     1,00     1,00     1,00     1,00     1,00     1,00     107     104       BA     57,9     0,0     1,024     1,00     1,00     1,00     1,00     1,00       LU     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0       LU     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0     -1,0	13	3	•	•	21	12	77	33	98 7	516	418	335	
40     81     74     70     66     62     60     56     51       40     48     46     42     41     34     36     36     35     35       40     48     46     42     41     370     340     1345     1075       40     723     612     516     430     370     204     264     262     239       40     1.31     128     122     120     117     115     107     107     104       PA     1.00     1.00     1.00     1.00     1.00     1.00     1.00     1.00     1.00       BA     57.9     0.     -4.     -10.     -10.     -10.     -10.     -10.       LU     -1.     -1.     -1.     -1.     -1.     -1.     -1.	*1	3	790	198	173	6+1	137	771	711	701	<b>*</b>	80	
40     48     46     42     41     39     36     35       40     48     46     42     376     1000     1162     1345     1075       40     723     612     516     430     370     244     262     239       40     202     190     180     169     169     149     149     131       97     97     97     122     120     117     115     109     107     104       94     1-00     96     1-00     97     1-02     102     104     107     104       95     1-00     96     1-02     1-02     1-02     1-02     1-02     1-02       10     -1-0     -1-0     -1-0     -1-0     -1-0     -1-0     -1-0     -1-0	٤2	Ş	81	42	20	99	79	ဒ္	50	*	21	74	
40     13     35     60     92     376     1000     1162     1345     1075       40     723     612     516     430     370     246     262     239       40     202     190     180     169     162     155     149     143     137       40     131     122     122     120     117     115     109     107     104       PA     1,00     0	2	3	84	40	45	14	34	26	36	35	35	33	
JO 723 612 516 430 470 320 244 262 239 JU 202 190 180 169 162 155 149 143 137 JU 202 190 180 169 162 155 149 143 137 JU 202 190 180 169 162 100 107 106 JU 1-100 JU -110- LU -11-	11	3	13	35	9	45	3.76	0001	7911	1345	1075	920	
JU 202 190 180 164 162 155 149 143 137 JU 131 128 122 120 117 115 104 107 104 PM 1.00 PM 1.00 PM 57.9 U. UC -410. LU -11.	82	<u>و</u>	723	915	516	430	370	370	594	262	239	R17	
130 131 128 122 120 117 115 109 107 104 107 109 100 110 1100 1100 1100 1100 110	67	3	202	190	180	164	162	551	651	143	137	134	
PT PLY3 P4 1.00 P4 1.00 B4 57.9 U. BF 57.9 U. C -410. LU -11.	<b>??</b>	Š	171	128	771	170	111	511	707	101	104	707	
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PA PLYN PA 1.00 BA 57.9 BF 57.9 UC -410. LU -11.	77	ď	1.00										
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		STATISTICS BASEU (URDINATES	SEU ON OPTIMIZATION REGION TES 40 THROUGH 541	AT10N REG H 54)	IUN		STATISTICS BASEU DN UPTIMIZATIUN REGIUN (URUINATES 40 THROUGH 54)
**************************************	04 UF	EQUIV DEPTH	EQUIV MEAN CENTER UF MASS	TIME TO CENTER OF MASS	LAG PEAK TIME OF C.M. TO FLOW PEAK	PEAK	TIME OF
PRECIPITATION EXCESS		0.313		44.00			
CUMPUTED HYDROGKAPH OBSERVED HYDROGKAPH	8331. 8342.	0.223	555. 556.	49.04 48.83	5.04	1179.	46.00
DIFFERENCE PERCENT DIFFERENCE	-11.	000-0-	-1-	0.21	0.21	-166.	-1-00
STANDARD DAJECTIVE FU	STANDARD ERROR OBJECTIVE FUNCTION	74.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERRUR ABSULUTE ERROR	62. 34.78	
STATISTICS BASED UN FULL HYDRUGKAPH (ORDINATES 1 THROUGH BO)		STATISTICS BASED (ORDINATES	BASED UN FULL HYI TES 1 THROUGH	UN FULL HYDRUGKAPH 1 Through 801	STATISTICS BASED UN FULL HYDRUGKAPH (ORDINATES 1 THROUGH 80)		
	SUM OF FLOWS	EQUIV	MEAN	TIME TO CENTER OF MASS	LAG C.M. TO C.M.	PEAK	TIME OF
PRECIPITATION EXCESS		0.413		44.00		:	; 1
LOMPUTED HYDRUGRAPH UBSERVED HYDRUGRAPH	13018.	0.348	163.	53.63	9.63	1179.	46.00
PERCENT DIFFERENCE	-3763.	101.0-	-47.	8.43	8.43	-166.	-1.00
STANDARD	STANDARD ERROR	35.	10 0 0 NA	AVERAGE	ABSOLUTE ERROR	18.	

						- •		UNIT	NYGRUĞR -PERIUU			ES					·- <del>· · · · · · · · · · · · · · · · · · </del>	
			42). lel4. 404.		404. 352.	100. 1240.	5159. 1461. 201.	3715. 424. 232.	3234. 604. 204.		2#1 10 17	7. •-	2452. 613. 133.	13 21 213	1.	1858. 464. 116.		
			131.		Je.	77.	67.	54.	٠٠.		•	٠.	34.	•	3.	29.		
•••	••••	•••••	••••••	••••	•••••	••••••		DRUGRAPH A				 LYM		******	******	******		******
•••	***	•••••	•••••	••••	******	•••••	•••••	••••••	•••••••	•••	••••	•••••	•••••		•••••	•••••	••••••	******
JĀ	PH 144	*****	ran ra	44 [N	F112.2	t MESS	Clime a	085 4		J.	MUN	HRMN	ORD	RAIN	ross	EXCESS	COMP 4	045
	Jul	444		<b>u.</b> Ju	4.40	J.00	y.	٠.	•			0100	41	0.04	0.00	0.00	į.	33.
45		1 300		0.01	0.01	0.00	••	••	:			0540	42	0.13	0.13 3.5 <b>9</b>	0.00	j. 3.	35.
?*		1130		1.54	0.52	J. J.	ý. 8.	10.	•			9300	43	0.50 0.48	0.17	0.31	135.	92
??		1300		4.45 FU.L	U.J5	J.30 J.30	<b>;</b> :	21.	:			4544	45	0.12	0.12	0.00	192.	376
		1400		9.12	0.12	4.00	<b>:</b> :	33.	·			U600	**	0.01	J.01	0.30	915.	1000
		1500		J. Je	3.34	0.30	ä.	180.	•			0100	41	0.00	U. 00	0.00	1174.	1102
24	JUL	1600		y. JI	0.01	0.00		514.	•	31	1.1	-	4.0	0.00	0.00	U.93	1105.	1345
84	JUL	Lrog		4.31	3.41	U. UU	1.	415.	•	31	JUL	8700	49	0.02	0.02	ဖု.လ	1015.	1075
24	JUL	1 600	10	J.OL	0.91	J.03	1.	335.	•			LJUG	34	0.00	0.00	0.00		650.
		[ And	4.6	J.JŁ	J.JL	J. JO	7.	266-	•			1100	51	0.00	0.00	0.00	110.	123
		\$000		<b>0.J</b> 0	0.00	0.00	7.	194.	•			1200	52	0.00	0.00	0.00	els.	+12
		\$100		0.00	3.00	0.00	<u>"</u> .	173.	•			1300	53	0.00	J.30	0.00	564. 307.	514.
		2200	• .	J.JJ	3.00	J.JJ	7.	149.	:			1400 1500	54 55	0.00 U.00	0.00	0.30	¥3.	370
		4 4 Q U		4.00	3.JJ 0.JJ	3.30	<b>6.</b>	124.				1400	56	J. 00	4.00	0.00	384.	320.
		2100		3.00	0.00	J.J0	•	112.				1703	57	0.00	0.00	0.00	334.	284
		4204		0.00	3.00	J.00	÷:	102.	•			1003	54	3.00	0.00	0.30	293.	202
		0130		3.00	2.00	U. 00	•.	94.	•			1900	54	0.00	J. UO	9.00	255.	239.
		0440		U.0U	3.00	J.00	••	85.	•			2000	ěJ	J. 60	0.00	g.3u	222.	210.
30	JUL	0900	21	U. U.	J. U3	J. JJ	••	al.	•	41	JUL	2100	• L	J. 00	U. 0U	4.00	194.	202.
10	JUL.	0000	:2	9.00	3.00	J.00	5.	14.	•			2200	42	J. 90	3.00	0.00	149.	190.
		J73J		<b>J. G</b> Q	0.40	J.J0	3.	70.	•			2300	• 1	3.90	0.00	0.00	147.	140
		9800		0.00	0.00	4.00	5.	44.	•			2000	• •	0.00	0.30	0.00	142.	149.
		0900		J. 00	0.00	4.00	5.	•2•	•			3100	65	3.00	U. UU	0.00 0.00	139. 135.	162.
		1000		0.00	3.00	٠. ٥٠	5. 5.	60. 56.	- :			0300	67	J.00 J.00	0.30	0.00	132.	[40]
		1100		0.00	0.00	J.00	3.	5				4444	14	J. 4J	3. 33	0.00	129.	143
		1300		u.30	0.00	0.00	7.	51.	:			4530	• 4	0.00	0.00	0.50	124.	137
		1 -00		0.00	0.00	J.J0	ś.	49.	·			Jeus	ta	0.00	0.00	0.00	123.	134
		1504		3.33	3.30	3.00	<b>4</b> :	48.	•			0700	71	0.00	3.33	0.00	120.	131.
		1000		J. JU	0.00	0.00	4.	46.	•			J 60U	12	0.00	0.03	3.00	117.	126
		1730		J. JU	3.00	0.33	••	42.	•			1403	7.6	0.00	0.00	0.00	115.	122.
IJ	JUL	LEGU		3.00	3.33	U. UO	٠.	41.	•	1	446	1003	74	J-00	4.03	0.30	112.	150
		1 407	35	0.00	J.U0	u.00	٠.	19.	•			1133	15	0.00	0.30	0.00	109.	117.
		\$000		0.00	0.00	3.50	٠.	34.	•			1200	74	0.00	0.00	0.00	107.	115
		\$100		J.0U	0.00	w. 00	••	30.	•			1300	11	0.00	0.00	0.00	104.	104
		1100		3.33	3.00	0.00	••	35.	•			1433	74	0.00	0.00	0.00	102.	107
		2 443		J.JU	9.00	ود.ه	••	35.	•			1300	14	0.00	0.00	0.00	100.	134
11	ALK.	UUIN	40	U. JJ	0.00	<b>u.u</b> 0	4.	11.	•	ŧ	AUG	LOOU	<b>e</b> u	U- 00	<b>U.JO</b>	0.00	<b>•</b> ?.	102

TOTAL MAINFALL . 2.24. IUTAL LUSS . 1.47. TOTAL EXCESS . U.32

PEAR FLUR 11ML MAXIMUM AVERAGE FLUW
(LF3) (MR) 6-MR 24-ML 72-MR 74-00-MI
1174- 40-30 (LF5) 407- 471- 174- 104(14CHE) 3-155 3-307 0-305 0-347
(4C-FF) 4dJ. 455- 1007- 1072-

COMPLATENT AREA + STUNG SQ 41

STATION SPLYA

	u.	20	w.	10) 0	MIFLUW,		045EK40 -CD8		9. LZ9	W. 140	Ju	). d		). <b>.</b>	• • •
	U.U PEP	4		J.	. 3	J. J	٠. ١	<b>.</b>		.0 4		.s a.	4 11 74	• 0.	EXCESS 0.0
\$78 VV	10												<i></i>		
\$41 100 \$41900	70 50		:		:	:	:	•	:	:	:	:		imm	
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\$45000	120		•	••••	:	• : •	• • • • •		: • • • •	:	:	:	• • • • •		•
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100200	140	•	•		•	•			•	•	•	:	•	•	
300 300 1004-0	190	•	:		:	:	:	1	:	:	:	:	•	•	: :
100500	510 ·						• • • •		·						
100 FOU	5 JU	:	:		:	:	:		:	:	:	:	:	•	: :
300400	24G 25U	•	•		•	•	•		•	•	•	•	•	•	
361000	310 200	:	:		:	:	:		:	:	:	:	:	•	: :
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301300	290	•	:		:	:	:		:	:	:	:	•	:	: :
301 400	310		•		•	•	•		•	•	•	•	•	•	•
301603	320				<u> </u>		• • • • •	• • • •	•	:	:	: • • • •	• • <i>• • •</i>		
101 100	340		•		•	•	:		•	•	•	•	•	• 1	
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- ^	25	2 C	CANASAWALTA CREEK Unit Graph (Clark)	A CREEK	6-22/23-72 AND 1.055 RATE	3-72 RATE (	CUNIFORM: UPTIMIZATION	UPTIMIZ	ATION				
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		STATISTICS BASED	CS BASED UN OPTIMIZATION RÖINATES 7 THRUUGH 45	PTIMIZATION REGION THRUJGH 45)	STATISTICS BASED UN OPTIMIZATION REGION (ORDINATES 7 THRUIGH 4>)		
		EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	C.H. TO	PEAK	TINE OF PEAK
PRECIPITATION EXCESS		3.192		22.05			
COMPUTED HYDROGRAPH UBSERVED HYDROGRAPH	100915.	2.701	2588. 2588.	25.15	3.70	4971. 5000.	19.00
UIFFERENCE PERCENT DIFFERENCE	-1.	000.0-	• 0 -	-0.62	-0.62	-29. -0.58	••00
STANDÁRD E OBJECTIVE FUNC	STANDARD ERRUR CTIVE FUNCTION	448. 471.	AVERAGE	AVERAGE AB Percent ab	ABSOLUTE ERROR ABSOLUTE EKROR	389. 14.99	
		STATISTICS BASED (ORDINATES		UN FULL HYDROGRAPH 1 THROUGH 85)			
	SUM OF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	C.M. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		3.192		22.05		į	
COMPUTED HYDRÜGRAPH OBSERVED HYDRÜGRAPH	130672.	3.447	1537. 1609.	34.00	11.95	4971. 5000.	19.00
DIFFERENCE PERCENT DIFFERENCE	-6116-	-0.164	-12.	-1.44	-10.44	-24.	00.9
STAND	STANDARD E480R	336.		AVERAGE	AB SOLUTE ERROR	250.	

	•			UNI	I MYDRUGHA	PH			
				44 EMD-0	F-PERIOD U	ROINATES			
217.	14/6.	122).	3059.	3324.	2430.	2541.	2211.	2007.	1770.
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125.	11.	¥1.		76.	67.	54.	52.	46.	<b>~0.</b>
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D7		***	UPL	***	fn22	tart22	com 4	085 4	:	DA	AUN	11 R 11 M	OKU	**	Luis	EXCESS	COMP 0	U45 J
71	ALPS	1.200	ı	3.30	0.04	J.00	44.	64.	·	24	JUN	J 10J	44	4.00	J. J.	U.00	1553.	2140.
		1300	ż	3.05	0.05	J.00	•••	70.	•			0400	45	0.00	0.00	0.00	1374.	1960.
		1400	•	4.45	0.05	9.00	45.	12.	•			1933	46	0.00	4.00	0.00	1211.	L740.
		1500		U.JL	J.01	3.33	<b>63.</b>	14.	•			1000	47	J.00	0.00	0.00	1370.	1445.
		1400	Š	4.30	0.00	u. 3a	62.	74.	•			1100	4.0	u.00	0.00	0.00	745.	1115.
		1 700		3.00	3.00	J. JJ	40.	63.	•			1500	49	U.JO	0.00	0.03	635.	1203.
		1400	7	4.00	0.00	J.J0	59.	.7.	•			1300	Su	0.00	0.00	0.43	144.	1166.
		1903	À	4.31	3.01	J. JJ	56.	102.	•			1400	51	0.00	U.00	9.00	111.	1113.
		2000	Ţ	4.60	0.60	9.00	56.	117.	•			1500	52	J. 04	4.01	0.05	7>4.	1063.
		2100	10	0.68	0.20	0.40	201.	370.	•			1600	53	0.15	0.01	0.14	141.	1010.
		2200	ii	0.49	0.01	4.43	1033.	880.	•			1700	59	0.06	0.01	4.35	433.	1030.
		2300	12	0.31	0.01	J.30	2330.	1940.	•			1400	22	0.02	0.01	0.01	1054.	1043.
			13	0.37	0.31	J.J6	3527.	1400.				1900	20	0.03	0.31	0.42	1144.	1125.
		0100	14	9.18	0.31	3.17	4149.	5000.	•			2000	57	0.02	0.01	J.01	1114.	1106.
		n \$ 00	15	3.10	3.31	J. 09	4324.	4/44.	•			\$100	54	0.00	J.00	0.03	1044.	1175.
		1111	i è	J.J.	3.01	J. 07	4337.	4010.	•			2200	59	0.00	0.00	0.30	954.	1162.
		3400	17	0.31	0.01	0.30	4349.	4430.				2300	6.0	U.00	0.03	0.00	652.	1113-
		3500	10	3.24	3.01	0.23	4603.	4550.	•			JJJ00	•1	0.00	0.33	3.30	149.	1050.
		0400	19	0.01	0.01	J.00	4431.	4670.	•			3100	62	0.00	u.00	0.00	740.	1030.
		9700	žo	J. 01	J.01	J.30	.971.	4340.	•			0200	63	0.00	0.00	0.33	762.	970.
		0800	21	1.03	3.00	4.00	4600.	401U.	·			2300	••	J.00	0.00	0.00	744.	920.
		0 400	22	3.30	0.00	0.00	4384.	1500.	·			4400	65	4.00	3.00	0.00	726.	840.
		1000	23	0.00	3.00	u.30	3034.	3083.	•			3500	44	0.00	0.00	0	709.	640.
	JUN		24	J. v0	3.33	v. vo	3141.	2015.	•			UBBU	•7	3.00	0.00	0.03	693.	ato.
	JUM		25	3.33	8.00	J.03	2646.	2325.	·			3730	••	9.00	U.00	0.00	676-	780.
	AUL		25	J. JO	J.00	J. JO	2419.	2040.	•			J800		U.U0	0.00	0.00	601.	750.
	JUN		27	J. 02	J. J1	3.01	2195.	1420.	·			3400	73	0.00	0.00	0.00	en5.	723.
	JUN		2.	0.01	3.01	0.33	1954.	1020.	·			LUQU	71	0.00	0.00	0.00	e30.	705.
		1000	24	3.33	J. 01	J.J2	1754.	1465.				1107	72	3.00	0.00	0.00	015.	674.
		1700	33	3.37	9.01	0.36	1620.	1390.	·			1230	73	0.00	0.00	0.30	401.	
	JUN		31	0.31	0.01	U.00	1552.	1315.				1300	74	0.00	0.00	0.00	567.	636.
		1900	52	3.17	3.31	3.10	1554.	1300.	•			1400	75	0.00	0.30	0.00	5/3.	620.
	AUA.		33	3.17	3.01	J. 16	1/11.	1405.	-			1500	ï	0.00	0.00	3.00	560.	.04.
	JUN		34	0.23	J.31	0.22	2084.	1/40.	•			1900	77	0.00	0.00	0.00	544.	546.
	JUN		15	0.23	3.01	J.22	2030.	2200.				1700	7 6	0.00	J.00	9.00	534.	544.
	JUN		36	v.03	0.01	U.02	3145.	2800.				1 403	79	0.00	0.00	4.00	321.	548.
	JUN		37	4.42	3.31	4.91	3354.	3530.	:			1 400	80	0.00	0.00	0.00	509.	524.
	JUN		34	9.01	0.31	3.01	3234.	3660.	:			1200	81	0.00	0.00	J. W	497.	508.
	JUN		17	4.90	3.00	0.00	2483.	3540.	:			2100	**	0.00	9.00	0.55	+#5.	493.
	JUR		45	J.00	J.00	J.03	2550.	3320.	:			2200	43	8.00	0.00	0.00	474.	486.
	JUR								:				44			0.00	443.	4/2.
	JUN		*1	3.30	0.00 J.JJ	J.JO	2257. 1992.	3020. 2700.	:			2300	4>	3.30	0.03	0.00	452.	465.
	J.W		42	J. UU					•	£3	304	<b>110</b> 0	->	0.00	u.vJ	5.50	476.	70,0
43	-	4440	7,	0.00	0.00	J.00	L759.	2400.	•									
									•									

TUTAL < 1. FALL = 0.05. TUTAL LUSS = 1.20. TUTAL EXCESS = 3.10

PEAR FLUE	FIRE			BYA MEPIRAN	RAGE FLUS	
16+51	[144]		6-HE	24-104	72-ma	84 . OJ-HA
4471.	1+.30	ICFSI	4615.	3171.	1743.	1553.
		I I NCHESI	0.741	2.057	1.435	3.490
		1 AC -+ T 1	2266.	6299.	Luede.	10//8.

CUMULATIVE ANEA . 57.40 SQ MI

EALES	0.	•.	14.7 PME	•.	١.	0.		ø.	•	•	wu.	54	4000.	•	juoo		1000.		Ñ.	Ìuõ	١.	•	
PALES			0.	9.0	•	0.	)	٥.٥	v	٥.	3.0	ı	J.0	•	٠.		9.3		J.J	J		J.	
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		AKISON OF COM	COMPAKISON OF COMPUTED AND OBSERVED HYDROGRAPHS	RVED HYDRO	GRAPHS		
		STATISTICS BA	FICS dased un uptimization region (UADINATES L'IHROUGH 15)	ATION REGION 15)	STATISTICS DASED UN UPTIMIZATION REGIUN (URDINATES L'THROUGH 15)	•	
		EQUIV	MEAN	TIME TO CENTER OF MASS	UM OF EQUIV MEAN CENTER C.M. TO PEAK TIME OF FLUMS DEPTH FLUM OF MASS C.M. FLOW PEAK	PEAK	TIME UF
PRECIPITATION EXCESS		0.807		4.51			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	26373.	0.706	1758.	9.79	5.28	3816.	8.00
DIFFERENCE Percent difference	-5-	000.0-	0-	-0.25	-0.25	-279.	00.0
STANDARU UBJECTIVE FUN	STANDARD ERRUR UBJECTIVE FUNCTIUN	279.	AVERAGE			186.	1
		STATISTICS BASEU (ORDÍNATES		ON FULL HYDKOGRAPH 1 THROUGH 55)			
1	SUM UF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.H. TO C.H.	PEAK FLOW	TINE OF PEAK
PRECIPITATION EXCESS		0.807		4.51			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	36280.	0.971	660.	15.32	10.81	3816.	8.00
DIFFERENCE Percent difference	-7350.	-0-197	-134.	-2.17	-2.11	-279.	0.00
STANDARD OBJECTIVE FU	STANDARD ERRUR OBJECTIVE FUNCTION	225.	AVERAGE	AV ERAGE P ERCENT	ABSULUTE ERROR	184.	

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	STA	= :	STICS BASED ON OPTIMIZATION REGION (UNDINATES 7 THROUGH 35)	ATION REGION ATION	NO		
	SUM OF FLOWS	EQUIV DEPTH	HEAN	TIME TO CENTER UF MASS		P EAK Flow	TIME OF PEAK
PRECIPITATION EXCESS		1.128		11.66			
COMPUTED HYDRUGRAPH OBSERVED HYDRUGRAPH	68046. 68147.	0.631	2346.	24.76	13.10	3539. 3691.	24.00
DIFFERENCE PERCENT DIFFERENCE	-101.	100.0-	-3.	0.27	2.07	-152.	00-1-
STANDARD E OBJECTIVE FUNC	ARD ERROR FUNCTION	249.	AVERAGE	AVERAGE AE Percent ae	249. AVERAGE ABSOLUTE ERROR 180. 242. AVERAĜE PERCENT ABSOLUTE ERROR 9.78	180. 9.78	
*	**************************************	STATISTICS BASED (URDINATES	STATISTICS BASED ON FULL HYDROGRAPH (URDINATES I THROUGH 110)	ON FULL HYDROGRAPH I THROUGH 1101			
		EQUIV OEPTH	MEAN	TIME TO CENTER OF MASS			TIME OF PEAK
PRECIPITATION EXCESS		1-128		11.66		į	
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	139634.	1.212	1184.	40.48	28.83 32.38	3539. 3691.	24.00
DIFFERENCE PERCENT DIFFERENCE	-9130. -6.53	-0.085	-63.	-3.55	-3.55	-152.	-1.00
STANDARD E	ARD ERROR	219.		AVERAGE AB	ABSULUTE ERRUR	183.	!

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	SUM UF FLOAS	EQUIV OEPTH	MEAN	TIME TO CENTER OF MASS	C.M. TC	PEAK	TINE OF PEAK
PRECIPITATION EXCESS		1.028		15.56			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	83096. 83107.	0.771	2374.	22.91	7.36	4078. 4299.	25.00
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		STATISTICS BASED (URDINATES	SASED ON FULL TES 1 THROUG	ON FULL HYDROGRAPH 1 THROUGH 67)	STATISTICS BASED ON FULL HYDROGRAPH (URDINATES 1 THROUGH 67)		
		EQUIV	MEAN	TIME TO CENTER OF MASS	C.M. TU	PEAK	TIME OF PEAK
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		STATISTICS BASED IDRDINATES	SED UN OPTIMIZATIUN REGIUN TES 8 THROUGH 18)	ATION REC 11 18)	STATISTICS BASED UN OPTIMIZATIUN REGIUN (ORDINATES B THROUGH 18)		
		EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	C.M. TU	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.544		13.24			
COMPUTED HYDRUGRAPH OBSERVED HYDRUGRAPH	1610.	0.204	146.	15.26	2.02	293.	16.00
DIFFERENCE Percent difference	0.10	0.000	•	0.03	0.03	-9.	-1.00
STANDAKD ERKOR 16. AVERAGE PERCENT ABSOLUTE ERROR 15. Objective function 18. Average Percent absolute error 21.74	JAKD ERKOR FUNCTION	16. 18.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERRUR	15.	
STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THROUGH 45)	,	STATISTICS BASED (URDINATES	SASED ON FULL FES 1 THROUG	ON FULL HYDRUGRAPH 1 THROUGH 451	I		
	SU4 DF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	C.H. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.544		13.24			
COMPUTED HYDROGHAPH UBSERVED HYDROGRAPH	4318.	0.549	96.	22.52	9.28 9.52	293.	16.00
DIFFERENCE PERCENT DIFFERENCE	-55. -1.25	100.0-	.1.	-0.24	-0.24	-33.	-2.00
STANDARD	STANDARD ERROR	20.	6	AVEKAGE	ABSOLUTE ERROR	14.	

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		STATISTICS BASED (ORDINATES	SEO ON OPTIMIZATION REGIOM IES 45 THROUGH 57)	ATION REG H 57)	STAFISTICS BASED ON OPTIMIZATION REGIOM (ORDINATES 45 THROUGH 57)		
TIME TO LAG SUM UF EQUIV MEAN CENTER C.M. TO PEAK TIME UF FLOMS DEPTH FLOM UF MASS C.M. FLOW PEAK	SUM UF FLOWS	EQUIV DEPTH	MEAN FLOW	TIME TO CENTER UF MASS	C. M. TO	PEAK FLOW	TINE OF PEAK
PHECIPITATION EXCESS		3.266		41.90			
COMPUTED HYDROGRAPH OBSERVEU HYDRGGRAPH	16359.	2.018	1258.	51.58 51.59	9.68 9.69	2237.	51.00
DIFFERENCE Percent difference	-32.	-0.004	-2.	-0.01	-0.01	111.	2.00
STANDAKD OBJECTIVE FUI	JARD ERROR E FUNCTION	123. 128.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR Absolute error	108.	
STATISTICS BASED ON FULL HYDROGRAPH  (ORDINATES 1 THROUGH 70)		STATISTICS BASED (ORDINATES	BASED ON FULL HYDROG TES 1 THROUGH 701	ON FULL HYDROGRAPH 1 THROUGH 701			
TIME TO LAG SUM OF EQUIV MEAN CENTER C.M. TO PEAK TIME OF	SUM OF	EQUIV	MEAN FLUM	TIME TO CENTER OF MASS	LAG C.H., TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		3.266		41.90			
CUMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	27608. 25580.	3.507	394. 365.	44.79 48.46	2.89 6.56	2237.	51.00
DIFFERENCE PERCENT DIFFERENCE	2028.	0.258	29.	-3.67	-3.67	111. 5.23	2.00
STANDARD	STANDARD ERROR	62.	7 CA URVA	AVERAGE Deptent	ABSOLUTE ERROR	32.	

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		EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.M. TU C.M.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.006		9.62			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	5041.	0.640	420. 403.	12.11	2.49	824. 838.	10.00
DIFFERENCE Percent difference	235.	970.0	17.	0.44	0.89 55.58	-14.	0.0
STANDARD UBJECTIVE FU		171. 178.	AVERAGE	AVERAGE	ABSOLUTE ERROR ABSOLUTE ERROR	141.	
		STATISTICS BASED (ORDINATES	JASED ON FULL HY. TES 1 THROUGH	ON FULL HYDRUGRAPH I THROUGH 30)	STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES 1 THROUGH 30)		
		EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.M. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.006		9.62			
CUMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	7719.	0.980	257. 238.	15.23	5.61	324. 838.	10.00
DIFFERENCE Percent difference	578. 8.10	0.073	19.	0.52	0.52 10.29	-14.	0.00
STANDARD	STANDARD ERROR	120.	3040000	AVERAGE	ABSOLUTE ERRUR	*	

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		STATISTICS BASED (URDINATES	SED UN UPTIMIZATES 5 THROUGH	GN UPTIMIZATION REGIUN 5 THROUGH 23)	STATISTICS BASED UN UPTIMIZATION REGIUN (URUINATES 5 THROUGH 23)		
•		EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	LAG C.H. TO C.H.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.826		10.08			
COMPUTFO HYDROGRAPH CRSERVED HYDROGRAPH	4918. 4936.	0.625 0.627	259.	15.28	5.19	404.	11.00
DIFFERENCE PERCENT DIFFERENCE	-18.	-0.002	-1.	0.22	0.22	8.2.00	0.00
STANDARD EPROP UBJECTIVE FUNCTION	STANDARD EPROP UBJECTIVE FUNCTION	31. 31.	AV ERAGE	AVERAGE PERCENT	ABSOLUTE ERROR Absolute error	25. 15.21	
STATISTICS BASED ON FULL HYDROGRAPH (ARDINATES 1 THROUGH 40)		STATISTICS BASEU (ARDINATES	BASED ON FULL HY( TES 1 THROUGH	ON FULL HYDROGRAPH 1 THROUGH 40)	STATISTICS BASED ON FULL HYOROGRAPH (IRDINATES 1 THKOUGH 40)		
	SUM OF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.826		10.08			
COMPUTEU HYDRUGRAPH OBSERVED HYDROGRAPH	6499.	0.825 0.818	162.	14.70 18.56	8.62 8.48	404.	11.00
DIFFERENCE Percent difference	61. 0.95	900.0	2.	0.14	0.14	8.2.00	0.00
STAND	STANDARD EPPOP	25.	24	AVERAGE	ABSOLUTE ERROR	18.	

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		STATISTICS BASED (ORDINATES	STATISTICS BASED ON UPTIMILATION REGION (ORDINATES 3 THROUGH 12)	ON OPTIMIZATION REGION 3 THROUGH 12)	IUN		
	• •	EQUIV DEPTH	MEAN FLOM	TIME TO CENTER OF MASS	TIME TO LAG  TOUTH FLOW OF MASS C.M. FLO.: PEAK	PEAK FLO:	TIME OF
PRECIPITATION EXCESS		0.715		4.52			
COMPUTED HYDRUGRAPH OBSERVED HYDRUGRAPH	4836. 4842.	0.614	4 8 4 4	7.76	3.24	878. 467.	6.30
DIFFERENCE PERCENT DIFFERENCE	-6. -0.13	100.0-	-1-	0.17	0.17 5.65	11.	00.00
STANDAPO DAJECTIVE FL	STANDAPO ERROR UBJECTIVE FUNCTION	74.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE FRROR	58. 14.86	
		STATISTICS BASED (ORDINATES	BASED ON FULL HYL TES 1 THROUGH	ON FULL HYDROGRAPH I THROUGH 30)	STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES 1 THROUGH 30)		
		EQUIV	MEAN	TIME TO CENTER OF MASS	C.M. TU	PEAK FLOW	TIME OF
PRECIPITATION EXCESS		0.715		4.52			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	6975.	0.886 0.929	233.	11.56	7.04	878. 867.	9.00
DIFFERENCE PENCENT DIFFERFNCE	-339.	-0.043	-11-	-0-06	-0.06	11.	00.0
STANDARD	STANDARD ERROR	46.	940974	AVERAGE	ABSOLUTE ERROR	31.	

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		STATISTICS BASED (ORDINATES	SED ON OPTIMIZATIUN KEGIUN Tes 1 Through 21)	ATIUN REG H 21)	STATISTICS BASED ON OPTIMIZATIUN REGIUN (ORDINATES I THRUUGH 21)	·	
		EQUIV BEPTH	MEAN Flow	TIME TO CENTER OF MASS	C.M. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.337		10.45			
COMPUTED HYDROGKAPH OBSERVED HYDROGKAPH	6704.	158.0	319.	13.36	2.91	510. 475.	13.00
DIFFERENCE PERCENT DIFFERENCE	61. 0.91	800.0	ě	0.25	0.25	35.	-2.00
STANDARD I		40. 40.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERRUR ABSOLUTE ERROR	35.	
		STATISTICS BASE	STATISTICS BASED ON FULL HYDROGRAPH (URDINATES I THROUGH 40)	HYUROGRAP H 401	STATISTICS BASED UN FULL HYDROGRAPH (URDINATES 1 THROUGH 40)		
		EQUIV DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	LAG C.M. TU	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.337		10.45			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	10336.	1.313	258.	18.95	8.50 8.89	\$10. 475.	13.00
DIFFERENCE PERCENT DIFFERENCE	-236.	-0.030	į	-0-39	-0.39	35.	-2.00
STANDARD	STANDARD ERROR	36.	(	AVERAGE	ABSOLUTE ERROR	32.	

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PRECIPITATION EXCESS		164.0		18.00			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	2853. 2854.	0.433	136. 136.	22.09	4.09	347.	20.00
DIFFERENCE PERCENT DIFFERENCE	-1.	-0.000	•0-	0.97	0.97 31.13	27. 8.38	1.00
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PRECIPITATION EXCESS		164.0		18.00			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	3631. 3886.	0.552	73.	25.52 25.56	7.52	347.	20.00
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		EQUIV DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	•	PEAK FLOW	TIME OF
PRFCIPITATION EXCESS		0.465		13.09			
CUMPUTED HYDRUGRAPH OBSERVED HYDRUGRAPH	2366. 2366.	0.359 0.359	131.	17.82	4.73	247. 233.	15.00
UIFFEGENCE PERCENT DIFFERENCE	0.00	000.0	ċ	0.21	0.21	14.	-2.00
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		STATISTICS BASED (ORDINATES	BASED ON FULL HYD TES 1 THROUGH	ON FULL HYDROGRAPH I THR(UGH 45)	STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES 1 THROUGH 45)		
	•	EQUIV DEPTH	HEAN Flow	TIME TO CENTER OF MASS	C.M. TU	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.465		13.09			
CUMPUTEU HYDRUG2APH OBSFRVED HYDRUGRAPH	3355. 3721.	0.510 0.565	75. 83.	22.21	9.12	247. 233.	15.00
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	Ĭ	STATISTICS BASED (ORDINATES	SED ON OPTIMIZATION TES 6 THROUGH 56)	ZATION REC GH 56)	STATISTICS BASED ON OPTIMIZATION REGION (ORDINATES 6 THROUGH 54)		
	SUM OF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	C.N. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		3.358		44.03			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	19221.	2.920	377. 378.	45.85	1.82	1886.	50.00
DIFFERENCE PF9CENT DIFFERENCE	-76.	-0.012	-1:	-0.39	-0.39	146.	2.00
STANDARD UBJECTIVE FU		98. 145.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE ERROR	57.	
		STATISTICS BASED LORDINATES	BASED UN FULL HYD TES 1 THROUGH	UN FULL HYDRUGRAPH 1 THROUGH 901	STATISTICS BASED UN FULL HYDRUGRAPH IORDINATES 1 THROUGH 901		
	SUM DE FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	C.N. TO	PEAK Flow	TIME OF Peak
PRECIPITATION EXCESS		3,358		44.03			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	23405 <b>.</b> 25277.	3.556 3.840	260. 281.	49.31 51.04	5.28 7.01	1886. 1740.	50.00
DIFFERENCE Percent difference	-1872. -7.41	-0.284	-21.	-1.73	-1.73	146.	2.00
STANDARD 08.JECTIVE FU	STANDARD ERROR DBJECTIVE FUNCTION	82. 131.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSULUTE FRROR	52.	

	UNIT	HYDAUGRAP	16			
	23 END-UF	-PERSUO ON	UINATES			
/82.	612.	440.	174.	294.	211.	101.

						44	DRINGPAPH AT	STATION		CLA	KUM:						
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	J393	L	U.00	0.00	0.00	₹.	≥.				0000	46	0.47	0.10	0.17	365.	30
SEP	8400	2	0.12	4.12	4.40	2.	3.				0100	47	0.35	0.10	0.25	731.	44
	0500	3	9.12	9.12	<b>u.</b> 03	z.	٠.				0200	4.5	0.59	0.10	0.49	1103.	150
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	0000	•	0.24	0.22	0.31	10-	14.				<b>0530</b>	51	0.47	4-10	0.37	1414.	16
	0100	7	3.12	J. 10	0.02	30.	45.				9600	52	. 0-12	0.10	4.02	1757.	6.70
	1000		0.00	0.00	U.40	37.	56.				0703	5.5	0.24	0.10	0.14	1464.	145
	1100	•	0.70	U. OU	u.00	36.	53.				0400	54	0.12	0.10	0.02	1275.	12
	1200	10	u.oj	9.00	0.00	35.	45.				3400	55	0.12	0.10	0.02	1025.	10
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	L400	12	0.0C	0.00	u.on	33.	33.				1170	51	0.00	0.00	0.00	<b>663.</b>	7
ser	1500	13	U.DO	0.00	0.00	32.	24.				1200	58	0.00	0.90	0.00	519.	5
	I 400	14	0.00	0.00	0.00	31.	26.				1 300	44	0.03	0.00	0.00	407.	•
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SE P	1 800	l b	0.00	3.00	<b>u.</b> 00	37.	22.	• 1	50	SEP	1500	٥L	0.00	0-03	0.00	250.	3
SEP	1 900	17	0.00	0.00	0.00	24.	le.	•	65	5 F.P	1400	62	0.00	0.03	0.00	194.	2
SEP	2000	1.0	0.00	0.00	0.39	28.	17.	•	65	SEP	1700	6.5	0.00	0.00	0.03	152.	Z
SEP	\$100	14	9.00	0.00	U.00	21.	17.	• 2	20	SEP	1400	44	0.33	0.00	0.00	117.	2
SFP	2200	20	u.0u	0.00	0.00	26.	17.	• •	16	SEP	L 400	65	0.00	0.03	9.00	43.	Z
SEP	2300	21	a.tz	0.10	0.02	25.	17.	•	16	SEP	2003	66	0.00	0.00	0.00	83.	L
SFP	9000	22	0.00	0.00	0.03	31.	16.	• 2	20	SEP	2100	67	0.00	0.00	0.00	€L.	L
SEP	8100	21	0.17	0.10	0.02	37.	27.	• ;	20	SEP	2200		0.00	J-00	0.00	78.	1
SEP	0200	24	J.00	0.00	<b>U.</b> 00	43.	35.	• 2	26	SEP	2300	69	0.00	0.00	0.00	76.	
SEP	<b>a 300</b>	29	J.00	0.00	0.00	41.	42.	• 2		SEP	0000	70	3.40	0.40	0.00	73.	ı
SEP	0400	26	0.12	0.10	0.02	<b>40.</b>	44.	• 2	27	SEP	0100	71	0.00	0.00	0.00	71.	ı
SEP	4563	21	0.12	0.10	9.02	57.	45.	•	17	SEP	0200	72	0.00	0.00	0.00	67.	ı
SEP	9460	28	5.24	U.10	0.14	155.	54.		"	SE#	0300	73	0.00	0.00	0.00	<b>67.</b>	ı
SEP	1703	29	0.24	2.10	0.14	115.	223.	•		SEP	400	74	0.00	J.00	U.00	64.	
SEP	U470	10	0.12	3.10	0.02	350.	401.	• 2	27	SEP	0500	75	8.00	0.03	0.00	42.	
SEP	0907	31	0.35	0.10	0.25	474.	446.		7	SEP	0630	76	0.00	0.00	0.00	60.	
56 +	[ 000	12	0.30	4.00	0.00	553.	393.		7	SEP	U 700	77	2.00	0.00	0.00	59.	
SEP	1100	3 5	3.12	0.10	0.02	446.	142.			SEP	3800	78	0.00	J-00	0.00	57.	
SEP	1200	14	0.12	0.10	0.02	317.	322.		11	SEP	0400	79	0.00	0.00	U.00	55.	
SEP	1 300	35	U.03	0.00	0.00	504.	254.	• 2	7	SEP	1000	80	0.00	0.00	0.00	53.	
SEF	1400	36	0.00	0.90	U. 70	242.	144.	• 2	7	SEP	1100	#L	0.00	0.00	0.00	51.	
SF+	1500	17	0.12	0.10	0.02	203.	140.	• 1		SEP	1200	42	0.00	0.00	0.00	50.	
SFP	1400	515	0.00	0.00	0.00	172.	141.	• 1	7	SEP	1300	43	0.00	0.00	0.00	48.	
	1700	19	0.12	0.10	0.02	lss.	200.				1400	84	0.00	0.00	0.00	47.	
SEP	1802	40	0.24	0.10	0.14	226.	310.				1500	85	0.00	0.00	0.40	45.	
	1900	41	9.12	0.10	0.02	247.	371.				1600	do	0.00	0.00	0.00	44.	
SEP	2300	47	0.12	0.10	0.02	252.	134.				1700	47	0.00	0.00	0.00	42.	
	2100	43	0.00	0.00	0.00	211.	274.				1400	48	0.00	0.00	J.00	41.	
	2200	44	3.33	3.00	0.00	105.	223.				1 400	49	0.03	0.00	0.00	40.	
	2300	45	U. 0U	0.00	0.00	129.	204.				2000	90	0.00	0.00	0.03	39.	

TOTAL MAINFALL . A.46. TOTAL LUSS . 3.60, TUTAL ERCESS . 3.30

PEAK FLIM	F1 4E			MAXIMUM AVE	FAGE FLOW	
(CFS)	[44]		6-4F	24-14	12-HP	64.90-MH
lute.	50.00	(CFS)	1550.	719.	314.	263.
		(INCHFS)	1.419	7.621	3.444	3.553
		(AC-FT)	112.	1424.	1901.	1755.

CHMULATIVE 40EA + 10.20 SQ ME

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I-1 LIMITS OF UPTIMIZATION

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		EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	C.M. TU	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.450		29.14			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	2638. 2620.	0.401 0.348	176.	32.95 32.82	3.22	467.	31.00
DIFFERENCE PERCENT DIFFERENCE	18.	0.003	:	0.13	0.13	-9- -1-35	00.0
STANDARC UBJECTIVE FL			AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE ERROR	13.	,
		STATISTICS BASED (ORDINATES	BASED ON FULL HY( TES _1 THROUGH	ON FULL HYDROGRAPH I THROUGH 50)	STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES _1 THROUGH 50)		
		EQUIV	MEAN	TIME TO CENTER OF MASS	LAG C.M. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.420		29.74			
COMPUTED HYDRUGPAPH OBSERVED HYDRUGRAPH	3399.	0.485 0.516	64. 68.	34.34	4.60	467.	31.00
DIFFEPENCE Percent difference	-201.	-0.031	i	-0.22	-0.22	-6.	0.00
STANDARD ERROR	ARU ERROR	13.	4	AVERAGE	ABSOLUTE ERROR	8	

WATERSHED MODELLING IN THE CHEMUNG AND UPPER SUSQUEHANNA RIVER BASINS USI. (U) ARMY MILITARY PERSONNEL CENTER ALEXANDRIA VA J C STYRON 20 MAY 83 F/G 8/8 AD-A129 964 3/4 UNCEASSIFIED NL



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		STATISTICS BASED (URDINATES	ON OPTIMIZATION 7 THROUGH 20)	ATION REC	STATISTICS BASED ON OPTIMIZATION REGION (URDINATES 7 THROUGH 20)		
	SUM UF FLOWS	EQUIV DEPTH	MEAN FLOM	TIME TO CENTER OF MASS	C.H.G.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.289		11.49			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	1711.	0.260	122.	14.52	3.03	222.	14.00
UIFFERENCE PERCENT DIFFERENCE	-2-	000*0-	-0-	0.10	0.10 3.58	12.	1.00
STANDAR OBJECTIVE FI		14.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE ERROR	20.32	
		STATISTICS BASED (ORDINATES	D UN FULL HY I THROUGH	UN FULL HYDROGRAPH I THROUGH 30)	STATISTICS BASED UN FULL HYDROGRAPH  (ORDINATES I THROUGH 30)		
	•	SUM OF EQUIV FLOWS DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	TIME TO LAG MEAN CENTER C.M. TO PEAK TIME OF FLOW OF MASS C.M. FLOW PEAK	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.239		11.49			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	2135. 2215.	0.324	71.	16.46 16.56	4.97 5.07	222.	14.00
DIFFERENCE Percent difference	-80.	-0.012		-0.10	-0.10	12.	1.00
STANDAR	STANDARD ERROR	13.	4 FR 4 GF	AVERAGE	ABSOLUTE ERROR	8.	

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		STATISTICS BASED (ORDINATES	BASED ON OPTIMIZA Nates 5 through	OPTIMIZATION REGION 5 THROUGH 15)	STATISTICS BASED ON OPTIMIZATION REGION (ORDINATES 5 THROUGH 15)		
		EQUIV DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	C.H. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.911		11.68			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	4738. 4706.	0.720	431. 428.	11.75	0.06	1026.	12.00
DIFFERENCE PERCENT DIFFERENCE	32. 0.68	0.005		0.10	0.10	-35.	1.00
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		STATISTICS BASED (ORDINATES	BASED ON FULL HY TES 1 THROUGH	ON FULL HYDRUGRAPH 1 THROUGH 40)	STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES _ 1 THROUGH 40)		
		EQUIV DEPTH	MEAN Flow	TIME TO CENTER UF MASS	C.H. TO	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.911		11.68			
CGAPUTED HYDROGRAPH OBSERVED HYDROGRAPH	7241. 8509.	1.100	181. 213.	16.24 17.76	4.55 6.07	1026. 1061.	12.00
DIFFERENCE Percent difference	-1268.	-0.193	-32.	-1.52	-1.52	-35.	1.00
STANDARD	STANDARD ERROR	138.	40.44	AVERAGE	ABSOLUTE ERROR	86.	

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	1	AKISUN UF CUM	COMPARISON OF COMPOTED AND DESCRIPED HYDROGRAPHS	VED MYDKI	COMPARISON OF COMPOLED AND UBSERVED HYDROGRAPHS		
		STATISTICS BASED (GROINATES	STATISTICS BASED ON OPTIMIZATION REGION (GRDINATES & THROUGH 20)	4 20)	<b>₽</b> つ		
	SUM OF	EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	<b>)</b>	PEAK FLOW	TIME UF PEAK
PRECIPITATION EXCESS		0.457		14.22			
CUMPUTED HYDRUGRAPH DBSERVED HYDROGRAPH	1429.	0.217	110.	16.94	2.72	256.	15.00
DIFFERENCE PERCENT DIFFERENCE	-2. -0.16	000-0-	0-	1.38	1.38	20.8	00.0
STANDARI OBJECTIVE FI	STANDARI) ERRÜR OBJECTIVE FUNCTION		AVERAGE	AVERAGE A Percent A	ABSULUTE ERRUR ABSULUTE ERROR	48. 58.40	
STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES 1 THROUGH 37)		STATISTICS BASED (GRDINATES	BASED ON FULL HYI TES 1 THROUGH	ON FULL HYDRUGRAPH 1 THROUGH 371			
	SUM OF FLOWS	EQUIV DEPTH	HEAN FLOW	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		254 0		14.22			
CUMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	2785.	0.423 0.316	75. 56.	21.55	7.33	256.	15.00
DIFFERENCE Percent difference	703. 33.75	0.107	.61	2.63	2.63 55.86	20.	00.00
STANDARD ERROR	ARD ERROR	. 7.4	, d	AVERAGE A	ABSOLUTE ERROR	36.	

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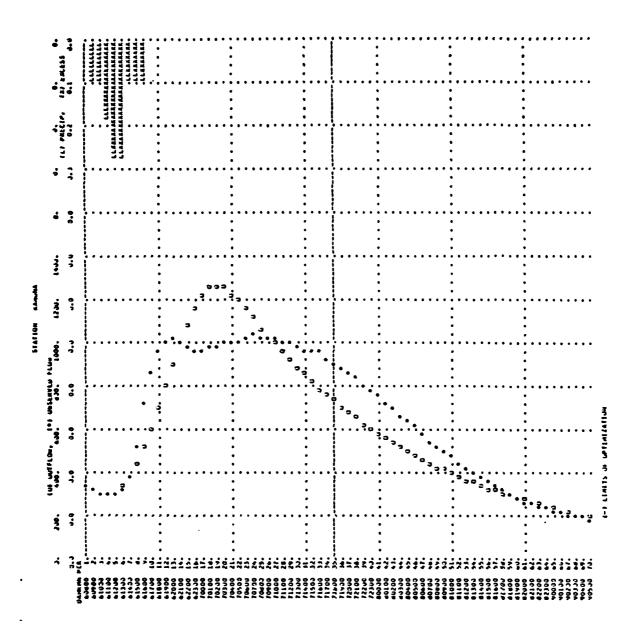
11		ARISON OF COM	COMPARISON OF COMPUTED AND OBSERVED HYDRUGRAPHS	RVED HYD	KUCRAPHS		
		STATISTICS BASEU (ORDINATES	SEJ ON UPTIMIZATION KEGIUN TES 1 THROUGH 35)	ATIUN KE!	STATISTICS BASEU UN UPTIMIZATIUN MEGIUN (URDINATES 1 THRUUGH 35)		
	L .	EQUIV	MEAN FLOM	TIME TO CENTER UF MASS	LAG C.M. TU S C.M.	PLAK FLOM	TIME UF PEAK
PRECIPITATION EXCESS		0. d64		5.60			
COMPUTED HYDROGRAPH UBSERVEU HYDROGRAPH	28968. 28953.	0.672	828. 627.	20.35	14.46	1208. 1037.	18.00
UIFFERENCE Percent o <i>iff</i> erence	15.	000 •0	•	-0-14	-0-14	231.	-5.00
STANDAKU DBJECTIVE FU		154.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERRUR	121.	
		STATISTICS BASEU (OKDINATES	SASEU GN FULL HYJKJG TES 1 THRUUGH 701	ON FULL HYDRUGKAPH 1 THRUUGH 70)	STATISTICS BASEU GN FULL HYVKUGAPH (GRUINATES 1 THRUUGH 70)		
		EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	LAG C.M. TC	PEAK FLUM	TIME OF PEAK
PRECIPITATION EXCESS		0.864		2.40			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	42844. 45275.	0.994	612. 647.	29.71	23.84	1208.	18.00
DIFFERENCE PERCENT DIFFERENCE	-2427.	-0.056	-35.	-1-0-	-0.79	231- 22-27	-5-00
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٠	-	3833	1	3.60	0.00	3.43	330.	330.	•		MAY	1477	50	3.33	دد.ه		101.	404.
•	-	3477	2	3.04	0.09	3.00	320.	314.	•	7	HAY	2000	47	0.00	0.44	دددذ	6/1.	0-
	MAT	1000	3	7-04	C.09	ەد . د	313.	105.	•	1	MAT	2130	3.	3.30	3.33	1.11	*><-	¥34.
٠	MAT	1103	•	0.14	3.04	0.15	303.	244.	•	,	HAY	2400	3.5	0.00	J. 43	3.30	•21.	430.
٠	MAY	1233	3	3,21	0.02	3.25	334.	294.	•	7	MAT	2300	• • • •	0.30	0.00	3.33	-44-	762.
•	MAY	1 160	•	3.27	0.02	0.25	331.	315.	•		MAY	2233	44	0.00		u. Ju	542.	152.
•	MAY	1433	,	J. 34	0- 92	0.07	374.	371.	•		MAY	3197	-2	0.03	3.03	4.10	>+0-	726.
•	MAY	1543	•	C. 09	6.42	0.07	435.	511.	•		MAY	0200	* 5	J. 33	3.00	3.00	>39.	. 90.
•	MAY	1600	4	9.44	0.42	3.37	511-	714.	•		MAY	0101	44	3.33	3.33	1.13	Siv.	445.
		1723	13	3.63	9.00	0. 33	e30.	<b>**</b> 0.	•		MAY	3-10	47	3.00	3-33	4.30	>00-	640.
•	444	しまうり	11	J. 0J	0.00	3.30	944.	954.	•		MAY	0530	**	0.00	0.03	3.30	441-	.10.
		1493	14	3. 83	0.00	9. )	<b>#J3</b> -	1304.	•		44 1	9603	~1	3.30	J-04	0.00	443.	5.0.
•	447	2 36 3	1.2	3.60	0.33	3.63	<b>401.</b>	1013.	•	•	MAT	9733	7.6	3.33	0.00	6	446-	547.
		2100	14	0.00	0.00	ں ۔ ہ	1904.	945.	•		TAY	3430	77	J-00	3.00	3.30	424.	>25.
•	444	5547	15	C.00	0.00	3.30	1043-	971.	•	•	MAF	ULU	>•	J.J0	3.03	4.30	413.	494.
		<b>4363</b>	14	3.00	6.63	J. 90	ile2.	Y64.	•		MAY	1333	>1	3.33	1-30	9.33	J46.	471.
		0003	17	J. 60	0.00	0.00	1214.	457.	•		MAT	1177	>4	9.33	J. 30	Jado	343-	447.
		1107	iu	J. 60	J. 00	0.00	(299.	97a.			MAI	1200	25	u. JJ	2-22	<b>i_0</b> J	367.	427.
		2563	15	3.03	3.00	0.00	1204.	745.	•	•	MAT	1300	>4	3.30	3.00	·-33	#35a	395.
		3130	23	0.00	9.00	J. 30	1257.	992.	•		MAT	1400	>>	J.00	تنوءو	3	142.	374-
		0403	21	3.63	J. 00	<b>U.</b> 00	1233.	1306.	•	•	441	1500	>6	3.00	J.JJ	4.40	329.	<b>357.</b>
		0273	22	3.03	0.00	3.30	1144.	1634.	•	4	447	1007	>1	J_0J	J.33	0.0)	347.	334.
		Jeuj	23	a.co	J. 00	J. 30	1153.	102/.	•	•	MAT	1/33	54	3.44	J.30	0.33	٠٥٥.	312.
		3733	24	J.CJ	3.30	J. 33	IIII.	LJ37.	•	•	MAY	1800	27	3.33	J_33	4	£89.	294.
		3493	23	3.00	0.60	ა.თ	1069.	1320.	•		May	1477	6.0	J.30	نال د یا	0-00	245.	202.
		0400	3.6	0.69	J.00	9.30	1050.	1050.	•			えいいり	• i	3.00	7-00	0.00	£12.	Ze>-
		1033	21	0. ÇJ	J. LO	3.03	441-	1927.	•	•	447	SIO	• 6	J.30	J	u	402.	255.
		1103	20	J. 0J	0. 00	9.00	<b>454.</b>	495.	•		MAY	2233	6.5	0	<b>u.</b> 00	u.0J	253-	292.
		1200	25	0.04	0.00	0.00	910.	499.	•			2100	• •	4.00	J. JJ	3.30	243.	234.
		1300	53	0.00	3.00	3-03	844.	985.	•		Ma T	$\mathbf{y}_{00}$	•5	J.UL	J.UJ	1.33	234.	224.
		1403	31	0.00	0.00	0.03	451.	704.	•			3103		0.33	3.00	0.33	223-	214.
		1203	32	0.00	0.00	<b>9.00</b>	<b>#19.</b>	Y04.				3200	• 7	3.35	3.00	4-33	£11.	204.
		1 400		3.03	0.00	0.33	744.	457.	•			2343		9.33	3.00	3.43	209.	LTV.
		1703	34	2.39	9-00	3.00	154.	423.	•		MAY	3433	<b>.</b> .	0.33	لاناءل	3.43	501-	175-
•	MAT	1601	"	c. Ju	9.00	0.00	7)1.	344.	:	٧	MAY	9>33	10	9-30	3.00	4.33	194.	iss.

TGTAL MAINFALL . 1.18, TOTAL LUSS . G.32, TOTAL ENCESS . J. e.

 PEAR FLUG
 TIME
 MARIMUM AVERAGE PLUG

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					MEC-1 INPUT	NP UT						PAGE	-
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	•	AKISUN UF CUM	COMPAKISON OF COMPUTED AND OBSERVED HYDROGRAPHS	KVED HVU	COMPAKISON OF COMPUTED AND OBSERVED HYDROGRAPHS		
		STATISTICS BASED (ORDINATES	SEJ ON OPTIMIZATION KEGION TES 5 THRUGH 501	ATION KEC	STATISTICS BASED ON UPTIMIZATION KEGION (OKOINATES 5 THKUUGH 50)		
	•	E-UIV DEPTH	MEAN FLOW	TIME TO CENTER UF MASS	C.M. TO	PEAK FLUM	TIME OF PEAK
PRECIPITATION EXCESS		3.125		21.26			
COMPUTED HYDROGRAPH DOSERVED HYDROGRAPH	85934. 85335.	1.993 1.980	1868. 1855.	33.55 34.55	12.50	3352.	30.00 28.00
DIFFERENCE PERCENT DIFFERENCE	599. 0.70	910°0	13.	***0-	94.7-	-1034.	2.00
STANC OBJECTIVE	STANDARD ERRUR OBJECTIVE FUNCTION	382. 430.	AVEHAGE	AVERAGE PERCENT	ABSULUTE EKROR ABSOLUTE EKROK	302. 21.23	
		STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES 1 THKOUGH 110)	BASED UN FULL HYDRUG TES 1 THKOUGH 110)	UN FULL HYDRUGRAPH 1 FHKOUGH 110)	STATISTICS BASED UN FULL HYDRUGRAPH (ORDINATES 1 THKOUGH 110)		
	SUM OF FLOWS	EQUIV	MEAN	TIME TO CENTER UF MASS	LAG C.M. TO	PEAK	TINE OF PEAK
PRECIPITATION EXCESS		3.125		21.26			
CUMPUTED HYDROGRAPH UBSERVED HYDRUGRAPH	130519.	3.028 3.602	1187.	46.01	24.76	3352.	30.00
DIFFERENCE PERCENT DIFFERENCE	-24775. -15.95	-0-575	-225.	-4.71	-4.71 -15.98	-1044.	2.00
STANDARD	STANDARD ERROR	433.		AVERAGE	ABSULUTE ERROR	357.	

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25 SEP	0000	1	3.33	0.00	3.00	31.	31.	• 21	SEF	0703	20	J.33	0.00	ں۔ ن	lolu.	2222.
25 56	3133	2	3.06	3.36	0.30	<b>3</b> J.	34.		SE	ردون م	51	3.30	J.JJ	0.00	1455.	2100.
\$> >E+		3	3.35	ა. თ	0.10	24.	۶2.			Crer 6	> 5	3.34	1.01	3.00	1405-	5151-
52 2F1		•	ي. نو	0.06	3.00	24.	34.			1202	24	J-0J	0.01	رد.،	1350.	2040.
25 251		5	0.00	0.06	3.33	47.	Je-			1133		1.11	0.03	٠.٠٥	1301-	2330.
25 566		,	J.06	0.Je U.Zl	0.30	27. 26.	30. 41.	- 61		P 1200	31	2.03	0.00	3-30	1253.	1940.
25 54		í	0.23	3.23	3.33	25.	¥7.			1497	#2 63	3.33	3.33	2-22	1103.	1034.
25 56		ç	2.19	3.14	3.35	15.	60.			1230	65	3.00	J.55	0-00	1120.	1766.
25 SE		LÓ	3.19	0.32	0.17	•••	97.			2 1630	62	0.30	J.J3	J.JJ	1379.	1711.
25 564		ii	3.30	J. 02	3.34	155.	111.			1133	••	3.33	3.00	0.33	1043.	1040.
25 St	1100	15	C.31	0.02	3.24	301.	337.	. 21	SE	7 1.03	6/	J. 0.	0.00	J.30	1002.	1010.
25 SFF		13	3.23	0.02	3.15	515.	441.	• 41	36/	P LYJU		J.J.	0	3.03	40>-	Lo 7ú-
25 SEP		14	3.35	0.02	2.31	150.	541.			P 2333	65	3.65	رن. د	4.33	424.	1>40-
25 SE		15	3. 11	0.01	3.33	445.	e75.			5 5177	7.0	3.33	3-03	0.30	*65-	1473.
25 SE		lo	3.06	0.02	3.34	13>0.	770.			2203	7.1	0.00	0.00	4	***	1426.
25 SEP		17	3.13	0. UZ J. OZ	J.10	1122.	642. 676.			6 5310	72	J.J.	J.UJ	J.JJ	#31.	1326.
25 561		iv	3.08	0.02	0.05	1230-	902.			b 0100	13	3.03	3.00	0.00	7/1.	1312.
25 521		23	3.04	3.02	3.37	159>-	753.			- 3233	75	3.00	0.00	0.30	1-3.	1200.
25 St		21	3.25	0.02	0.23	1394.	1034.			טנני ל	ió	3.30	3.03	J.JU	716.	1432.
25 564		22	3.25	0.02	2.23	1530.	1125.			P 1411	77	3.30	3.33	v.30	047.	1200.
25 SEF		23	3.19	3.02	3.17	1732.	1334.			2500	7.	J. JU	3.03	3.03	604.	1153.
25 SEF		24	3.13	0.02	0.13	1959.	tezs.			7 3033	14	0-00	0.00	0.00	b40.	1111-
20 St/		25	3. 19	0.02	0. 17	2166.	2037.	• 20	Sêl	2703	• 3	3.33	JJ	3.03	ele-	10/6-
24 56		26	3.28	3.02	0.25	2430.	2222.			. 0.07	e i	0.00	3.33	3.33	>94.	1055.
50 PF		27	3.21	0.02	0.19	2710.	2407.			6 04C7	84	0.00	3.00	0.30	>12.	494.
50 2F1		24	3.13	0-05	0.13	2994.	3,50.			1200	63	J.J0	3.03	J.J3	>51-	457.
50 PC		29	3.14	0.02	5.15	3202.	4380.			6 117º	- 4	J. 40	3.30	4-33	>31-	726.
20 SE		30	3.38	0.32	3.35	3314.	3900.			- 1230	#5	0.00	3.10	9-00	>ii.	evo.
24 SEF		32 31	3.35 3.11	0.u2 3.u2	3.31 3.39	3372. 3346.	1248. 1024.			1303	••	J.Ju	0.33	4.33	473.	esj.
20 51		35	2.09	0.02	3.07	3322.	2047.			7 1930	J/	0.00	3.63	0.33	37.	ele. 766.
20 56		34	3.33	0.02	3.00	1247.	2766.			1900	37	2.00	U.JJ	U.JJ	****	140.
24 St		15	J. Ji	0.31	3.33	32.5.	2646.			1/33	7.	J. J.	33	3.35	*44.	705.
26 SF		30	3.03	0. 33	3.33	3172.	2033.			1000	91	J.CL	3.33	3.30	434.	675.
Zo sti	1500	11	3.30	3.03	J. u3	3370.	2647.			P 1435	ψž	J.J.	3.30	J.J.	344.	647.
26 Sc#		1.3	3.00	3.03	1.33	2454.	2010.		36	2335		3	3.00	3.33	117.	D25.
20 52/		34	3.00	J. 60	J.JJ	2450.	2064.			2133	**	J.0J	J.03	0.33	345.	>41.
20 St		•3	3.00	0.03	0.00	2746.	2601.			, 5537	*>	J. 40	J.00	3.33	<b>356.</b>	>45.
20 Sef 20 Sef		41	33	0	3.40	2645.	2004.			5,20	70	3.33	1.00	3.33	234.	>30.
20 567		42 43	3.60	0.00	J.30	2546.	2604.			بربد م	7/	9.00	J.03	٠.٠٠	127.	507.
20 569		44	0.00	3.00	3.33	2455. 2365.	2098. 2098.			, 2527	78	0.00	4.10	4.33	115.	****
20 56			0.00	3.00	3.33	2270.	2081.			2233	144	0.66 0.00	3.33	3.33	305. 243.	455.
20 361		46	0.00	3.00	0.00	2179.	2004.			0413	121	0.00	J.JJ	J.JJ	243.	431. 463.
20 SEF	2200	47	3.00	0.00	0.00	2114.	2004.			2 2200	122	3.03	3.35	2.13	216.	361.
20 SEF		**	3.63	J.00	0.00	2036.	2014.			2003	1	J.J.	٠.٠٥	6.35	401.	307.
27 56		**	3.00	0.00	0.33	1402.	2543.			0110	lus	J.JÜ	0.00	3.33	259.	347.
27 SE		53	3.00	3.33	J. JJ	LW9J.	2550.	. 29		1030	135	U-30	3.33	0.33	251.	333.
27 SE		>1	3.33	3.03	3.30	1851.	2532.			, 7472	100	0.00	J	6.18	245.	310.
27 569		52	3.00	3.03	0.30	1724.	2425.			ردرا م	Lyl	J.00	0.33	3.33	430-	30=-
27 SLF 27 SCF		<b>53</b>	3.33	0.63	0.03	1643.	2395.			1179	19.	3.00	٠.১১	3.33	220.	244.
27 567		54 22	J. 60 J. 60	0.00 J.JJ	0.33	1620.	2350.			1235	139	U. CO	0.00	0.00	221-	240-
,.,	2040	••	,,,,,	0.00	3.33	1500.	2262.		25,	P Liui	110	0.00	0.30	6.30	215.	24>.

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		STATISTICS BASED (URDINATES	SED UN UPTIMIZATIUN KEGIUN TES 5 THRUUCH 40)	ATION RE	ISTICS BASED UN UPTIMIZATIUN REGIUN (UKUINATES 5 THRUUGH 40)		
	SUM UF	EQUIV	MEAN	TIME TO CENTER UF MASS	D LAG C.M. TC S C.M.	PEAK FLUM	TIME OF PEAK
PRECIPITATION EXCESS		2.002		14.09			
CUMPUTED HYDROGRAPH UBSERVED HYDROGRAPH	41464. 41721.	0.962	1152.	25.83 26.15	6-15	1827.	23.00 19.00
DIFFERENCE Percent difference	-237.	-0.005	-1-	-0.32	-0.32	50.	.0.4
STANDAF OBJECTIVE F	JARD EKKÜK FUNCTIUN	1.54.	AVEKAGE	AVEKAUE PERCENT	154. AVENAVE ABSULUTE ERRUR 114. 163. AVENAGE PERCENT ABSULUTE ERROR 9.35		
	-	STATISTICS BASI (URDINATES	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES I THROUGH BO)	HYJRUGAA! H 803	на		
	SUM UF FLOWS	ELUIV OEPTH	MEAN	TIME TO CENTER OF MASS		PEAK FLUW	TIME UF PEAK
PRECIPITATION EXCESS		2.002		14.69			
COMPUTEU HYDROGRAPH UBSERVEU HYDROGRAPH	76175.	1.707	952. 905.	40.50	20.87	1627.	23.00
DIFFERENCE Peacent difference	3781.	3.088	.1.	2.19	2.19	50.	30.4
STANDAK P. 18.15.T1VF	AKU EKKOR FUNCTION	207-	AVENACE	AVERAGE	ABSULUTE ERKUR	172.	

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		1 900	10	3.14	3.01	0.19	254.	254.	•			1170	<b>5.</b>	J. GJ	3.03	J.0J	871-	1055.
		2003	11	0.00	0.00	0.00	300.	407.	•			1203	51	0-06	0-03	6-00	914-	1021-
		\$100	12	0.19	3.01	0-19	490.	583.	•			13.00	56	0.30		0-00	440.	994.
		5502	13	0.10	0.01	3.34	051.	143.	•			1433	53	0-60	0.03	0.33	765.	¥57.
		5300	14	9-10	0.01	0.04	414.	454.	•			1270	>>	3.10	J.UL	0-09	994.	420-
		0000	15	C-10	3.01	3.04	401.	9/1.	•			LOUU	> 5	J_00	0.30	0.00	1911-	440.
		3100	10	3.10	0-01	3.09	Heo.	1176.	•			1100	54	J.Ci	0-00	2-30	1030.	854.
		0200	17	C- 19	0-01	0-19	1314.	1329.	•			1833	51	3.00	v-03	0.00	1452-	414-
		0 300	10	J. 00	စ. လ	3.00	1464.	1491.	•			1933	>4	J. 66	J.03	J-0J	1057.	782.
		3400	15	3.03	0.00	J. JJ	1505.	1060.	•			2330	2,5	2-03	0.33	3-03	1049-	752.
		3503	Zu	3.00	0.00	0.00	16/4.	1777.	•			21 JU	•	2.01	0-60	0.00	1031-	734.
		3643	21	3.00	0.00	0.30	1744.	1777.	•			2200	• 4	3.33	0.60	J-0J	1010-	700.
		2703	22	0.63	0.00	0.33	1744.	1040.	•			5300	• <	3.75	0.00	0-00	987.	•75.
		1800	23	0.00	0.00	0.00	1424.	Lodü.	•			ಎಂ	• 4	0.30	3.,,	0.30	942.	<b>645.</b>
		0.00	24	8. CO	0.00	0.00	1427.	1530.	•			6103		3.33	J.JJ	0.33	935.	<b>•20.</b>
		1000	25	3.03	0.00	J. 30	iell.	1402.	•			05 00	•>	0.33	0.60	0.33	403.	>92.
		1100	20	0.00	0-00	0.30	1770.	1473.	•			00 و ل		0.00	0.00	0-00	870-	505.
		1500	27	3.00	0.30	0.00	1731.	1491.	•			<b>&gt;-</b> 33	• 7	0.00	J.0J	0.33	<b>636.</b>	547.
		1 300	24	3.31	3.60	9.03	1673.	1491.	•			<b>35 30</b>	• •	3.30	4-33	0.03	#0#-	523.
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		1063	31	3.00	J.00	0.30	1491.	1491.	•			0=33	11	0.00	3.50	0.30	722.	451.
		1133	32	0.00	0.00	0-00	1442.	1442.	•			<b>0</b> ~JJ	14	3.00	3.30	0.03	+74-	431.
		1430	93	3.00	0.00	0.01	1340.	1473.	•			LJUJ	(,	3.63	9-03	٠.٠٠	470.	411.
		1.33	34	c.co	0.00	3.03	1334.	1464.	•			1100	14	0.03	3.,0	0.0-	****	391.
		5077	35	3.00	0.00	0.00	1540.	1455.	•			1500	15	0-20	0-00	J.03	<b>622-</b>	374.
		2133	36	0.00	3.00	0.00	1243.	1437.	•			1333	/•	0.33	3-03	3.30	•30-	367.
		2200	27	3.00	3.00	٥. ن	Llyd.	1417.	•			1400	!!	J. 0J	0.00	0.00	570-	434.
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	Ÿ	STATISTICS BASED (CHOINATES	) UN UPTIMIZATIUN KEGIGN ; 5 THKUUCH 541	ATLUN KEG H 543	STATISTICS BASED UN UPTIMIZATIUN KEGIGN (GRÜINATES 5 THKUUCH 54)		
	SUM OF FLOWS	Equiv 06.PTH	HEAN	TIME TO CENTEK UF MASS	LAG U.M. TU	PEAK FLUm	TIME OF PEAK
PRECIPITATION EXCESS		3*029		24.50			
COMPUTEU HYDRÖGRAPH OBSERVED HYDRÜGRAPH	104121-	2.082	1928.	42.25	12.75	4225. 3938.	45.00
UIPFERENCE PERCENT DIFFERENCE	.00.0	0.00.0	·	1.32	1.32 11.56	1.29	0.00
STANDAK UBJECTIVE F	AKD ERRUR FUNCTION	317.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERRUR ABSULUTE ERRUR	.224. 16.43	
		STATISTICS BASED (ORDINATES	SED ON FULL HYDRUG S 1 THROUGH 114)	ON FULL HYDRUGRAPH I THRUUCH 114)	STATISTICS BASED ON FULL HYDRUGNAPH (ORDINATES 1 THROUGH 114)		
		EQUIV OEPTH	MEAN FLOW	TIME TO CLNTER OF MASS		PEAK FLUm	IINE UF PEAK
PRECIPITATION EXCESS		3.029		24.50			
COMPUTED HYDROGRAPH OBSERVEJ HYDROGRAPH	155692-	3.113 2.810	1300.	52.12	23.25	4225. 3934.	45.00
UIFFERENCE PERCENT DIFFERENCE	15169.	0.303	133.	10.6	3.07	247.	00-0
STANDAK	JAKO EKROK	342.	3244324	AVERAGE	ABSULUTE ENROR	244.	

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154.	720.		451.	<b>622.</b>	>41.	265.	5 30 .	>14.	
467.	443.	722.	402.	J43.	365.	147.	331.	115.	300.
244.	272.	254.	241.	235.	224.	214.	204.	196.	16>.
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4 SEP	1433		9.03	0.00	0.00	254.	2>4.	:	21	SEP	0303	>=	9.00	0.00	0.00	2070.	2391
• SEP	1900	Ž	3.00	0.00	0.00	241.	250.	•			0400	37	3.03	3.00	3.30	2550.	2174
	2033	3	3.93	0.00	0.00	234.	242.	•			0503	. 0	0.00	3.55	J.J0	4429.	1971
. SEP	2100	4	4.00	0.00	0.00	231.	432.	•			J6 U0		3.33	3.30	0.30	2310.	1765
4 SEP	2200	•	3.00	0.00	J. 3J	223.	224.	•			1733	62	0.00	3.05	0.30	4405.	1011
SEP	2300	•	3.30	0.00	0.00	210.	210.	•			3433		3.00	0.00	3.30	2101.	1+55
SEP	0000	,	0.00	0.00	0.00	204.	204.	•		SEP			3.00	0.00	3.33	2302.	1320
S SEP	0110		0.00	0.00	0.03	٤٥٤٠	236.	•			1000	65	3.00	3.03	1.33	1407.	1503
SEP	0200	9	3.00	0.00	3.03	Lv5.	202.	•			1100		3.00	0.00	0.00	1017-	1095
SEP	4300	16	3.03	0.00	0.00	144.	206.	•			1570	67	0.35	0.00	3.33	1731.	1002
	0433	ii	0.03	0.03	0.00	142.	210.	•			1300		3.30	3.00	3.30	1047.	921
	0503	īž	0.02	0.02	3.30	174.	224.	•			1400		3.33	0.00	3.30	1272.	855
	0400	13	3.11	0.05	3.04	172.	230.	i			1500	70	4.33	0.30	3.00	1997.	
	0733	14	0.16	0.05	0.11	i;;;	201.	•			1030				0.00		405
	0803	iš	3.03	3.00	3.30	1 42.	367.				1113	71	3.30	3.00	0.00	1427.	703
	3900	ió	3.20	0.35	0.15	217.	108.	:			1400		0.00	3.00		1359.	729
	1003	17	5.12	0.05	0.07	265.	637.	-			1433	7.5	3.00	0.00	5-10	1245-	648
	1100	iė	3.11	0.0>	0.36	324.	951.	:				15	3.00	4.03	3.33	1234.	062
	1233	19	3.22	0.05	0.10	920. 411.	1173.				2000	/5	3-00	3.00	0-30	1170-	•7•
	1 303	žó	3.13	0.35	3.35	210.	1293.				5100	/6	9.00	3-00	3.33	1125.	•••
	1400	21	C. 06	0.0>	0.01			:			2233	"	3.00	3.03	2.30	1007.	•>1
	1500	22	0.03	0.03	3.00	633.	1383.	:			2300	7.	3.00	0-00	3.00	1017.	•3/
	1603	23	3.17			151.	1422.	-			0133	/4	5.30	٠.٠٠	0-00	969.	623
	1730	24		0.05	3.12	603.	1+34.	•			3133	•0	3.30	0.00	3.33	923.	•04
	1633	25	0.06	0.05	3.02	1010.	1437.	•			0200	e L	9-00	J.00	0.33	440.	544
	1470			0.05	0.15	1134.	1434.	•			23.22	44	9.33	0.30	3-11	434.	5/3
		24	0.11	0.35	0.04	1251.	1413.	•			3430	• 3	0.00	0.03	0-30	544.	550
	2000	27	9.15	0.05	0.10	1302.	1422.	•			0>00		3.00	4.00	3-33	/el-	545
	2100	20	3.04	0.0	0.00	1444.	1473.	•			0000	<b>&gt;</b>	0.00	3.60	9.33	725.	530
	5500	29	2.04	0.35	7-05	1555.	1535.	•			0100	••	0-00	0.00	0.00	ear-	513
	5307	30	3.10	0.05	0.05	1636.	1578.	•			0400	• 7	0.00	0.00	0-00	<b>454</b> -	499
	3003	31	0.11	3.35	3.00	1706.	1035.	•			3430	••	J. JJ	3.33	3-30	<b>621.</b>	484
	0103	35	3.23	0.35	3.17	1772.	1722.	•			1303	47	J-03	3.00	3.30	544.	463
	2520	33	C. 34	3.05	0.31	1851.	1797.	•			1133	40	0.00	U.00	0.03	>70-	442
	0300	34	0.+3	0.05	0.38	1 764.	1035.	•			1500	71	0.30	J.00	0.30	543.	420
SEP		35	3.10	0.05	3.11	2139.	1857.	•			للذا	76	3.30	0.00	0.33	51/.	340
SEP		36	0.32	0.05	0.21	2246.	1905.	•			1-00	¥3	0.00	0.00	0.30	***.	360
SEP		37	3.14	0.05	9. LL	2493.	1965.	•			1530	**	0.00	3.30	J-J3	413.	367
SEP		36	9.22	0.05	0.10	2124.	2163.	•			1000	*5	3.33	00	J.00	447.	354
	0=00	39	3.14	J. 25	3.11	2963.	2498.	•			1100	**	4-03	3.03	33	420.	143
766		40	3.01	0.31	3.00	3270.	2404.	•	40	SEP	1=33	71	3	J.JJ	0-77	406.	336
SLP		41	9.10	J. 05	3.13	3511.	<b>3224.</b>	•	24	SE P	Lvuu	٧.	0.00	3.03	0.30	s47.	327
<b>463</b>	<b>4230</b>	57	3.00	3. 33	0.33	2434.	2614.	•			1130	114	3.33	0.00	0.00	207.	230.

TOTAL RAINFALL . 4.46, TUTAL LOSS . 1.43, TUTAL EXCESS . 3.03

 PEAR PLUS
 TIME
 MAXIMUM AVERAGE FLUS

 (CFS)
 (MM)
 6-MR
 20-MR
 72-MR
 113.JJ-MR

 0223 05.JO
 (CFS)
 6134 30.00
 1997 1370 

 (FMLMES)
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 (AL-FI)
 2052 0759 11860 12840

CUMULATIVE AREA - 11.50 Su HE

## STATION CLAIRA

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10						HEC-1 I	INPUT						PAGE
NEWTUWN CREEK 10-09-76  UNIT GRAPH (CLARK) AND LOSS RATE OPTIMIZATION LUNIFORM)  LOUIT GRAPH (CLARK) AND LOSS RATE OPTIMIZATION LUNIFORM LOSS RATE OPTIMIZATION LUNIFORM LOSS RATE OPTIMIZATION LUNIFORM LOSS RATE OPTIMIZATION LOSS RATE OPTIMIZATI	E E	10.	1	2	3	,	· · · · · · · · · · · · · · · · · · ·					10	
LUNIT GRAPH CLEAKE AND LOSS KATE OFTIMIZATION CONTEUKNIT  4.2 65  E. 3.44  S. 42 65  E. 3.44  S. 42 65  E. 3.44  S. 42 6.00  S. 60 0.00  S	-	2	ž	ENTUWN CR	EEK 10-0	9-16	:						
FL 3.44  AF .00 .00 .00 .10 .10 .00 .00 .00  .20 .30 .00 .00 .00 .00 .00 .00  .20 .30 .00 .00 .00 .00 .00 .00 .00  .10 .10 .10 .10 .00 .00 .00  .10 .10 .10 .10 .00 .00 .00  .10 .10 .10 .10 .00 .00 .00  ELMINA 32 32 32 32 32 32 32 32 34  15 110 156 164 46 46 46 45 55 60 64  16 110 156 110 110 100 99 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 100 99  178 120 110 110 110 90  178 120 110 110 110 90  177 10 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	~ ~	2 =	5	AIT GRAPH		AND LUSS	KATE Q	INICAL	IND NOI	7 CX X 2			
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32       32       32       32       33       34         35       37       44       46 <t< td=""><td>4</td><td>*</td><td>40.14</td><td></td><td></td><td></td><td></td><td></td><td>FRACE</td><td></td><td></td><td></td><td></td></t<>	4	*	40.14						FRACE				
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123   114   116   110   104   103   100   99   99   104   115   118   118   220   463   670   865   1055   1288   1435   1555   1765   1940   2042   2105   2100   2055   1940   1803   1678   1505   1313   1120   965   865   792   692   652   620   593   555   530   694   695   695   692   692   693   695   69	- T	3	2	011	156	lod	167	75.1	152	142	136	130	
104     115     138     178     220     463     670     865     1055       1248     1435     1555     1765     1940     2042     2105     2100     2055       1900     1803     1678     1565     1313     1120     965     865     792       692     652     620     593     555     530     494     455     448       402     356     350     318     306     298     290       272     268     260     252     248     0     0     0       EL     5P     AF     47       AF     47     .06     .47       1.00     0     0     0     0       77.5     0.     .47       -9.     -12.       -1.     -1.	2 2	9	123	811	110	017	709	101	000	3	66	100	
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3.4	33	רח											
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	****	STATISTICS BASED (UNDINATES	STATISTICS BASED ON UPTIMIZATION REGID (URDINATES 42 THROUGH 65)	LATION REST	REGION		***************************************
	SUM DF FLOWS	EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	C.N. TO	PEAK	TIME OF
PRECIPITATION EXCESS		1.388		38.62			
CUMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	31,577.	0.627	1307.	50.52 56.44	17.90	2155.	56.00
DIFFERENCE PERCENT DIFFERENCE	-1-	-0.000	.0.	0.01	0.07	50.	00.0
STANDARD	ARD ERROR FUNCTION	162. 149.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERROR ABSULUTE ERROR	128.	
	***************************************	STATISTICS BASED (ORDINATES	3ASED ON FULL HY FES I THROUGH	ON FULL HYDROGRAPH 1 THROUGH 95)	STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES 1 THROUGH 95)	**************************************	000000000000000000000000000000000000000
	SUM OF FLOWS	EQUIV OEPTH	MEAN	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK	TIME OF PEAK
PRECIPITATION EXCESS		1.388	;	38.62			
COMPUTED HYDRUGHAPH OBSERVED HYDROGRAPH	66285. 49627.	1.325	698. 522.	56.02	17.40	2155.	56.00
DIFFERENCE PERCENT DIFFERENCE	16658.	0.333	175.	-3.98	-3.98	50.	00.0
STANDARU UBJECTIVE FUI	IND ERROR FUNCTION	153. 183.	AVERAGE	AVERAGE	ABSOLUTE ERROR	95.	

			•		1 140 °C 40	PERLOD				<del></del>		
	14.	Z+2.	519.		1253.	1541.	Less.	2107.	2230.	2214.		
	2124.	2001.	1441.	1742.	1664.	1501.	1445.	1404.	1325.			
	6477.	1104.	1345.	985.	920.	0/5.	825.	111.	132.			
	650.	413.	5/4.	544.	513.	444.	454.	429.	40>.	301.		
	354.	134.	114.	101.	244.	267.	252.	237.	724.	ai.		
	Lvv.	107.	176.	100.	157.	140.	119.	131.	124.	117.		
	110.	lus.	74.	47.	47.	62.	ìi.					
	41.	37.	>4.	51.	40.	42.	44.	40.	34.			
	39.	12.	10.	20.	20.	25.	24.	22.	21.			
	17.	17.	10.	16.	15.		• • • •	•••				
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KT 1500			0.00	30.	ñ.	·	9 OLT 1500	51		0.04 0.11	1146.	12
LT TABLE			uo 0.00	29.	12.	•	4 OCT 1900	52	0.10	5.08 - 5.5i -	1433.	14
£1 1703			JO J.00	28.		4	4 067 1700	51		0.04 0.01	1044.	is
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KT 2000						:		>>				
C1 5000				25.	33.		4 OF 1 5000	54		0.00 0.00	2147.	50
				24.	34.	•	a OCL STAN	57		0.04 0.01	2155-	51
ET 2200			00 0.40	24.	34.	•	4 UCT 2200	Sid		0.00	\$100.	51
FL 5300			09 0.01	23.	35.	•	9 UCT 2300	57		0.00 0.00	2034.	20
CT OGOS			08 0.11	31.	47.	•	TO OFL GOOD	60		0.00	1954.	19
CT V100			36 0.21	64.	44.	•	10 nC1 0100	41		0.00	1001.	19
CT U200			.00 0.00	141.	**.	•	TO OCT 9500	62		0.00 0.00	1767.	10
CT U30U			30 3.30	237.	44.	•	TO OCT 0300	43		0.00 0.00	1675.	10
CT 0430			0.01	349.	53.	•	10 OCT 0400	••		0.00 0.00	1545.	150
LT U500			J4 0.01	470.	55.	•	LO OCT 0500	4>		0.00	1497.	13
CT 0400	14 0	o v.	JO J.30	>86.	60.	•	10 ACL 3000	46		u.oo u.oo	1413.	LL
CT 0700			JO 0.00		44.	•	10 001 0100	•7		0-00 0-00	1332.	9
CT Jugg			JU U.00	750.	64.	•	TO TO A OBOD	48		0.03 0.00	1255.	
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	SUM OF FLUMS	EGUIV	HEAN	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK FLOW	TIME OF PEAK
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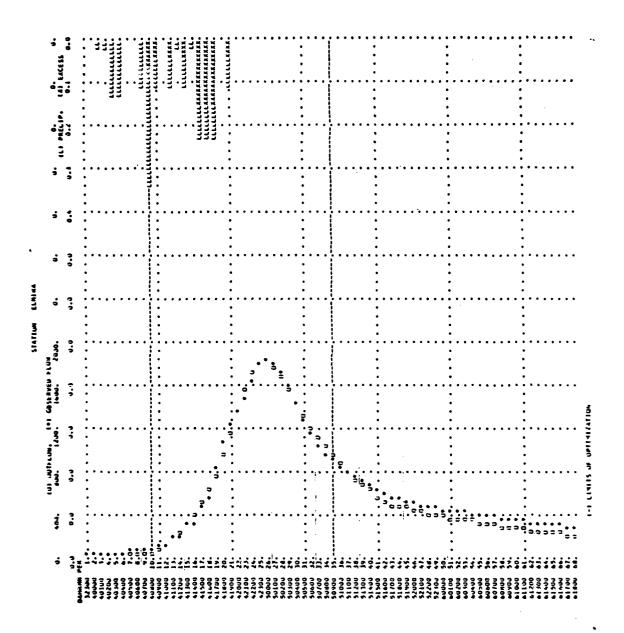
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		STATISTICS BASED (DRDINATES	ES LO THKOUGH	UN UPTIMIZATION REGIUM LO THROUGH 35)	STATISTICS BASED UN UPTIMIZATION REGIUM (Urdinates 10 through 35)		
		EDCIV	MEAN	TIME TO CENTER OF MASS	UM OF EQUIV MEAN CENTER C.M. TO PEAK TIME OF FLOWS DEPTH FLOW OF MASS C.M.	PEAK FLUM	TIME OF
PHECIPITATION EXCESS		9.714	:	16,33			i
CUMPUTED HYDRUGRAPH UBSERVED HYDRUGRAPH	25830. 25673.	0.516	993.	25.81 25.74	77.6 84.6	1829.	25.00
DIFFERENCE PERCENT DIFFERENCE	157.	0.003	٥	10.0	0.07	-13.	20.0
STANDARD UBJECTIVE FUN	ARU ERROR FUNCTION	54. 54.	AVERAGE	AVERAGE ABO	ABSOLUTE ERROR ABSOLUTE ERROR	**************************************	
		STATISTICS BASED (URDINATES	ASED ON FULL HY ES 1 THRUUGH	ON FULL HYDRUGRAPH 1 THRUGH 68)	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THRUUGH 68)		
		EQUIV DEPTH	MEAN	TIME TO CENTER UF MASS	UM DF EQUIV MEAN CENTER C.M. TO FLOW PEAK TIME OF FLOWS DEPTH FLUW UF MASS C.M. FLOW PEAK	PEAK FLOW	TINE OF PEAK
PRECIPITATION EXCESS		0.714		16.33			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	39765.	3.795 0.834	585. 614.	34.17	16.89	1829.	25.00
DIFFERENCE PERCENT DIFFERENCE	-1961-	650.0-	-53.	-0.95	-0.95	-13.	00.0
STANDARU	ARU ERROR	57.		AVERAGE	ABSOLUTE ERROR	<b>.</b> 8+	!

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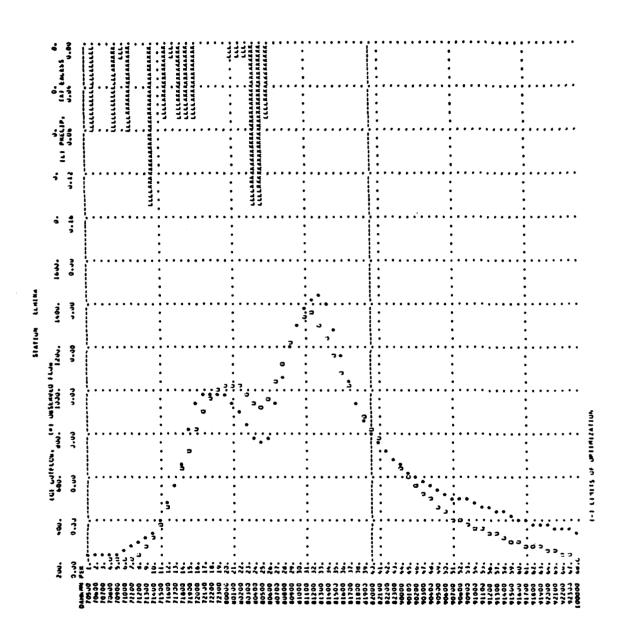
(AC-FT) 351. 2325. LUJULAT (VE AKEA # 17.5) 34 MI



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4U 508 501 492 480 468 452 444  4U 398 301 492 480 468 452 444  PT EL SP AF HCK PM AF HOR PM -85 -15  6A 77.5 0.  BF 248. 410. 1.0340  CLU5005	17	9	752	712	5/0	169	620	100	760	żοż	545	575	
43 398 349 382 370 367 369 554 PT EL SP AF HCK PM .46 .06 .47 .01 PR AF HOR PM .85 .15 dA 77.5 0. dC -815. LU5005	22	<b>3</b>	508	105	492	480	468	452	***	431	<b>/</b> I <b>/</b>	406	
PT EL SP AF PN .46 .06 .47 PN .85 .15 BA 77.5 0. BF 244. 410. 1.0340 UC -815.	23	3	348	395	382	370	367	300	324	745	၁	0	
PW .46 .06 .47 PR AF HOR BM .85 .15 BM 77.5 0. BF 244. 410. 1.0346 UC -815.	<b>*</b> 2	PT	£L	g.S	AF	HCK							
PR AF HOR BA 77.5 0.2 BF 244. 413. UC -815. LU5305	25	9	94.	90.	24.	10.							
PM .85 .15 BA 77.5 0. BF 248. 410. UC -815. LU5005	20	A d	AF	HOR									
64 77.5 0. 8F 244. 410. UC -815. LU5005	2.7	Z.	.85	.15									
8F 244 413. UC -815. LU5305	24	ΨO	17.5	0									
UC -815. LU5005 22	2.5	a.	244.	410.	1.0340								
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		*********	CONTRIBON OF CONTOLLO AND UBSERVED HYDRUGRAPHS	EKVED MYD	OLICIO ON COLOCULO DE COLOCULO		
	-	STATISTICS BASED (OKUINATES	SED UN UPTIMIZATION KEGIUN IES 1 THRUUGH 40)	ZATION KEU	STATISTICS BASED UN UPTIMIZATION MEGIUN (OMJINATES 1 THROUGH 43)		
	SUM UF FLOWS	EQUIV	MEAN	TIME TO CENTER UF MASS	C. 1. TO	PEAK FLOW	TINE OF
PRECIPITATION EXCESS		0.736		17.04		<u> </u>	
COMPUTED HYDROGRAPH OBSERVEJ HYDROGRAPH	31599.	0.632	790.	25.24 25.16	8.20 8.11	1351.	31.00
UIFFERENCE Pekcent difference	-98. -0-31	-0.002	-2.	3.08	0.08	-63.	20-1-
STANDARD EKROK 81.  AVEKAGE ABSOLUTE ERROK 6.2.  AVERAGE PERCENT ABSOLUTE ERROK 8.2.	STANDARD EKROK OBJECTIVE FUNCTION	81. 65.	AVERAGE	AVEKAGE PERCENT	ABSULUTE ERRUR ABSULUTE ERROR	62. 8.22	
	***************************************	STATISTICS BASED (URDINATES	JASEU ON FULL HYDROGIES I THROUGH 631	ON FULL HYDROGRAPH I THROUGH 681	STATISTICS BASED ON FULL HYDROGRAPH (URUINATES I THROUGH 68)		
	SUN OF FLONS	EQUIV DEPTH	MEAN FLU#	TIME TO CENTER OF MASS	LAG C.M. TU	PEAK	TIME OF
PHECIPITATION EXCESS		0.736		11.04			
COMPUTED HYDROGRAPH DBSERVED HYDROGRAPH	43047. 45521.	0.861	633. 664.	32.25 33.52	15.21	1351.	31.00
UIFFERENCE Percent difference	-2474.	-0.349	-36.	-1.21	-1.27	-83.	-1.00
STANDARD OBJECTIVE FUN	ARD EARDA FUNCTION	87.	and and and and and and and and and and	AVERAGE /	ABSULUTE ERRUR	72.	

												•	5 580	1247.	1150.	***	, , ,	3	152.	717.	2	• > 7 • • > 7 • • • • • • • • • • • • •	7.74	>	505.	545.	525.	208		• 00	+0+	• 55.			•0•	7.00		170	307.	340.	754	7.4.	•••••						
												•••••••	COMP	1155.	1001	2101		, ,	176.	727.	0	637.	7.5	523.	.00.	.76.	.30	200	140.	307.	355.	•	120	110.	.649	20%	200	747	253.	244.			•••••••						
	2417.	1716.	410.	307.	3			20.	}			•	E ACESS	70.0	3.0	0.0	200	00.0	0.00	0.30	00°	9	000	2	0.00	00.0	20.0	0 :	9	20-5	0.0	30.0	000		37.3	7.0	200	7	0.0	0.00	07.7	2.5	••••••						
	.74	.10		787				;	•				5507	00.0	00.0	9	200	0	00.0	0.00	0.0	2 0	00.0	00.0	3.03	0.00	00.0	9 6	00.0	00.0	3.0	30.0	900	000	60.0	0.00	30	20.0	00.0	0.0	0	9	********						
	. •												RAIN	0.30	30.0	9 :	200	0.0	3.30	000	٥٠ ٠	900	2	3,0	0.00	2.00	00.0	3	20.0	0.00	0.0	0.0	900	9	0.00	9	9	207	3.0	0.0	0.0		•						
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165		.63.	.22	11:	7			<b>*</b>			ELMINA	•	JA MUN MKAN																										-	(077 1	٠.	_	******			Ĭ	•	2 4	
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-PERIOD	2047	1 599	3	**	707	701	3	2			11 STALLUN		• •	•	•	• •	•	•	•	•	• •	•	•	•	•	•	• •	• •	•	•	•	• •	• •	•	•	• •	•	•	•	•	• •	• •	•			72-114		3 76	
100 100 100 100 100 100 100 100 100 100	1067.	1713.	<b>#</b> 61.	432.	217.	10%	55.	28.			HYDRULKAPH AT	••••••	7 590	248.	246.	26.5	248.	.962	273.	250	176	391.	+11.	503.	656.	816.		100	996	412.	465		777	760.	766.	.076	1,000	1296.	1343.	1420.		13461	*********	0.35, TUTAL EXCESS -		AVEKAUE F 72	•		
	1142.	1835.	422.	+63.	233.	7.7	24	20.		*************	DAN	*************	DEST.	248.	240·	224.	219.	218.	420.			346.	*09*	552.	.10	127.	216		1000	1025.	1029.		937.	454.	404		1225.	1 304.	1350.	1351.			••••••	0.35. 10	1	MAXINGM AVERAUL	1065.	7117	,
	73.	1961.	.984	4.00.	244.	125.	63.	32.					EXCESS	80.5	86		0.32	00.0	90.0	9.0	2	20.0	6.0	0.03	0.05	o. 05	60.0	900	3.0	3.00	9.6	33	0.13	0.13	50.0	36	90	00.0	00.0	00.0	3		•••••••	A LUSS .		#X-0	1249.	63.0	
	363.	.901	.25.	532.	267.	134.	.;	÷	17.			•	1055	3.	300	30	8	0.01	70°0	000	300	6.0	0.02	10.0	2.02	20.0	300	9 9	9.0	0.00		500	0.05	0.05	2.02	9 ?	88	3	o. 3	0.0	96	3	********	1.09. TGTAL LUSS			(CFS)	10-F13	
		•	_										H I	0.00	0	90.0	0.0	0.0	3	9	3 -	. 3	0.07	3.01	2.63	6.0	36	9	0.00	9.0	5.5	5 6	2.15	3.15	2.0	3 6	50	3.0	00.0	0.0	9 6	3	******					: =	
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										•			HAR	2500	000	0400	2040	3	2	3	27	1503	1,000	1 700	250	0063	3 5	2200	2 300	2000	000	250	000	2500	0000	3	250	200	301	227		3	•••••	TURAL	100	(CFS)	<u>:</u>		
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~	3	Š	IT CRAPH	UNIT GRAPH (CLARK)	AND LUSS KATE UPTIMIZATION (UNIFORM)	KATE UP	TIMIZAT	TON CONT	DKM)				
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52	H.F.	78.	420.	1.0346								
70	S	-10.	-13.									
27	יר יר	-:	08									
79	77											

	3						
	•	STATISTICS BASED (UKDINATES	ON UPTIMIZATION KEULUN 1 THRUUGH 251	ATION KEN			
	*	SUM LF ELUIV MLAN CENTLK FLOMS OF MASS	MLAN	TIME TO CENTER OF MASS	LAG PEAK TINE OF C.M. TO FLOW PEAK	PEAK FLOm	TIME OF PEAK
PRECIPITATION EXCESS		0.927		0			
COMPUTED HYDROGKAPH UBSERVED HYDRUGKAPH	29709.	0.594 0.595	1146.	15.00	α.υ.α υ.υ.α	2081.	12.00
UILFERENCE PERCENT UIFFERENCE	-37.	-0.001	-1-	44.0	0.44	-34.	-3.00
STANDARD UBJECTIVE FU		205.	AVERAGE	AVEKAUE PERCENT	ABSULUTE ENRUR Absulute enrun	171.	
		STATISTICS BASED (ORDINATES	D UN FULL	UN FULL HYDKUGKAPH I THKÜUCH 41)	STATISTICS BASED UN FULL HYDKUGKAPH (ORDINATES 1 THKÜÜCH 41.)		
		Equiv DEPTH	MEAN FLUM	TIME TO CENTER UF MASS	TIME TO LAG SUM UF EQUIV MEAN CENTER C.M. TO PEAN TIME UF FLOWS DEPTH FLUW UF MASS C.M. FLUW PEAK	PEAN FLOW	TIME OF PEAK
PRECIPITATION EXCESS		0.927		6.69			
CUMPUTEJ HYDROGRAPH UBSERVED HYDROGRAPH	41019.	0.820	1000.	20.25 18.75	13.56	2041. 2115.	12.00
DIFFERENCE PEKCENT DIFFERENCE	3616.	0.072	* :3 2	1.50	1.50	-34-	-3.00
STANDARD OBJELTIVE FU	ARD ERROR FUNCTION	219.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERRUR ABSOLUTE ERRUR	194.5	

											••••••		•	2 2 2	1100.	400		170-		654	•13.	\$65.	**	510.	• • • •	• 66.	.37.	*1.	404	349.	374.	304	351.		;			
											•••••••		••••••	CLAP .	1372.	12.00	11/6.	1117.	1050	941.	430.	:	634.	708.	;;	703.	.693	626.	592.	254	524.		471.		,			
	1441.		74.	10	104	•		÷	÷		••••••			ERCESS	0,00	2.00		20.0	3.0	77.0	20.0	0.0	20.0		00.0	00.0	00.0	0.00	2.0	60.0	20.0	20.0	00.0		;			
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165	2371.					:	.,	.;	۶.		••••••	רראואא	••••••	AUTO MERIN		-	0050 >	_		_																		
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Mi Colase	2406.	746.	. 36	144			:	;	٠. د.			STATION		• • •	••	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•		. •		•	
-	2165.		*	240				÷	25.		**********	HYDRUGHAPH AT STATION	***************************************	7 787	78.	~~	7.	P.1.7	155.	467.	405.	1253.	1341.	74.47	1054.	1.797.	1411.	2011.	2099.	2115.	20.30	1930.	1703.	1231	1295.	7637		
	::::	191	.92.	776	153.	,	•	,	۶۷.	.5.	**********	H	***************************************	COMP	78.	<b>.</b> 2	73.	\$	124.	265.	>31.	446	1242.	1047.	1905.	2045.	2041.	2050.	1970.	1867.	163.	1004.	1571.		14031	•		
	1112.	7/01	275	-00			;	<u>.</u>	. P.Z	•	••••••		•••••••	EXCE SS	0.00	20.5	00.0	90.0	0.23	0.23	21.0	٠. د . د	0.10	0.0	0.0	3.0	o.6	0.00	o	00.0	2.0	0.0	00.0					
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(-) LIMITS OF OPTIMIZATION

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- 7 ~	235	N O	NEATOWN CREEK 10-28-81 Unit Graph (Clark) And LOSS RATE UPTIMIZATION (UNIFURM) 60 260CT81 0500 83	EEK 10-2 (CLARK) 0500	8-81 ANU LOSS 83	RATE UI	TALIMITAT	ION CONT	FURMJ			
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	1	COMPARISON OF COMPUTED AND OBSERVED HYDROGRAPHS	UTEU AND OBS	ERVED MYDROG	COMPARISON OF COMPUTED AND OBSERVED HYDROGRAPHS		
		STATISTICS BASED (URDINATES	SED UN UPTIMIZA TES 25 THRÜUGH	UN UPTIMIZATION REGIUN 25 THRUGH 651	STATISTICS BASED UN UPTIMIZATION REGIUN (URDINATES 25 THROUGH 65)		
		EQUIV DEPTH	HEAN	FINE TO LENTER OF MASS	LAG C.M. TO C.M.	PEAK Flow	FINE UF PEAK
PRECIPITATION EXCESS		1.531		45.69			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	50546.	1.011	1233.	52.76 51.47	10.07	3012.	53.00
DIFFERENCE PERCENT DIFFERENCE	-6-0-01	0000-	•0-	0.89	0.89	-32.	00.0
STANDARD DBJECTIVE FUN		184.	AVERAGE	AVERAGE AB	ABSOLUTE ERROR ABSOLUTE ERROR	137. 20.56	
		STATISTICS BASED (URDINATES	SASED ON FULL MY TES I THROUGH	ON FULL HYDRUGRAPH 1 THRUUGH 83)	STATISTICS BASED ON FULL HYDRUGRAPH (URDINATES 1 THRUUGH 83)		
		EQU IV DEPTH	MEAN	TIME TO CENTER UF MASS	246 C.M. TO C.M.	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.531		45.69			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	70019.	1.400	844.	57.80 55.60	15-11	3072.	53.00 53.00
PERCENT DIFFERENCE	5462.	0.109	• 99	7.20	2.20	-32.	00.0
STANDARD OR LECTIVE FILM	ARD ERROR	197.	7.52 A 7.74	AVERAGE	ABSULUTE ERROR	136.	

				UNIS	T HYDROGRA	Pri			
				92 END-0	F-PERIOD U	ROIMATES			
64.	sid.	454.	1052.	1478.	1 801.	2144.	2124.	2 140.	2251.
2044.	1975.	1858.	1748.	1445.	1540.	1434.	1370.	1289.	1211.
1145.	10/4.	1066.	***.	445.	842.	144.	745.	IQL.	<b>660.</b>
421.	504.	550.	517.	447.	454.	+31.	405.	341.	359.
334.	318.	249.	261.	245.	249.	23%.	220.	201.	LYS.
184.	173.	143.	153.	144.	135.	127.	120.	115.	134.
100.	94.	84.	45.	īd.	74.	6 V.	45.	41.	54.
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SA RON	HRM	URD	RAIN	LOSS	EXCESS	COMP 0	UBS 4	:	DA	MUN	HRMM	URD	. RAIN	LOSS	EXCESS	COMP 0	085 4
24 UCT		ı	4.00	0.00	U. 00	34.	34.	·	27	JCT	2300	4.5	0.07	9.04	3.01	354.	341.
24 OC1		2	0.01	0-91.	0.00	33.	34.	•			0000	44	0.20	0.04	9.12	344.	322.
24 OCT		5	9.00	0.00	0.00	32.	34.	•	28	UC T	0100	45	0.34	9.04	0.30	335.	419.
26 OC 1		4	9.00	U.00	0.00	31.	34.	•	28	uC I	JZGG	44	4.41	0.04	0.34	445.	<b>630.</b>
24 OC I		5	9.00	0.00	0.00	<b>30</b> .	34.	•			3303	47	0.41	9.04	0.34	713.	1164.
26 OCT		•	0-04	0.09	0.00	29.	35.		24	UCT	U400	46	0.25	0.04	0.17	1045.	1372.
24 OCT		1	4.60	4.00	0.00	28.	35.	•	28	UC T	J500	49	0.04	U.08	0.01	1537.	1479.
24 OCT		•	9.00	6.30	0.00	21.	36.	•	24	OLT	0000	Su	0.04	0.04	0.00	2315-	1733.
26 OC T	1 300	9	0.00	6.00	0.00	26.	J4.	•	28	DC T	0700	51	0.05	0.04	0.00	2457.	2182.
26 OCT		LO	0.00	0.00	0.00	25.	36.	•	24	UL F	0800	>4	9.03	9.00	0.00	2804.	2113.
50 OC L	L 500	11	0.00	0.00	U.69	24.	37.	•	26	OC T	0900	53	0.00	0.00	0.00	3012.	3034.
26 OCT	1600	12	0.00	0.00	0.00	23.	37.	•	26	OC T	LOUD	54	0.01	U.U1	0.00	3072-	3104-
26 DCT	1 700	13	9.00	J-00	0.00	23.	37.	•	28	OLI	1100	55	0.00	0.00	0.00	3010.	304 %-
24 OCT	1800	14	0-04	0.04	0.00	22.	37.	•	26	OCT	L200	54	0.00	0.40	0.30	2073.	2955.
26 OLT	1900	15	0.00	0.00	0.00	21.	37.	•	20	OC F	1300	57	J. 00	0.00	0.00	2711.	2748-
24 OCT	2000	İ	0.00	0.00	0.00	21.	38.	•			1400	54	0.00	0.00	0.00	2551.	2552.
26 OCT	2100	17	0.00	0.00	0.00	20.	39.	•			1500	59	0.00	0.00	0.00	2401.	2357.
26 OCT	2200	10	0.00	0.00	0.00	20.	41.	•			1000		0.00	0.00	0.00	2259.	2147.
26 OCT		19	0.09	0.08	0.01	ži.	42.	ė			1700	ě.	U.00	0.00	0.00	2120.	1445.
27 OCT		20	0.01	0.01	_0.00	24.	45.	•			1400	62	0.00	6.00	0.00	2000-	1745.
27 OCT		Zi	0.09	0.08	0.01	29.	48.				1400	43	0.00	0.00	0.00	1882.	1562.
27 OCT		22	0.17	0.04	0.09	46.	52.				2000	84	0.00	0.30	0.00	1771.	1367.
27 OCT		23	0.03	0.03	0.00	ii.	54.				2100	65	0.00	0.00	0.00	1407.	1256.
27 001		24	0.10	0.08	0.03	121.	60.	•			2200	66	J. 00	0.00	0.00	1564.	1143.
27 OCT		25	0.08	0.08	0.00	174.	71.	·			2300	67	0.00	0.00	9.00	1476.	1049.
27 OCT		26	9-10	0.08	0.03	233.	100.				3000	•4	9.00	0.30	0.00	1388.	974.
27 OCT		27	0.09	0.06	0.01	291.	281.	•			9100	69	0.00	0.00	0.00	1301.	920.
ZT UCT		28	0.12	0.08	3.04	340.	412.	•			9200	70	3.00	J.0J	0.00	1229.	466.
27 OCT		29	0.00	0.00	0.00	39 7.	523.	·			0300	71	0.00	0.00	0.03	4157.	\$22.
27 001		30	0.00	0.00	0.00	434.	625.				9444	72	0.00	0.00	0.00	1084.	776.
27 061		31	0.05	0.00	0.00	459.	722.	ě			0500	ï.	0.00	0.00	0.00	1024.	739.
27 OCT		32	0.01	0.01	0.00	475.	153.	·			4640	75	0.00	0.00	0.00	980.	703.
27 OCT		33	0.00	0.00	0.00	482.	728.				2700	75	0.00	3.00	0.00	947.	445.
27 001		34	J.00	0.00	0.00	440.	***	•			0600	7.	3.00	4.00	9.00	410.	636.
27 OCT		35	0.00	0.00	0.00	468.	591.	:			0400	ii	0.00	0.00		445.	613.
27 001		36	0.00	0.00	0.00	452.	526.	:			1000	74	9.00	0.00	0.00	855.	591.
27 001		37	0.00	0.00	ø.00	437.	447.	:			1100	79	0.00	0.00	0.00	627.	572.
27 001		36	0.00	9.00	0.00	422.	427.				1200	40	0.00		0.00	749.	345.
27 OCT		39	0.00	0.00	0.00	408.	389.				1200			J. 00			510.
27 001		40	J.08	0.00	J.00	374.	359.	:			1400	6L	9.00	0.00	0-30	712.	475.
27 OCT		41	0.00	0.00	4.00			:						0.00	0.00	141.	
27 001		42	0.07	0.00	1.00	301.	334.	•	24	•	しちいひ	43	0.00	0.03	0.00	122.	446.
er we	4440	44	0.01	0.07	3.00	56 <b>4</b> .	316.	•									

TOTAL RAIMFALL . 3.31, TOTAL LOSS . 1.76, TOTAL EXCESS . 1.53

PEAK FLOW	TIME			MAXIMUM AVE	RAGE FLOW	
1CF51 1072.	(MR) 53.00	(CF5)	6-HR 2013.	24-1 <b>6</b> 0 2037.	72-14R 763.	82.0J-HR 649.
		(INCHES)	0.347	0.487	1.347	1.342
		IAC-FT1	1434.	4079.	3732.	5755.

CUMULATIVE AREA - 17.50 SQ ME

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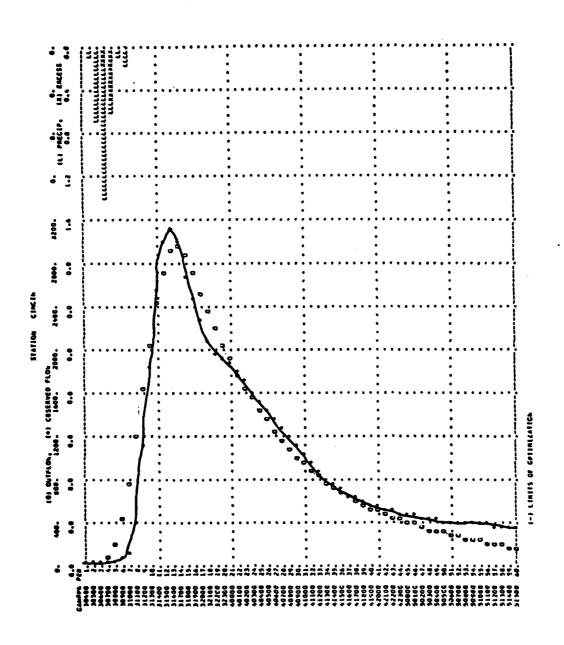
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PG       LINC       3.60         PG       PLN       3.17         PG       PLN       .04       .05       .10       .00       .00         PI       .05       .44       .91       .40       .05       .10       .00       .00         PI       .05       .44       .91       .40       .05       .01       .01       .01       .00       .00         PI       .05       .04       .07       .01       .01       .01       .01       .01       .00       .00       .00         RK       CINCIN       53       52       52       57       66       102       419       1130       10         UU       554       3010       2675       2460       2265       2090       1964       1180       10       40       428       44       44       44 <td>•</td> <td>9</td> <td>ZI J</td> <td>1.31</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td>	•	9	ZI J	1.31										
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UU     2549     3010     310     2675     2460     2265     2090     1964       UU     1873     1810     1705     1600     1516     1432     1348     1260     1180       UU     1024     947     872     408     746     700     652     595       UU     546     525     504     486     473     461     449     437     428       UU     401     395     389     383     384     374     372     366       PM     180     64     ACLA     .02     .02     .02     .03       BA     167     0     -13     .0163       LU     -1     -1     -1     -1       LU     -1     -1     -1	•	9	53	53	25	25	57	99	107	614	1130	1824		
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STATISTICS BASED ON CPTIMIZATION REGION (ORDINATES 1 THROUGH 60)		STATISTICS BASED (ORDINATES	STATISTICS BASED ON CPTIMIZATION REGION (ORDINATES 1 THROUGH 60)	ATION REG H 60)	STATISTICS BASED ON OPTIMIZATION REGION Iordinates 1 Through 601		
9	SUM OF FLCNS	EQUIV DEPTH	MEAN FLOW	TIME TO CENTER OF MASS	LAG C.M. TO C.M.	PEAK FLOW	TINE OF PEAK
PHECIPITATION EXCESS		0.627		4.80			
CCMPUTED HYDROGRAPH CBSERVED HYDROGRAPH	59776.	0.630	986.	22.64 24.63	17.84	2970. 3110.	13.00
DIFFERENCE PERCENT DIFFERENCE	-0-0-	000*0-	01	-1.99	-1.99	-140.	1-00
STAN	STANDARD ERROR OBJECTIVE FUNCTION	197.	AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE ERROR	135.	

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3 MAR 1200	<b>-</b> <u>-</u>	9.6	9 6	9 6	1620.	1130.	• •	A MAR		6 ¢		96	8 6	572.	595
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-		167	334	364	317	3.84	165	165	165	391	395	
9		195	349	405	614	420	454	420	474	416	*04	
2		400	399	345	391	399	4 36	255	680	970	1140	
07		1475	9661	2846	4000	5143	2800	5946	2944	8165	5750	
17		5535	5428	5330	5358	5386	5346	1475	5036	4760	4431	
77		4070	3790	3530	3765	3067	285+	4697	87.47	2310	7166	
<b>53</b>		2059	1961	18/7	1793	1723	1653	1540	1534	1481	1430	
<b>5</b> 7		1304	1336	1584	1555	1504	1169	1126	1040	1048	101	
25		<u> </u>	LINC	DERU								
٥2		81.	99.	•10								
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		STATISTICS BASED UN OPTIMIZATION REGIUN (URDINATES 35 THROUGH 65)	UN OPTIMIZATION REGIUN 35 THROUGH 651	ATION REGI H 65)		•	
:		SUM OF EQUIV MEAN CENTER FLUMS DEPTH FLOM OF MASS	MEAN FLOW	TIME TU CENTER OF MASS	C.H. TO PEAK	•	TIME OF PEAK
PRECIPITATION EXCESS		7.04.7		37.18			
COMPUTED HYDROGRAPH UBSERVED HYDROGRAPH	117938.	1.243	3804.	52.18 52.53	15.00	6350.	00-84
DIFFERENCE PERCENT DIFFERENCE	-566.	-0.006	-18	-0.35	-0.35	404. 6.79	2.00
STANDARU OBJECTIVE FUN		371. 392.	AVERAGE	AVERAGE A Percent A	ABSOLUTE ERROR ABSOLUTE ERRUR	322. 13.20	
		STATISTICS BASED ON FULL HYDROGRAPH (ORDINATES 1 THROUGH 90)	D ON FULL HY 1 THKOUGH	ON FULL HYDROGRAPH I THROUGH 90)			
:		EQUIV OEPTH	MEAN FLOM	TIME TO CENTER UF MASS	•	PEAK FLOM	TINE OF PEAK
PRECIPITATION EXCESS COMPUTED HYDRUGRAPH UBSERVED HYDRUGRAPH	180040.	1.802 1.802	2000.	55.1d 55.1d 56.14	18.00	6350.	48.40
DIFFERENCE PERCENT JIFFERENCE	9097. 5.32	960*0	101.	16.0-	-0.97	404. 6.79	2.00
STANDARI OBJECTIVE FI	STANDARD ERROR	243.		AVERAGE	ABSOLUTE ERROR	158.	

				UNI	I HYDROGRA	PH			
				117 END-0	F-PERIOD U	ROINATES			
154.	582.	1202.	1937.	2000.	3251.	1632.	3710.	3561.	3414.
1259.	1104.	2466.	2429.	2699.	2574.	2454.	2342.	2234.	2152.
2033.	1940.	18>0.	1/05.	1684.	1404.	1534.	1401.	1394.	1310.
1269.	1210.	1154.	1101.	lusu.	1002.	456.	912.	470.	630.
142.	755.	120.	407.	425.	425.	596.	>64.	543.	51#.
444.	./1.	444.	424.	404.	390.	312.	355.	339.	323.
308.	244.	284.	207.	255.	243.	232.	221.	211.	202.
197.	183.	175.	167.	159.	152.	145.	156.	132.	iżo.
120.	114.	109.	104.	44.	95.	90.	46.	84.	76.
75.	71.	64.	65.	42.	59.	56.	54.	<b>51.</b>	49.
4/.	45.	42.	41.	34.	37.	35.	34.	32.	31.
24.	26.	21.	25.	24.	23.	22.			

						MY	DKUGRAPH A	. 57411	UN	LIN	CEM						
*****	•••••	*****	••••••	******	********	********	••••••	••••••	****	••••	•••••	•••••	*******	••••••	*******	•••••	•••••
JA MÖN	MAM	_0¥9_	HAIM	-1035	Excess	COMP U	005 4		DÁ	ÑÜN	MRMN	URJ	'RALA"	1025	EXCESS	COMP S	OBS Q
7 06.1	1400	ı	0.00	0.00	0.00	115.	115.	•	9	OC F	1600	46	0.12	0.08	J.03	>259.	5404.
7 00 7	2000	Z	0.12	0.10	0.02	114.	215.	•			1700	47	U. 00	0.00	0.00	544#.	5944.
7 007		3	0.12	3.44	0.43	124.	118.	•			[800	48	0.00	0.33	0.00	<b>6217</b> .	5848.
/ UCT			0.00	0.00	0.00	152.	121.	.•	•	UCT	1 400	49	0.00	0.00	0.00	6350.	2414.
	2300	` `3	3.30	D.04	0.00	Løb.	125.	•			2000	50	0.00	0.00	0.00	6274.	5750.
6 467		•	4.12	3.04	0.01	224.	139.	•			5100	51	0.00	0.00	0.00	6064.	5562.
# CC !		,	0.12	0.34	3.33	282.	156.	•			2200	52	0.40	0.00	0.30	5809.	5428.
# OCT		•	0.12	0.04	0.03	344.	115.	•			2300	53	0.00	0.00	0.00	5552.	5330.
a uct		. 4	0.00	0.00	0.00	423.	204.	•			0000	54	0.00	U. 0U	U-00	5299.	5358.
8 OCT		- †ô	0.00	0.00	0.00	493.	248.	•			0100	55	0.00	U.00	0.00	5056.	5340.
# ULT		- TI	2.00	3.00	u.00	554.	297.	•			0200	56	0.00	0.00	0.00	4425.	5386. 5247.
# OC F		12	0.00	0.00	0.00	401.	334.	•			3300	57	0.00	0.00	0.00	4604. 4393.	5034.
8 UCT		13	0.30	0.00	0.00	627.	364.	•			0400	58	0.00	0.00	0.00	4192.	4760.
a uct		14	0.00	0.00	J.00	624. 614.	377. 384.				3500	59 60	J.0J	0.00	3.00	4300.	**31.
# UCT		16	0.00 0.12	0.00 0.00	U.00 U.03	607.	391.	•			0700	•1	0.00	0.00	J.00	3817.	4090.
# OC 1		17	3.03	0.00	0.00	544.	391.	•			0100	•ż	0.00	0.00	0.00	3643.	3/90.
4 001		14	3.00	0.00	0.00	590.	391.	•			0900	6.5	0.00	0.00	0.00	3474.	3530.
8 007		19	0.00	0.00	0.00	545.	391.	•			1000	64	0.00	0.00	0.00	3317.	1545.
8 OCT		20	3.03	0.03	3.30	584.	395.				1100	65	0.00	0.00	U-00	3165.	3067.
d uci		31	0.00	0.00	U.00	584.	195.	·			1200	44	0.00	0.00	U.00	3021.	2854.
# OC 7		22	0.12	0.00	3.03	567.	399.				1300	67	0.00	0.00	0.00	2883.	2454.
uc i		75	0.00	u.00	4.00	586.	402 ·				1400	4.6	0.00	0.00	0.00	2751.	2478.
# UCT		24	0.00	0.33	0.00	585.	413.	•			1500	64	0.00	0.00	<b>U.</b> U0	2020.	2310.
6 OCT		25	J.0J	0.00	J.J0	547.	420.	•			1600	70	0.00	0.00	U.00	2506.	2166.
8 001		26	a.au	4.60	0.00	590.	424.				1700	71	U-00	0.00	U. 00	2391.	2059.
6 UCT		27	J.U0	0.00	0.00	590.	428.	•			1 800	72	U.00	J. 00	0.00	2282.	1401-
& UCT		20	U.00	0.00	u.03	501.	424.	•			1400	73	0.00	0.00	0.00	2178.	1877.
8 UCT		21	J.00	0.00	U.00	571.	414.	•	"Ìò	ÜČT	2000	74	0.00	0.00	0.00	2074.	1793.
9 001	0000	10	u.00	0.00	3.00	544.	409.	•	10	UC I	4100	75	0.00	0.00	0.04	1964.	1723.
9 007	0100	31	0.00	0.00	0.00	555.	404.	•	10	120	2200	76	0.00	0.00	0.00	1844.	1453.
9 001	<b>#230</b>	32	0.00	0.00	0.00	544.	397.	•	ĹO	UCT	2300	77	J. 00	0.00	0.00	I tos.	L590.
9 061	J 500	33	4.12	g. us	0.33	534.	395.	•			0000	7#	0.00	0.00	0.00	1725.	1534-
4 001		34	0.12	0.04	U. U3	529.	391.	•			0100	14	0.00	0.00	v.00	1047.	1461.
T UCT		35	0.24	0.04	0.15	524.	399.	•			0500	B)	0.00	0.00	9.00	1572.	1436.
9 001		36	J. 60	0.00	U.00	6L7.	434.	•			0300	8 L	0.00	0.00	9.00	1500.	1384.
4 OC1		37	J.12	0.04	U. 33	749.	532.	•			0400	82	0.00	0.03	U.00	1435.	1334.
9 156.1		38	J.24	4.04	U.15	427.	640.	•			3500	43	0.00	0.00	0.00	1367.	1294.
4 001		39	9.24	0. 14	0.15	1164.	870.	•			0000	84	0.00	0.00	0.00	1305.	1525-
• u.i		40	0.24	9.00	U.15	[462.	1146.	•			0700	85	0.00	0.00	0.00	1200.	1204 -
TOO P		41	0.35	0.38	0.27	1440.	1475.	•			0400	#6	J. 65	0.00	0.00	1100.	1144-
9 061		42	J. 35	0.06	0.27	2323.	1996.	•			0900	87	J.00	0.00	0.00	1135.	1124-
+ ULT		**	0.4/	J. UN	0.39	2944.	2446.	•			1000	44	0.00	0.00	0.00	1064.	1090-
9 001		44	0.24	0.08	0.15	3711.	4000.	•			1100	84	0.30	0.00	0.00	1035.	E048-
4 001	1 200	45	u.00	0.00	J.00	4520.	5143.	:	11	ULT	1500	90	0.00	0.00	0.00	v48.	1014-

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		STATISTICS BA	TICS BASED UN UPTIMIZATIUN REGION (URDINATES I THROUGH 50)	ATIUN REG	STATISTICS BASED UN UPTIMIZATIUN REGION (URDINATES L'THROUGH 50)		
		EQUIV DEPTH	MEAN	TIME TO CENTER OF MASS	SUM OF EQUIV MEAN CENTER C.M. TO PEAK TIME OF FLOW PEAK	PEAK FLOW	TIME OF PEAK
PRECIPITATION EXCESS		1.645		16.18			
COMPUTED HYDRUGRAPH DBSERVED HYDROGRAPH	106233.	1.120	2125.	31.22	15.03 14.94	3461. 3391.	32.30
OIFFERENCE PERCENI DIFFERENCE	- 787.	800*0-	-16.	0.10	0.10	70.	7.00
STANDARD OBJECTIVE FU	STANDARD ERRUR OBJECTIVE FUNCTION	89. 82.	AVERAGE	AVERAGE PERCENT	) ERRUK 89. AVERAGE PERCENT ABSOLUTE JAROR 70. Inctiun 82. average percent absolute eardr 5.31	70. 5.31	
		STATISTICS BASED (URDINATES	IASED ON FULL HYDRUG ES 1 THROUGH 701	ON FULL HYDRUGRAPH I THRUUGH 701	STATISTICS BASED ON FULL HYDRUGRAPH (ORDINATES 1 THRUGH 70)		
		EQU IV DEP IH	MEAN	TIME TO CENTER UF MASS	C.H. TO	PEAK FLOW	TINE OF PEAK
PRECIPITATION EXCESS		1.645		16.18			
COMPUTEU HYDRUGRAPH OBSERVED HYDROGRAPH	141390.	1.490	2020.	38.08	22°12 21°91	3461.	32.00
UIFFERENCE PERCENT OIFFERENCE	-5-	-0•000	•0-	0.21	0.21	2.06	1.00
STANDARD	ARD ERROR	91.		AVERAGE	ABSOLUTE ERROR		

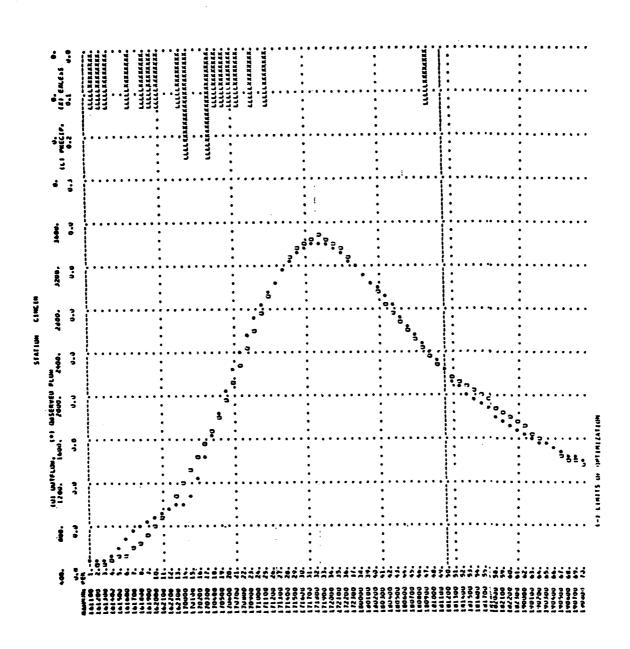
	UMII NTOKUGRAPH	
60	ENU-OF-PERIOD ORDINATES	
	ACT OWF . 0.44	
	100 OHE - 0.77	

				AG	LUME - 0.9	y			
37.	134.	244.	464.	674.	447.	413/.	1 340 .	1617.	1864.
żūśü.	2223.	2354.	2447.	2494.	2476.	2407.	2320.	2249.	2174.
2102.	2032.	1964.	1449.	1635.	17/4.	1715.	Lebel.	1643.	1549.
1448.	1448.	1400.	1353.	1308.	1265.	1222.	1142.	1142.	1104.
1064.	1932.	448.	964.	932.	901.	#7L.	842.	814.	101.
761.	/36.	711.	487.	445.	442.	421.	600.	540.	561.
544.	>24.	507.	440.	4/4.	458.	443.	428.	414.	400.
367.	375.	361-	344.	336.	320.	315.	30>.	293.	205.
275.	206.	257.	244.	241.	231.	225.	217.	210.	203.
1 76.	190.	1#3.	177.	171.	100.	LaO.	155.	150.	145.
140.	135.	131.	126.	122.	110.	114.	110.	107.	103.
LUJ.	94.	45.	40.	47.	44.	41.	79.	76.	74.
71.		65.	44.	<b>62.</b>	60.	54.	50.		52.
31.	99. 49.	41.	46.	44.	43.	41.	.0.	<u>54.</u> -	37.
34-	35.	34.	34.	32.	30.	24.	24.	24.	27.

						н	DROGRAPH A	T STAT	t ON	ÇLN	Cim				-		
*****	•••••	*****	••••••	••••••	•••••	•••••	••••••	******		••••	*****		*******	*****	••••••	********	*******
DA MUN	HRAN	ORU	RAIN	F022	EXCESS	COMP W	065 9		DA	MON	MARN	ORD	RAIN	LUSS	ExCESS	COMP Q	UBS 3
16 OCT		1	3.00	0.00	3.00	460.	460.	·			2200	36	0.00	0.00	0.00	J354.	#301 ·
		2	0.13	J.05	U. 08	455.	444.	•			2300	37	0.00	0.00	9.00	3284.	3254.
16 DCT		•	0.13	0.05	0.06	459.	440.	•			0000	36	J. 00	0.00	0.00	3210.	3193.
LG OCT	1400	•	0.13	0.05	J. 08	475.	512.	•			0100	34	J.00	0.00	0.00	3127.	3130-
L& OCT		5	0.00	0.00	0.00	502.	585.	•			0200	40	0.00	0.00	u-00	<i>3</i> 038.	3058.
14 OCT		•	J. JU	0.00	0.00	534.	675.	•			0300	41	0.00	0.00	0.00	2947.	2984.
TO UCL		_ I	. 0.13	0.05	0.00	583.	<i>1</i> 57.	•			0+00	42	0.00	0.00	0.00	2854.	2902.
to OCT			0.00	3.00	0.00	<b>638.</b>	#15.	•			<b>U500</b>	4.5	J.00	0.00	6.00	2163.	2622.
14 OCT		¥	U.13	0.05	0.00	704.	855.	•			0600	44	9.00	0.00	0.00	2675.	2734.
16 ULT		10	0-13	0.05	0.06	783.	<b>#90.</b>	•			0 700	45	0.00	0.00	J.00	2540.	2644-
10 OCT		11	0.15	0.05	U. 04	àle.	931.	•	14	OL T	9400	46	J.00	0.00	0.00	2507.	2558.
IN UCT		15	0.00	0.00	J.JQ	477.	969.	•	1.6	OC T	9900	41	0.00	0.00	0.30	2428.	2410-
16 UCT		11	0.00	0.00	0.00	1082.	991.	•			1000	48	0.13	0.05	0.08	2354.	2342.
L7 OCT		14	0.13	J.05	0.04	1149.	LUL4.	•			LLOU	49	9.00	9.00	0.00	4266.	2302.
FS UCL		15	0.2>	U.U5	0.21	1305.	1084.	•	18	UC T	1200	50	0.00	0.00	U. 00	2228.	2222.
17 OCT		16	J.05	0.00	U.00	1433.	1234.	•			1300	51	0.00	9.00	0.00	21/3.	2134-
17 OCT		11	<b>0.33</b>	0.00	0.00	1540.	1+36.	•	1.8	OC T	1400	52	U-00	0.40	0.00	2122.	20/3.
17 007		10	0.25	3.45	4.21	1647.	1053.	•			1200	53	0.00	0.40	4.00	2075.	2010.
17 001		14	.U-13	0.05	0.06	1619.	1454.	•	4.6	UCT	1-00	54	J.00	0.30	3.00	2031.	1401-
L7 OCT		20	0.13	ひ。ひち	0.36	1900.	2045.	•	1.0	OL T	1700	55	0.00	8.00	0.00	1941.	1705.
A DCT		21	0.13	J. US	4.04	2113.	2230.	•	16	OL T	F # 00	56	0.00	0.00	0.00	1952.	1870.
L7 OCT		22	4.13	0.05	0.04	2215.	2398.	•	1.6	UCT	1900	57	U. 03	0.03	0.00	1914.	1435.
TA OCL		23	<b>u.00</b>	4.03	J. 30	2438.	2550.	•	1.0	UCT	2000	54	0.00	0.00	0- Ou	1674.	1793.
LT UCT		24	0.13	0.35	J.38	2549.	2710.	•	18	UL T	2100	54	J.00	0.00	0.00	1834.	L7Se.
17 OC.1		25	J.03	J.03	0.00	2753.	2844.	•	10	UCT	4500	60	0.00	0.30	0.00	1/93.	1725.
LT INCT		24	0.13	J. JS	U. 08	2900.	2464.	•	10	OC F	2300	-1	0.00	0.00	0.00	1750.	ledi.
13 061		21	0.00	0.00	J. JO	3040.	3058.	•	19	OC T	0000	62	0.00	0.30	0.00	1706-	Late.
L/ ULT		23	3.30	U.JU	0.00	3144.	J157.	•	14	UC T	9100		3.30	0.00	0.00	1457.	tell.
17 OCT		2+	J.00	0.00	J.00	3277.	3234.	•	19	JC T	0490		0.00	0.00	0.00	1000.	1576.
17 UC 7		30	0.00	J. JO	U- DO	3356.	3319.	•	19	OL T	0100	65	0.00	U. 00	U.00	1555.	1541.
L7 OCT		31	u.00	3.30	U. 34	1414.	J364.	•	19	100	0400	60	U.00	9.00	0.00	1500.	1507.
IT OLT		32	0.00	0.00	J.J0	1450.	3591.	•	19	OL T	0500	67	0.00	J. 00	6.30	1459.	1981-
17 UCT		3 5	U.00	0.00	U.00	1401.	3391 .	•	19	OC T	0600		0.00	0.00	U- 00	1413.	1455.
17 OCT		34	J. 00	0.00	0.03	3445.	1362.	•			0700	44	0.00	0.00	0.00	1364.	1422.
11 001	210-)	3>	0.00	0.00	J. JU	1400.	1346.	•			0 600	70	0.00	0.00	0.00	1362.	1391.
								•									

FOTAL MAINHALL . 2.42, FCTAL LUSS . U.Fd. TUFAL EXCESS . 1.44

PEAR PLUM	TIME			MARIMUM AVE	RAGE FLOW	
(CFS)	****		6-M4	24-HM	72-HR	64.0J-HK
14-1.	52.00	(CFS)	\$422.	3044.	2030.	2034.
		(INCHES)	0.416	3.773	1.461	1.441
		14C-F11	1697.	a034.	11610.	Heto.



25 25 26 26 26 27 26	767477	5ca-10444N2	LINE
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HUCUR 1.000 147. 126.	CINCIN 926 1800 2986 2358 2358	01 60 60 61 60 61 61 61 61 61 61 61 61 61 61 61 61 61	
1600. 121.	920 2246 3040 2270 1639	015ELIC XI 04APR78 2 2 35 1.26 -92	
1.0163	926 2590 3067 2174 1597	VER /	
	931 2798 2101 255 1555	-05-78 AND LOSS KATE GPTIMIZATION (UNIFUKM) 52 -00 -10 -10 -10 -2 -2 -2 -2	laareen 2. 2. 2. 2. 2. 2. 2. 2. 2. 2. 4. 2. 2. 2. 2. 2. 2. 2. 0. 2. 2. 2. 7. 7. 2. 2. 2. 2. 4.
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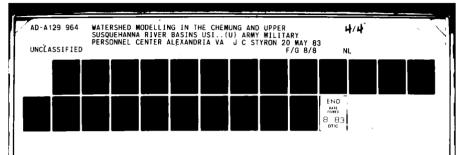
10 UNSELIC KIVEN 4-05-74  10 UNIT UMAPH (CLARK) AND LOSS KATE CPTIMIZATION (UMIFUKM)  11 LO 04APR74 1500 52  10 1 2  10 2 35  10 1 2  10 2 35  10 1 2  10 2 35  10 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0								:			
UNIT GRAPH (CLARK) AND LOSS MATE CETIMIZATION (UNIT-UR, UNIT GRAPH (CLARK) AND LOSS MATE CETIMIZATION (UNIT-UR, UNIT GRAPH (CLARK) AND LOSS MATE CETIMIZATION (UNIT-UR, UNIT-US)	<u>-</u>	J	.7	••••••			•••		, p		01
CIN 1.2b DERU .92 MUCUR CIN 1.2b DERU .92 CIN 1.2b DERU .92 MUCUR CIN 1.2b DERU .92 LO .10 .10 .10 .10 .10 .10 .10 .10 .10 .10	~ :		OTSELIC M	IVER 4	1-05-76	2 214 2	14/14/10	T TOTAL OF THE	1 1		
CIN 1.26  CIN 1.26  CIN 1.26  DERU .92  .00 .00 .10 .10 .10 .10  .10 .10 .00 .00 .00 .00 .00  .10 .10 .00 .00 .00 .00 .00  .10 .10 .00 .00 .00 .00 .00 .00  CINCIN 920 926 931 947 986 1036 1  1800 2246 2590 2798 2918 2997 2906 2966 2596 2590 2798 2918 2998 2970 2782 2  2966 3040 3047 940 1036 1  1800 2246 2590 2774 2918 2999 2970 2782 2  2968 1639 1597 1555 1520 1494 1462 1  1888 1639 1597 1555 1520 1494 1462 1  1888 1639 1597 1555 1520 1494 1462 1  1988 1639 1597 1555 1520 1494 1462 1  1988 1639 1597 1555 1520 1494 1462 1	-	9	04APR78	1300	52						
CIN 1.26 DEKU .92 HUCUR .00 .00 .00 .10 .10 .10 .10 .00 .00 .00 .00 .00 .10 .10 .00 .00 .10 .10 .10 .10 .10 .00 .00 .10 .10 .10 .10 .10 .10 .00 .00 .10 .10 .10 .10 .10 .10 .00 .00 .10 .10 .10 .10 .10 .10 .10 .00 .1	<u> </u>		~	•	1						
CIN 1.2b  DERU .92  HUCUR .00 .00 .00 .10 .10 .10  .10 .10 .00 .00 .00 .00  10 .10 .00 .00 .00 .00 .00  CINCIN 920 926 931 947 98c 1036 1  1800 2246 2590 2794 2997 2906 2795  2356 2270 2174 2916 2997 2906 2  2356 2270 2174 2101 2031 1961 1492  1889 1639 1597 1555 1520 1494 1402  1384 1378 0 0 0 0 0 0  CIN DERU .54 .54  HUCUR 1.00 0 0 0 0  147. 0.  147. 0.  147. 0.  147. 0.  147. 0.  148. 140. 0.  149. 140. 140. 140. 140. 140. 140. 140. 140	5	· S	% %								
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2354 2270 2174 2101 2031 1961 1498 1688 1639 1597 1555 1520 1494 1462 1384 1378 0 0 0 0 0 CIN DENU -46 -54 HUGUR 1.00 0. 147. 0721.	7			3067	3040	2959	2870	2182	2674	2566	2470
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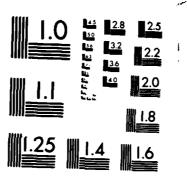
		כנוחליאה נין כן נוחשה מובים אינט נופפה ארוט אלטה ניניפו איני נופפה ארוט אלטה נינים אינים נופפה אלום אלטה היום א	יטובט אינט יטניס	אלואח אלואי	COMPAKISSIA DE COMPUTED ALO CUSSENVED MYDROGRAPHS		
STATISTICS BASED ON OPTIMIZATION REGION (URDINATES 5 IMKUDUM 35)		STATISTICS BASED ON OPTIMIZATION KEULUN LURULNATES 5 THKUUCH 351	EU ON OPTIMIZAT	Allun KEU H 351	אס		
	SUA UF FLOWS	ELCIV	AEAN FLUA	TIME TO CENTER OF MASS	LAG C.M. TU C.M.	PEAK Flus	TIME OF PEAK
PRECIPITATION EXCESS		9.403		10.17			
COMPUTED HYDRUGRAPH LESEAVED HYDROGRAPH	73473. 73461.	0.774	2370.	21.30 21.28	11.10	3265.	14.30
UIFFERENCE PERCENT OIFFEMENCE	9.00	000°0	•	20.0	0-02	198.	02.4-
STANJAKO UBJECTIVE FUN		1+7. 152.	AVEKAUE	AVERAGE /	ABSULUTE ERKÜR ABSULUTE ERKÜR	114. 5.31	
		STATISTICS GASCO (DAUINATES	JASEJ UN FULL HYJKUG ES I THKUUGH 521	UN FULL HYJKUGKAPH 1 THKUUGH 52)	STATISTICS GASLU UN FULL HYDRUGRAPH (GRUINATES I THRUUGH 52)		
		E0.01 V	N. 4 Jr	TIME TO CENTER	LAG L.A. TC	PEAK	TIME OF
	FLOWS	JEPIH	FLUM	UF MASS	C.A.	FLGN	PEAK
PRECIPITATION EXCESS		3.463		10.19			
COMPUTED HYDROGRAPH OBSERVEJ HYDROGRAPH	104504-	1.102	2010.	26.46 26.38	16.27	3205-	14.03
DIFFERENCE Percent difference	157.	0.002	÷	a0.0	80.0	198.	00.4-
STANDARU	STANDARD ERKUR	114.		AVERAGE	ABSULUTE EKKUK	73.	

																	•••••		•	ă		2706.		44.70	2358.	4479.	217.	2101.	2031	1071	701	7.44	1737	1000	1634.	1547.	1555	1520	1462	1430.	1404	1397.	134.	1374.	•						
											•						**************			9		2045.	75.78	2370.	2326.	7449.	71 la.	£107.	2041			- 4261	70.0	1080	1034.	1592.	1506.	151	14.42.	140	1445	1471.	1399.	1376.	••••••						
	2823.	1436.	1320.	.114	679-	450.	244.	201.	1.30.	43.	65.	<b>.</b> 5•	31.		•		••••••			2 4 7 4 3		300	3	20.0	00.0	.0.	0.00	00.0	3	7	5		900	0.00	00.0	20.0	00.0	0.0	9	00.0	00.0	0.00	00.0	0-0	••••••						
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INATE	3131	4168	7007	1020	4	7	3.29	770	155.	2	Ξ	2	*	. ~	ì		:	CINCIN			į	A P	100	A A	A	AP K	A	APA	¥	A P		¥ 5	1	4	APA	A.A	¥.	APA		4	¥ d	¥	APR	A K				-20	4361	¥90°1	85-2
UNIT HYDROLRAPH D-OF-PERIOU OND	•						7	•	161.		٥	7.	· •		•		•••••	NOTE	******		5	•	• •	` ^		^	**	•	•	•	٠.	<b>.</b>		, ,	•	٠	•	•	0 4	•	, a	•	•	9		9.0		21.			
HVD		ž	3	Ó	2	•	*	~	=	=	_	•		, -	•			11 51/		•	•	• •	•	•	•	•	•	•	•	• •	• •	• •	• •	•	•	•	•	• •	•	•	•	•	•	• •		. 55		1 F	. 17.	. 96	. 7 .
TAND-OF	2575.	337.	603.	107	754	517.	355.	245.	10%	114.	78.	24.					•	HYDRUGRAPH AT STATIUN		91	3	986	200		. 246	900.	1036.	1144.	1324.	1569.	200	.0477	2700	2918.	2447	2968.	2936.	2834.	.0167	£ 706.4	300.	30.0	.4567	2610.		0-11, TUTAL EXCESS -		MAKINUM AVENAUE FLUM	₹	-	•
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MICROCOPY RESOLUTION TEST CHART NATIONAL BUREAU OF STANDARDS-1963-A

		STATISTICS BASED (ORDINATES	STATISTICS BASED UN UPTIMIZATION REGIUN (ORDINATES 15 FMRUUGH 40)	UN UPTIMIZATIÜN REGIÜN 15 FHRUUGH 40)	• •		
		EQUIV	HEAN FLOW	TIME TO CENTER OF MASS	•	PEAK FLOM	TIME UF PEAK
PRECIPITATION EXCESS		1. 324		11.25			
CUMPUTED HYDRUGRAPH NBSERVED HYDRUGRAPH	100917.	0.845	3881.	28.76 28.95	11.51	5698.	24.30
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		EQUIV	HEAN	TIME TO CENTER OF MASS	LAG C.M. TU C.M.	PEAK FLUW	TIME OF PEAK
PRECIPITATION EXCESS		1.324		17.25			
CUMPUTED HYDROGRAPH UBSERVED HYDRUGRAPH	149007.	1.248	2129.	35.70 35.08	18.45	5698.	24.00
DIFFERENCE Percent difference	9087. 6.49	0.076	130.	0.62	0.62	250.	0.00
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		7000	3	0.04	0.04	J. 00	72.	75.	•	•	NUV	1100	34	0.00	9.00	0.00	2502.	34
		9100	4	10.0	0.01	0.00	70.	15.	•	•	MJA	1200	19	0.00	0.00	0.00	1121.	31
		9200	5	J.02	0.02	U. 00	47.	76.	•	•	NOV	1300	40	0.00	0.00	0.00	3153.	24
		<b>4340</b>	•	0.07	J.07	0.00	67.	74.	•			1400	41	0.00	4.00	0.00	2991.	24
١	UV	0400	7	0.02	0.02	0.00	66.	76.	•		NOV	1500	42	9.00	0.00	0.00	26 54.	24
١	OV	U500	5	0.08	0.06	J.02	44.	77.	•			Leou	43	0.00	U. 00	0.490	2643.	22
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		0700	10	0.05	0.05	J. 00	110.	80.	•		NOV	1400	45	0.00	0.00	0.00	2425.	14
		0400	11	0.06	0.06	0.00	147.	81.	•	•	MOV	1900	44	0.00	0.00	0.00	2301.	11
		9900	12	0.06	0.04	0.01	195.	45.	•		NUV	2000	47	0.00	0.00	0.00	2104.	14
		1000	13	0.09	0.00	0.03	248.	89.	•	•	NOV	2100	48	0.00	0.00	0.00	2072.	11
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		1400	19	0.23	0.06	0.17	2274.	2300.	•	LO	NOV	0300	54	0.00	0.00	0.30	1513.	11
į	04	1700	20	0.13	0.06	0.07	3080.	3335.	•	10	NOV	0400	55	0.00	0.00	0.00	1434.	10
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		STAFISTICS BASED (URDINATES	SED UN UPTIMIZATIUN REGIUN TES 10 TARUUGH 55)	ATION REU 4 551	STAFISTICS BASED UN UPTIMIZAFIUN REGIUN (Urdinates 10 thruuch 55)		
		EQUIV JEPTH	MEAN FLOW	TIME TO CENTER UF MASS	C	PEAK FLUM	TIME UF PEAK
PRECIPITATION EXCESS		3.583		27.28			
COMPUTED HYDROGRAPH OBSERVED HYDROWRAPH	340783.	2.854	7408. 7432.	37.5u 37.5u	10.28	13774.	41.00
UIFFERENCE PERCENT DIFFERENCE	-1104.	*00*0-	-24•	90.0	0.06	-446. -3.13	-1.00
STAND UBJECTIVE	STANDARU ERROR UBJECTIVE FUNCTION	440. 422.	AVEKAGE	AVERAGE PERCENT	ABSULUTE ERKUR ABSULUTE ERKOR	350.	
STATISFICS BASED ON FULL HYDROGRAPH (DRDINATES I THROUGH 84)		STATISTICS BASEU (URDINATES	BASEU ON FULL HYDROGRAPH TES 1 THROUGH 88)	HYDROGRAF 1 Be)	T.		
	SUM OF FLOWS	EQUIV	SUM OF EQUIV MEAN CENTER FLOWS DEPTH FLOW UF MASS	TIME TO CENTER UF MASS			TIME OF PEAK
PRECIPITATION EXCESS		3.583		27.28			
COMPUTEU HYDRUGRAPH URSERVED HYDRUGRAPH	442682. 445048.	3.70d 3.728	5030. 5057.	43.85 43.91	16.57	13774.	41.00
DIFFERENCE PENCENT DIFFERENCE	-2366.	-0.020	-27.	-0.01	10.0-	-446.	-1.30
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	9100	•	0.03		0.30	490.	466.	·			2100	48	0.00	0.00	0.00	10774.	1008
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	9000	•	0.02		J.00	474.	457.	•			2 100	50	3.00	0.00	0.00	4514.	975
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	1000	13	J.26		0.24	1143.	1447.	•			0600	57	J. 00	0.00	0.00	6087. 2706.	621
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	1000	;;	9.07		J. J4	>624.	5308.	•			1200	43	0.00	0.00	J.00	4135.	340
	1700	20	4.16		0.13	5705.	\$514.	•			1300	64	0.00	0.00	4.40	30/6.	345
	1800	21	3.20		0.23	5/65.	5710.	•				45	4.00	0.00	0.00	3637.	343
	1900	_22	9,04		0.05	5913.	5934.		27	SEP	1500	**	0.00	J. 00	0.00	3411.	323
	\$000	23	3.63	0.03	0.00	6135.	4014.	• -	21	SEP	1400	•7	0.00	0.00	0.00	3199.	301
	2100	24	0.01		0.00	6398.	4044.	•	27	SEP	1700	64	3.00	0.00	0.00	3001.	243
SEP	2200	25	J. 00	0.03	0.03	6623.	6046.				1800	69	0.00	0.00	0.00	5612-	219
	5 MG	26	0.01	3.31	J.J0	6726.	5444.				1900	73	0.00	0.00	0.00	2441.	270
	9000	21	0.0/		U. 04	6664.	5962.			SEP		71	J.00	J.UJ	J-00	2478.	254
	0100	20	_2.02		0.00	<b>4461.</b>	6044.	•		SEP		15	U. UO	0.00	0.00	2325.	243
	JZOJ	24	9.11		J. 06	6205.	6u60.	•		SEP		73	0.00	0.00	0.00	2142.	234
	0300	30	0.17		0-14	6032.	eves.	•		ie P		74	J.30	0.00	v.00	2044.	22
	U40U	31	3.03		0.00	5963.	6022.	•		SEP		75	0.00	0.00	0.00	1973.	211
	0500	32	3.10		J.07	5466.	5434.	:		icf		76 71	u.00	0.00	0.44	1930. 1867.	514
	0600	33	3.31		0.20	6044.	>436.	•		SEP		14	0.00	0.00	0.00	1846.	194
	0700	3.	J. 38		0.35	<b>6455.</b>	4004.	:		SEP		75	0.00	0.00	0.00	1804.	194
	9400	35	0.46		0.43	7144. 8145.	0564.	·			J500	80	0.00	0.00	0.00	1766.	10
	1000	36 37	9.10 9.30		9.07 9.27	9336.	1532. Bozd.	;		SEP		41	J.00	0.00	0.00	1727.	i
	1100	36	0.25		J.22	10062.	14400.			SEP		42	0.00	0.00	J. 55	1447.	10
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	1 300	40	0.09		0.00	12943.	13350.	•		SEP		i,	0.00	U.00	0.00	1016.	17
	1400	41	0.07		0.04	13562.	14110.	•		SEP		#>	0.03	Ü.00	J. JO	1581.	16
	1500	*2	3.00		J.00	13/14.	14105.	•		SEP		Va.	U.U0	0.00	u.0J	1540.	10.
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SEP																	

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PEAR FLOW TIME RANGE FLUW
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LUMULATIVE AREA . 185.00 SO M

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I-F LIMITS OF OPTIMIZATION

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PI         .34         .01         .00	92	ā	71.	.26	67.	.26	90.	ò	Š		70	10.
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UD         457         760         1277         2194         3266         4256         4948         5536         6074           UD         6876         6764         6172         5356         4852         4540         4408         4360           UD         4058         3761         3347         2947         2550         2329         2168         2057         1947           UD         1774         1672         1616         1526         1446         1416         1345         1500         1247           PI         CAN         NEMVAL         HUGUR         MFPL         1025         1004         976         948         920           PR         -18         -31         -31         -31         -31         -31         -31         -40         976         948         920           BF         456         -31         -31         -31         -31         -31         -31         -31         -31         -31         -31         -31         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32         -32 <td>32</td> <td>3</td> <td></td> <td>137</td> <td>135</td> <td>137</td> <td>747</td> <td>141</td> <td>641</td> <td>173</td> <td>718</td> <td>243</td>	32	3		137	135	137	747	141	641	173	718	243
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		STATISTICS BASED (ORDINATES	ED ON OPTIMIZATION REGIUN ES 35 THKOUGH 651	F1UN REG 65)	STATISTICS DASED UN UPTIMIZATIUN REGIUN (ORDINATES 35 THROUGH 65)		•
电转换 计电传统 医慢性 医骨骨 医骨骨骨 医骨骨骨 医骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨骨		EQUIV OE 9 I H	MEAN FLOW	TIME TO CENTER OF MASS	LAG C.M. TO C.M.	PEAK FLOM	TIME OF PEAK
PRECIPITATION EXCESS		1.290		42.92			
CUMPUTED HYJRÜGKAPH JASERVEU HYDRÜGKAPH	107045.	0.897	3453.	53.67	10.75	6956.	50.00
DIFFERENCE PERCENT DIFFERENCE	-541.	-0.005	-17.	0.21	0.21	-14.	-1.03
STANDARU UBJECTIVE FUI	STANDARD EHROR DAJECTIVE FUNCTION	284. 304.	AVERAGE	AVERAGE PERCENT	ABSULUTE ERRUR ABSULUTE ERRUR	225.	
erceccanoscoscoscoscoscoscoscoscoscoscoscoscosco	**************************************	STATISTICS E	STATISTICS BASED UN FULL HYDRUGRAPH (UNDINATES 1 THROUGH 90)	YDRUGRAP	STATISTICS BASED UN FULL HYDRÜGRAPH  (OKDINATES 1 THROUGH 90)		•
*** *** *** *** *** *** *** *** *** **	SUM UF FLONS	EQUIV DEPTH	MEAN FLUM	TIME TO CENTER OF MASS	LA6 C.A. TU	PEAK	TIME OF PEAK
PRELIPITATION EXCESS		1.290		45.45			
COMPUTED HYDROGRAPH OBSERVED HYDROGRAPH	146909.	1.231	1632.	59.11 54.34	16.19	6882. 0956.	50.00 51.00
DIFFERENCE PERCENT DIFFERENCE	1.765.	910.0	19.	0.17	5.02	-14.	1.00
STANDARU	DARU ERROR	209.	6.00	AVERAGE	ABSOLUTE ERROR	122.	

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	700		723.	440.	434.	401.	565.	532.	504.		i.	443.		
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OCT 1700	į	0.00	0.00	0.00	45.	45.	•	9 OCT 1400	44	0.12	0.04	0.04	3714.	425
OCT LEGO	—- ≱	0.01	0.01	J. 00	44.	45.	•	9 UCF 1500	47	0.01	0.01	U.00	4634.	494 553
UCT 1900	3	0.00	0.00	0.00	43.	45.	ě	9 007 1400	40	9.09	0.07	0.01	5424. 6528.	401
NC1 5100 NC1 5000	🏃 🗀	0.04	0.04	0.U0 0.U0	42. 41.	44.	- :	9 OCT 1700	4 9 50	0.00 J.00	0.00	0.00	6677.	650
001 5100	?	0.00	0.00	J.00	40.	**.	•	4 OCT 1900	51	0.00	0.00	0.00	6882.	487
OCT 2300	<del></del>	0.04	0.06	0.00	39.	<del>- 37.</del>	<del></del> -	9 OCT 2000	- 32	0.01	0.01	9.00	4424.	49
0000	i	0.10	0.10	U.00	39.	44.	•	9 JCT 2100	53	0.00	0.00	0.00	6261.	670
OCT 0100		0.12	0.12	0.00	34.	51.	•	9 UCT 2200	54	0.00	0.00	0.00	5494.	617
UCT 0200	10	0.04	0.04	0.00	37.	51.	•	9 OCT 2300	55	0-00	0.00	0.00	5548.	535
OC 1 0300	11	0.00	0.00	0.00	36.	51.	•	10 001 0000	56	0.04	0.04	0.00	5219.	445
OC 7 0400	15	0.01	0.01	J.00	35.	51.	_•	10 OCL 0100	51	0.00	0.00	0.00	4909.	454
OCT 0500	13	0.06	0.04	J. 00	34.	51.	•	10 001 0200	56	0.00	0.00	0.00	4618-	440
OCT 0400	13	0.00	0.00	0.00	34.	51. 52.	•	10 UCT 0300	59 40	0.00	0.00	0.00	4344. 40 <b>8</b> 4.	420
DC1 0800	15	0.00	0.01	0.00	33. 32.	52.	:	10 OCT 0500	40 41	0.00	0.00	9.00	3844.	405
001 0900		0.00	0.00	0.00	-38.	×.	-	10 UCT 0400	62	0.00	0.00	0.00	3616.	374
DCT 1000	14	0.03	0.03	0.00	31.	54.	•	10 OCT 0700	•3	0.00	0.00	8.00	3402.	334
0011 T3d	19	0.00	0.00	U.U0	30.	54.	•	10 041 0400		0.00	0.00	8.00	3200.	- 294
DCT 1200	20	0.01	0.01	0.00	30.	55.	•	10 JCT 0900	65	0.00	0.00	0.00	3010.	255
OCT [300	21	U.00	0.00	0.00	29.	57.	•	10 UCT 1000	66	0.00	0.00	0.00	2032.	232
DCT 1400	25	0.00	0.00	J.00	24.	40.	•	to oct fra	67	0.00	0.00	0.00	2664-	214
DCT LSOO	23	0.04	0.06	0.00	20.	46.	•	10 OCT 1200	**	0.00	0.00	0.00	2506.	205
DCT 1400	-24	0.03	0.03	0.00	27.	79.	<u>•</u>	10 OCT 1300	69	0.00	0.00	0.00	2357.	174
DCT 1800	25	0.00	0.00	0.00	26.	89.	•	10 OCT 1400	70 71	0.00	0.00	0.00	2046.	174
0091 130	27	0.00	3.00	3.00	25.	121.	:	10 OCT 1900	72	0.00	0.00	0.00	1963.	147
DCT 2000	28	0.00	0.00	0.00	25.	139.	:	10 OCT 1700	73	0.00	0.00	0.00	1444.	161
DC1 5100	- ž,	3.00	0.00	Ú.00	24.	(39.	- E	10 UCT 1800	74	0.00	0.00	0.00	1737.	152
DC1 2200	Ĵó	0.06	0.06	0.00	24.	139.	_•	LO UCT 1400	75	0.00	0.00	0.00	1634.	146
DCY 2300	31	0.00	0.00	0.00	23.	139.	•	10 OCT 2000	76	0.00	0.00	0.00	1537.	141
OCT 0000	32	0.00	0.00	0.00	23.	137.		10 OCL 5100	77	0.00	0.00	0.00	1446-	134
ICT ULGO	33	0.00	0.00	0.00	22.	135.	•	TO 0C4 5500	76	0.00	0.00	0.00	1361-	130
0020 730	34	0.00	0.04	0.00	55.	137.	•	TO OCT 5300	79	0.00	0.00	0.00	1280.	124
CT 0300	35	0.10	0.10	.0.00	21.	142.	:	11 001 0000	60	0.30	0.00	0.00	1204.	121
XT 0400 XY 0500	36	0.10	0.10	0.00	21.	147.	<u>-</u> -	TE OCT OSOO	#2 #2	0.00	0.00	0.00	1000.	ii:
CT 0400	36	0.17	0.07	U.00	121.	173.	:	11 001 0200	43	0.00	0.00	0.00	1064.	101
X1 6700	39	3.07	0.09	0.00	224.	218.		11 001 0400	**	3.00	0.00	0.00	1041.	101
ET 0000	40	0.13	0.04	0.04	347.	243.	•	11 201 0500	85	0.00	0.00	0.00	LOLd.	102
DCT 8900	41	0.19	0.09	0.11	549.	457.	•	11 OCT 0600	86	0.00	0.00	0.00	¥94.	100
KT 1100	62	3.25	0.04	0.17	611.	760.		LL UCT 0700	87	0.00	0.00	0.00	974.	71
	43		0.09	0.21	1205.	1277.	•	11 001 0800	66	J.00	J.00	0.00	952.	94
CT 1200	44	0.41	0.09	0.32	1791.	2194.	•	LL OCT 0900	89	0.00	0.00	0.00	932-	92
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		ANISON OF COM	CLINYAKISUN UF CUMPULEU AND UGSEKVED HYDRUGKAPHS	KVEU HYUK	CHAPANISUM UT COMPOLED AND GOSERVED HYDROGRAPHS		
		STATISTICS BASED U (ORDINATES	ED UN UPTIMIZATION KEGIUN ES 5 THKOUGH 30)	A1100 NEU H 301	N UPTIMIZATION KEGIUN 5 THKOUGH 30)		
	SUM OF FLUMS	EQJIV DEPTH		FIME TO CENTER UF MASS		•	TINE OF PEAK
PRECIPITATION EXCESS		1.532		17.12			
COJPUTED HYDRUGKAPH OBSERVED HYDROGKAPH	96061. 95526.	0.805	3695.	20.93	3.21	04 1d. 6494.	14.00
DIFFERENCE PERCENT DIFFERENCE	535. 0.56	90°°	21.	-0.02	-0.02	-16.	00.0
STANDARD 03JECTIVE FUN			AVERAGE	AVERAGE PERCENT	ABSOLUTE ERROR ABSOLUTE ERROR	114.	
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	SUM UF FLOWS	TIME TO EQUIV MEAN CENTER DEPTH FLOM OF MASS	MEAN	TIME TO CENTER UF MASS	LAG C.M. TO C.M.	PEAK	FIME OF PEAK
PRECIPITATION FXCESS		1.532		17.72			
COMPUTED HYDROGKAPH OBSERVED HYDROGKAPH	196727.	1.648	2342. 2363.	36.35	18.63	647d.	19.00
DIFFERENCE PERCENT DIFFERENCE	-1804.	-0-015	-21.	-1.32	-1.32	-16.	00.0
STANDARU ERRUK CHJECTIVE FUNCTION	ARU ERRUK	304.	11 0 W TU D W	AVERAGE	Ad SULUTE ERRUR	232.	

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ier ie			0.07	0.07	9.00	350.	360.	•		P 1000	44	0.05	0.03	0.01	1483.	1794
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EP 23		ii	0.12	0.03	U.09	4157.	1097.	•		1700	53	0.01	0.01	0.00	1807.	2015
EP 00		15	0.35	0.03	0.32	1544.	1712.	•		EP 1600	54	0.11	0.01	9.04	1402.	2049
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EP OZ		14	0.13	0.45	0.10	2997.	3176.	-		EP 2000	36	0.00	0.00	0.00	1798.	2042
EP 03		15	0.14	0.03	0.14	3952.	4014.	•		EP 2100	51	0-00	9.00	0.00 9.00	1823. 1835.	2176
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EP 07		19	0.03	0.03	0.00	6433.	6302.	•		E 0100	61	0.00	0.00	0.00	1458.	5530
EP 08		20	0.00	0.40	3.00	6470.	6494.	•		P 4240	62	0.00	0.00	0.00	1370.	7147
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SEP LL		23	0.00	0.00	0.00	5824.	6004.	•		P 4500	45	0.00	0.03	0.00	1430.	1911
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UCT 1400		01 0.01	0.00	<del>- 424.</del> 528.	372.	<del>-</del> -	18 OCT 0500	- <del>50</del>	0.00	0.00	0.00	2041.	200
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Appendix D -- Statistical Results of Verification Storms

Basin/Storm	% Lag C.M. to C.M.	% Peak Flow	Time to Peak	Avg % Abs Error	0BJ Func
Butternut Ck. 10-18-75	1.37	-14.92	ο.	15.39	248
Canasawacta Ck. 7-31-71	5.73	130.85	3.	238.17	1514
Corey Ck. 10-09-76	3.35	9.79	-2.	37.96	62
Elk Run Ck. 8-06-78	44.73	109.13	Ο.	118.80	197
Five Mile Ck. 9-25-77	-10.92	-40.37	2.	31.17	529
Newtown Ck. 10-09-76 10-28-81	1.17 38.41	28.91 9.49	-1. 0.	32.62 37.62	450 331
Ostelic River 10-09-76	9.85	-20.67	5.	27.60	1223
Owego Ck. 10-09-76	3.15	36.59	0.	47.24	1903

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